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PRELIMINARY REPORT ON KIKIKTAT MOUNTAIN: A KLIPPE OF BASALTS, CHERTS, AND  
CALCAREOUS ARKOSIC SEDIMENTS IN THE NORTHERN BROOKS RANGE, ALASKA

By

Diana Nelson Solie

Alaska Division of  
Geological and Geophysical Surveys

June 1986

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794 University Avenue, Basement  
Fairbanks, Alaska 99709

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INTRODUCTION

Kikiktat and West Kikiktat Mountains, collectively known as Kikiktat Mountain, on the west side of the Killik River on the north flank of the Brooks Range, are an isolated pair of basalt-capped mountains in the Killik River quadrangle. Interlayered shales and calcareous arkosic sandstone together form an allochthonous klippe. This klippe is the uppermost of the imbricated thrust sheets of the Brooks Range.

The tholeiitic basalts consist of plagioclase, clinopyroxene, opaques, and alteration minerals, and are silica saturated to oversaturated. The interlayered cherts are generally greyish and recrystallized. Correlation of these basalts with ultramafic and mafic ophiolitic complexes elsewhere in the Brooks Range is tempting, but inconclusive based on available data. The underlying sedimentary section contains interbedded shales and sandstones. The sandstones are very calcareous and contain clastic grains of quartz and feldspar. Plant stem fragments, and invertebrate microfossils are common in the sandstones, and suggest a Paleozoic age for the sediments.

REGIONAL SETTING

Of the many thrust sheets that comprise the Brooks Range, the structurally uppermost tectonic package consists of an ophiolitic sequence of ultramafics, cumulate gabbros, basalts and cherts. This sequence is thought to have obducted from the south in Juro-Cretaceous times, over a minimum distance of 120 km (Roeder and Mull, 1978). Two discrete terranes, one of gabbro and dunite, the other of basalt, were stacked in inverse order during obduction (Mull, 1982). The mafic volcanics and gabbros of the Angayucham terrane in the southern Brooks Range, and several ophiolitic klippen (for example, Siniktanneyak Mountain, Misheguk Mountain, and Avan Hills) in the western Brooks Range, as well as the Christian Complex in the eastern Brooks Range, and the Innoko and Tozitna terranes and Rampart Group of the southern Yukon and Koyukuk basins have all been attributed to this obducted thrust package (Dillon, 1985; Roeder and Mull, 1978). On the basis of stratigraphic and structural setting, and composition of basalt and chert, it may be possible to correlate the rocks of Kikiktat Mountain with the basaltic portion of this ophiolitic thrust sheet. However, little data is available with which to make definitive correlations.

FIELD OBSERVATIONS

Of the nearly 2,000 vertical ft of topographic relief of Kikiktat Mountain, approximately 1,500 ft is comprised of basalt with scattered lenses and discontinuous layers of chert. The basalt is dark green, locally magnetic, has brownish-reddish weathered surfaces, is unfoliated and blocky,

<sup>1</sup>DGGS, 794 University Ave., Basement, Fairbanks, Alaska 99709.

with extensive brittle fracturing. In general, deformation has made individual flows indiscernible, though forms suggestive of pillows are not uncommon. A well-exposed but inaccessible layer with columnar jointing is visible on the eastern cliffs of West Kikiktat. There is an overall layered pattern to the basalts, and intervals of chert indicate that the Kikiktat basalts, and intervals of chert indicate that the Kikiktat basalts occurred at least in part as oceanic flows. No gabbros or ultramafics were seen in the unit. The overall sense of layering in the mountain is in the form of a synclinal warp.

One to 2 mm round bluish-greenish milky blebs occur in basalts locally, especially near chert horizons. These are interpreted as alteration resulting from fluids driven from the cherts upon eruption of the basalts. Secondary minerals in fractures and veins are common; these include clear euhedral quartz crystals up to 2 cm long, and pale green prehnite crystals in aggregates 1 to 2 cm thick, and calcite. In several areas, maroon, black, and green shaley horizons occur, with associated turquoise-blue cherty zones. Rare tiny feldspar crystals and lithic fragments resembling lapilli suggest that these are tuffaceous layers between the basalt flows.

The cherts within the basalt unit are typically black to grey, but green-grey, blue-grey, maroon, and turquoise-blue cherts are also present. All tend to have a white-weathered surface. No chert horizon was seen to extend uninterrupted for more than about 30 to 50 meters. This is probably due to deformation within the unit during obduction. Broken nodules and blocks of chert surrounded by basalt may be a result of disruption during emplacement of the basalt flows. Elsewhere, folding of chert lenses and intense shearing in surrounding basalt suggests that the cherts have been structurally broken and deformed after the basalts cooled.

Deformation is predominantly brittle on Kikiktat Mountain. Intense brittle fracture is prevalent in the basalts, with and without secondary mineral infilling. Shear bands within basalts and cherts are thin (.5-2 mm), with irregular crosscutting orientations. Mineral fibers across fractures are bent, but not deformed, indicating growth in a changing stress field. Some fault planes are filled with non-foliated cataclasite and reaction minerals indicating fluid flux along the fault zone.

Below the basalts and cherts of Kikiktat and West Kikiktat Mountains is a unit of shales and calcareous arkosic sandstone, which was measured to be at least 30 meters thick. There is no evidence for faulting along the contact between this unit and the overlying basalt. The uppermost sandstone becomes silicified and grades into a chert at the contact; there is an increase in iron oxide at the contact in both chert and basalt, and a dark homogeneous margin in the basalt is probably an altered chill zone. A small (1 cm) clast of calc-sandstone was seen incorporated within the lower basalt, probably plucked up when the basalt flowed over the sediments. Figure 3 is a measured section of this unit, exposed in 'Dutweiller Canyon' on the western side of Kikiktat Mountain. The sandstones are grey, calcareous, fine- to coarse-grained, with clasts of quartz and feldspar. Shaley rip-ups are common. Upper surfaces of some beds are undulatory; plane and ripple laminations, laminated cross-beds, starved ripples and burrows are all features of the sandstone beds.

The shales are grey, fissile, and phyllitic near the structural bottom of the unit. Some shales are calcareous; sandstone lenses and limestone nodules are common. Abundant plant fragments occur in some sandstone beds. Only plant stems of indeterminate age were found. Invertebrate fossils are abundant in the calcareous sandstone layers.

Deformation of the calc-sandstone unit is characterized by small-scale kink folding in the shales (wavelength about 1 cm), slip along bedding planes - which is perhaps responsible for disruption of thin lensoid sandstones in the shaley layers, and by later brittle normal slip faults. Deformation seems to increase toward the lower part of the unit, and though not exposed, the lower contact of the sandstone-shale unit with underlying basalts is probably a fault. Structurally beneath the sedimentary package are more basalts and cherts. This lower section of basalt appears to be only a few flows thick with well-developed slickensides between many of the layers, suggesting structural repetition. Wherever exposed, the lower contact of the lowermost basalt is faulted, supporting the interpretation of Kikiktat Mountain as a thrust klippe. The klippe overlies Paleozoic sedimentary rocks which are familiar from the local stratigraphy, including interbedded black chert and limestone, and blue-grey chert. On West Kikiktat, dark green conglomerate and sandstone with clasts of black and blue chert occur between basalt layers. This layer is similar in appearance to nearby Cretaceous conglomerates, and is interpreted as a fault sliver within the basalt sheet, though contact relations are not clear.

#### PETROGRAPHY AND ANALYTICAL DATA

##### Basalt

The Kikiktat basalts are modally quite uniform throughout the map area. They consist primarily of plagioclase, clinopyroxene and opaques. Secondary green micaceous alteration minerals (celadonite?) and titanite, some in euhedral forms, suggest pseudomorphing of primary olivine. Other alteration minerals include iron oxide, chlorite, biotite, calcite, sericite, epidote, and opaque. The degree of alteration varies, with the freshest samples having  $\leq 1$  percent secondary minerals, and the least fresh being nearly all alteration minerals. No glass is recognizable in thin section. Results of point counts of six thin sections (400 counts each) are shown in table 1, and plotted in figure 4. Average modal percentages are : 53 percent plagioclase (predominantly labradorite), 37 percent augite, 0.05 percent opaques (including magnetite, ilmenite(?), and pyrrhotite), and 0.05 percent possible olivine pseudomorphs. The basalts are aphanitic to fine-grained, some with phenocrysts of clinopyroxene and/or plagioclase up to about 5 mm. There is no foliation; textures are generally intersertal. Major oxide analyses of nine representative basalts from Kikiktat Mountain and two similar basalts from a knob east of the Killik River are shown in Table 2, with accompanying CIPW norm calculations for each sample. Four of the 11 samples are quartz normative, the other seven are olivine normative; none are nepheline normative. According to Yoder and Tilley's (1962) classification scheme, the samples are tholeiites and olivine tholeiites. A variation diagram plotting  $\text{Na}_2\text{O} + \text{K}_2\text{O}$  versus  $\text{SiO}_2$  shows a spread in alkalis over a small range in  $\text{SiO}_2$ . Using the Macdonald and Katsura (1964) classification the samples span the

tholeiite to alkali basalt fields. On a triangular AFM diagram (fig 6), the samples show a high relative proportion of iron, typical of the tholeiitic suite. Microprobe analyses of clinopyroxenes shows them to be typical augites. When pyroxene compositions are plotted on the discriminant diagrams of Nisbet and Pearce (1977), they fall within both the ocean floor basalt field and the fields representing pyroxenes from all tectonic settings, making tectonic interpretation based on pyroxene composition ambiguous. The possibility of alkali mobility in these basalts cannot be discounted due to the presence of at least small amounts of alteration in all of the basalt thin sections.

#### Calc-arkosic Sandstones

Thin sections of sandstone layers from the unit directly beneath the basalts reveal clasts of quartz, alkali feldspar, plagioclase, opaque, up to 45 percent modal carbonate, and trace amounts of zircon, ±biotite, ±white mica, ±chlorite, ±titanite, and hematite. Small altered volcanic rock fragments occur in some layers. Alkali feldspar grains include microcline and perthite, subrounded, up to 3 mm diameter. Quartz is highly strained, with extensive subgrain development and recrystallization of some grains. Plagioclase is present in variable amounts, and tends to be smaller than the alkali feldspar grains. There is no metamorphic foliation evident in the samples. Most sandy layers contain invertebrate fossil fragments. These include foraminifera, bryozoa, brachiopods, trilobites, ostracods, and pelmatozoan debris. Micritic filling within bryozoa fragments indicates that at least some of these fossils have been reworked.

#### DISCUSSION

The fossil assemblage present in the calc-arkosic sandstone unit strongly suggests a late Paleozoic age for these sediments. This is permissive evidence, though obviously not definitive, for correlation with the Carboniferous Nuka Formation further west in the Brooks Range. The similar-looking Nuka Formation, at its type locality at Nuka Ridge, consists of quartz, alkali feldspar, calcite, opaques, ±biotite, ±hornblende, ±zircon, ±fossils. Source rocks for the clastic grains of the Nuka Formation and the Kikiktat sediments are granites; the Kikiktat source was apparently a subsolvus granite, probably containing biotite ±muscovite. The type Nuka rocks appear to be derived from a hypersolvus granite (no calcic plagioclase), containing biotite ±hornblende. Microcline from the latter rocks has been dated by K-Ar at 1.2 - 1.8 Ba. (Tailleur, 1985).

The Kikiktat sediments were deposited in a Paleozoic marine setting, probably continental shelf, perhaps continental slope. The overlying basalts and cherts were also formed in a marine environment, although the occurrence of columnar jointing suggests local subaerial cooling. The age of crystallization of the Kikiktat basalts remains uncertain. Cretaceous K-Ar ages of a similar allochthonous mafic complex to the west have been reported by Boak and others (1985). Similar ages for corresponding metamorphic contact aureoles lead the authors to the interpretation that the K-Ar ages of the Misheguk sequence gabbros and diorites represent resetting due to post-igneous thermal metamorphism. It seems clear at Kikiktat that the

basalts flowed over the arkosic sandstone and shales---the existence and duration of any time hiatus is unknown.

The allochthonous nature of the sandstone±shale and basalt±chert package on Kikiktat Mountain is indicated by faulted contacts of the lowermost layers. The overall structural style is brittle deformation, with locally extensive shear zones. The oceanic character of the basalts is evident by the interlayering with chert. Whether these rocks are part of an ophiolite sequence is not definitive, based on the limited assemblage present, but they could conceivably be part of a dismembered ophiolite. Trace element data are forthcoming with which to further characterize the Kikiktat Mountain basalts. Correlation with other mafic bodies in the Brooks Range must await more detailed structural, petrologic and chemical data of the various occurrences.

#### ACKNOWLEDGMENTS

Field work for this study was done during June - July 1985, as part of the Alaska Division of Geological and Geophysical Surveys' regional studies in the Brooks Range. Gil Mull was the chief geologist, and I thank him for the opportunity to work on the Kikiktat Mountain rocks. Jerry Siok was of great help in measuring the sedimentary section, and I thank him also for information about the type Nuka Formation. I also want to acknowledge Fred Read, Ed Simpson, Isabel Montañez and Julie Herman at Virginia Polytechnic Institute for help in identifying microfossils, and Todd Solberg, also at VPI, for help with the microprobe analyses. The ADGGS geochemical laboratory did major oxide analyses of the Kikiktat basalts.

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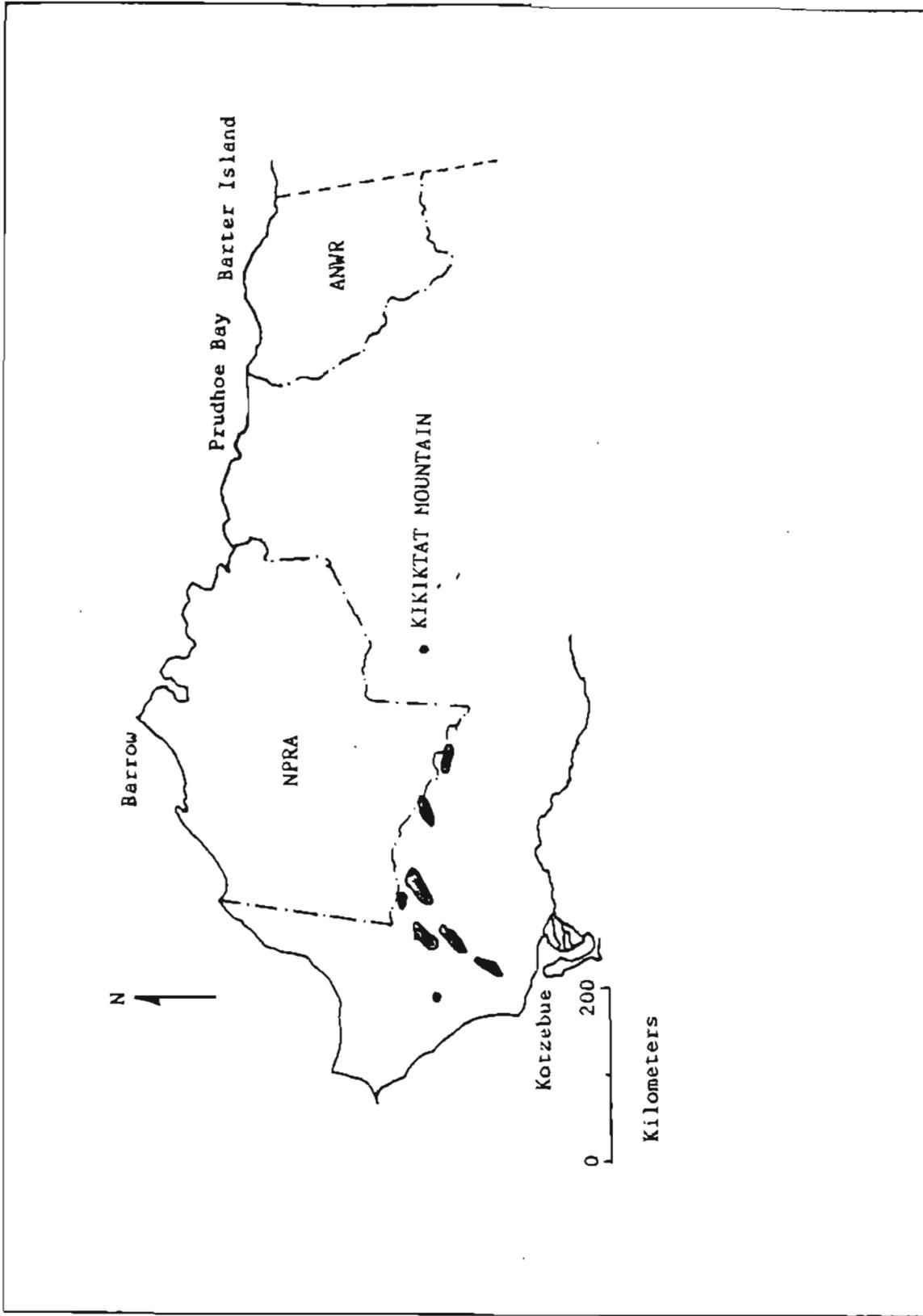
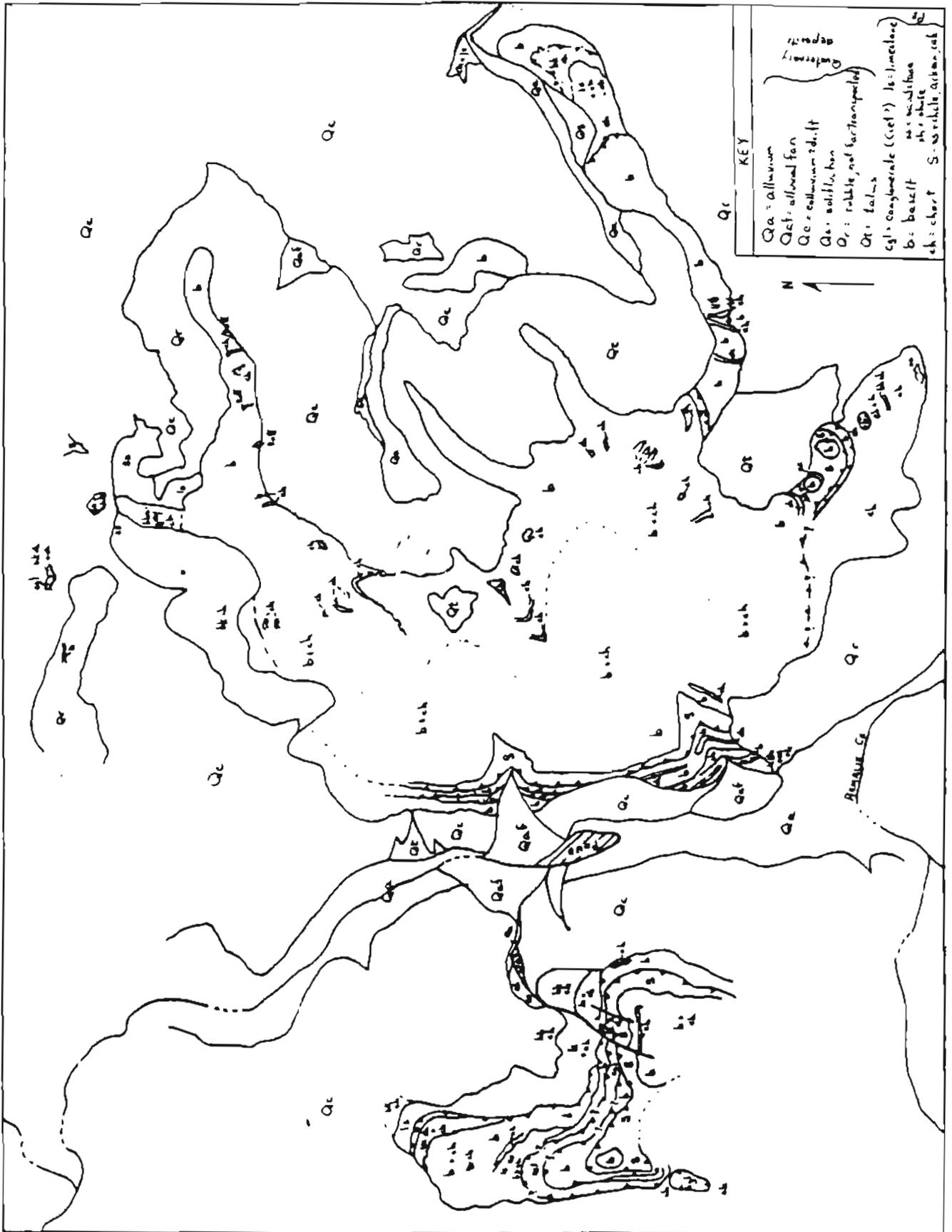


Figure 1. Location map showing Kikiktat Mountain. Shaded areas are other Brooks Range ophiolitic bodies.



KEY

Qa	: alluvium
Qc	: alluvial fan
Qe	: alluvium, d.f.t
Qg	: siltstone
Qj	: rubble, not transported
Qk	: talus
cg	: conglomerate (Cret) ls-limestone
b	: basalt
sh	: shale
sk	: chert
S	: ss-siltstone, siltstone

Figure 2. Map of Kikikrat Mountain area.

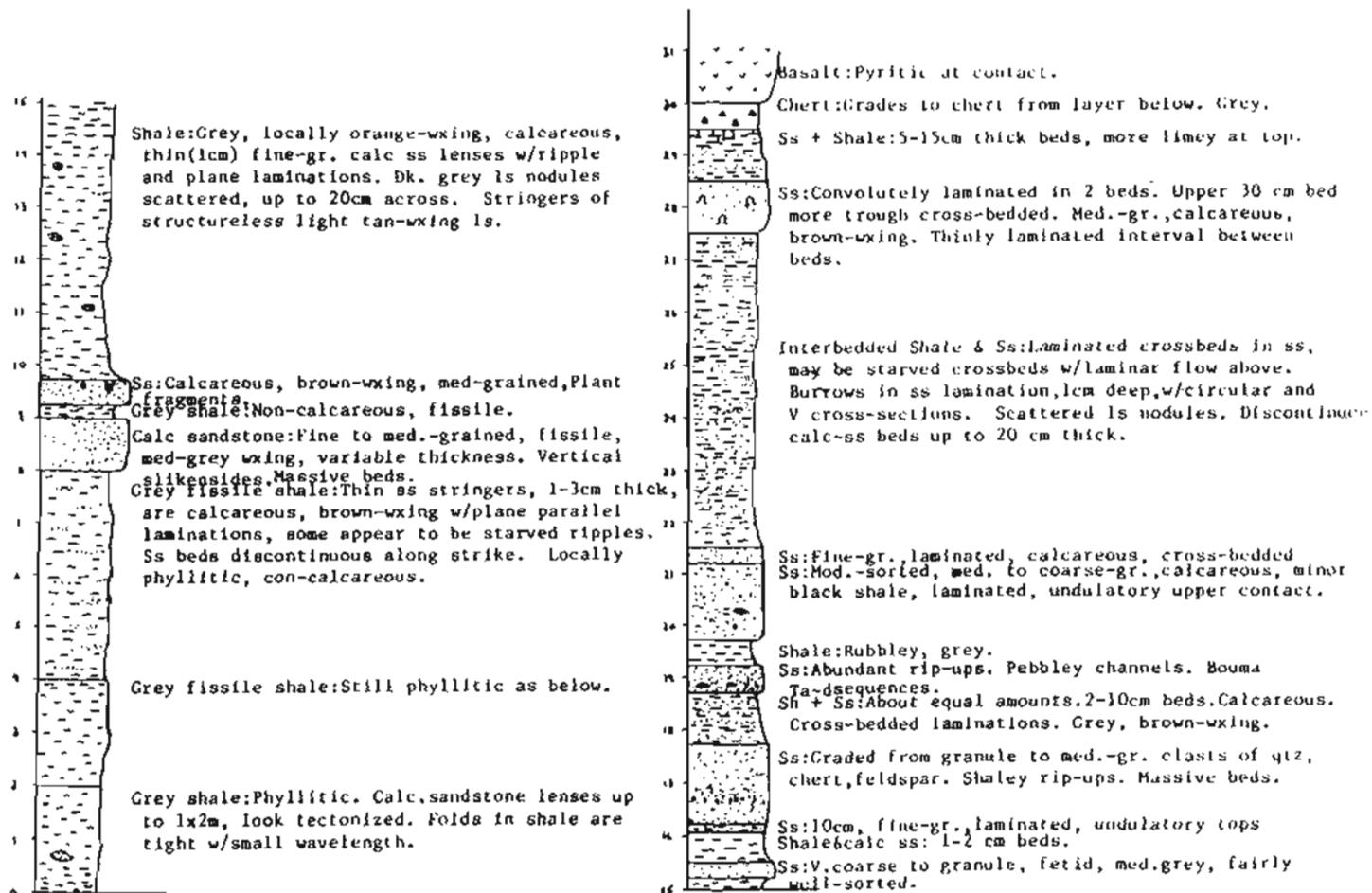


Figure 3. Measured section of sedimentary section underlying Kikiktat Mountain basalts.

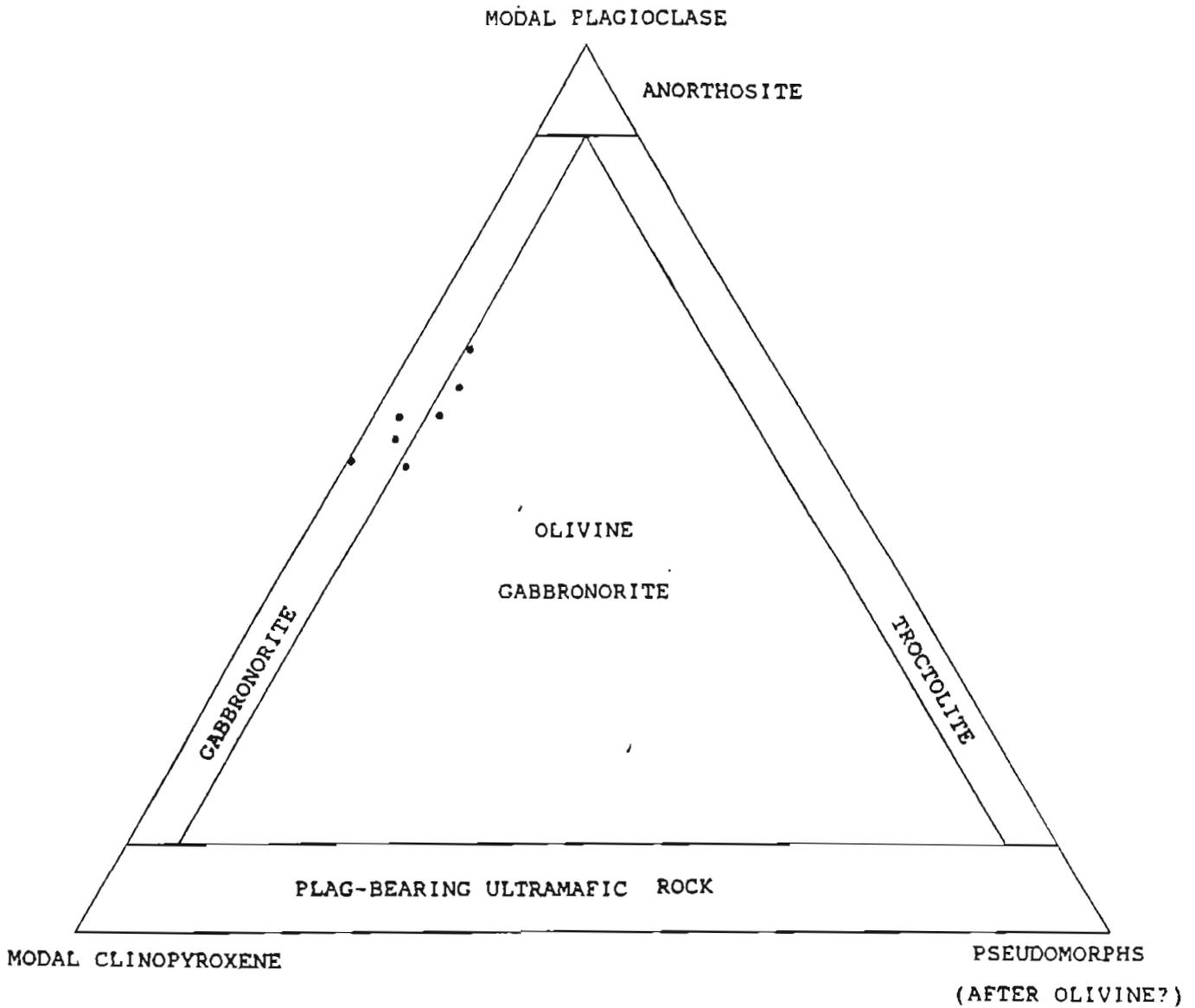


Figure 4. Basalt modal point count data, plotted in terms of plutonic classification scheme (Streckeisen, 1973).

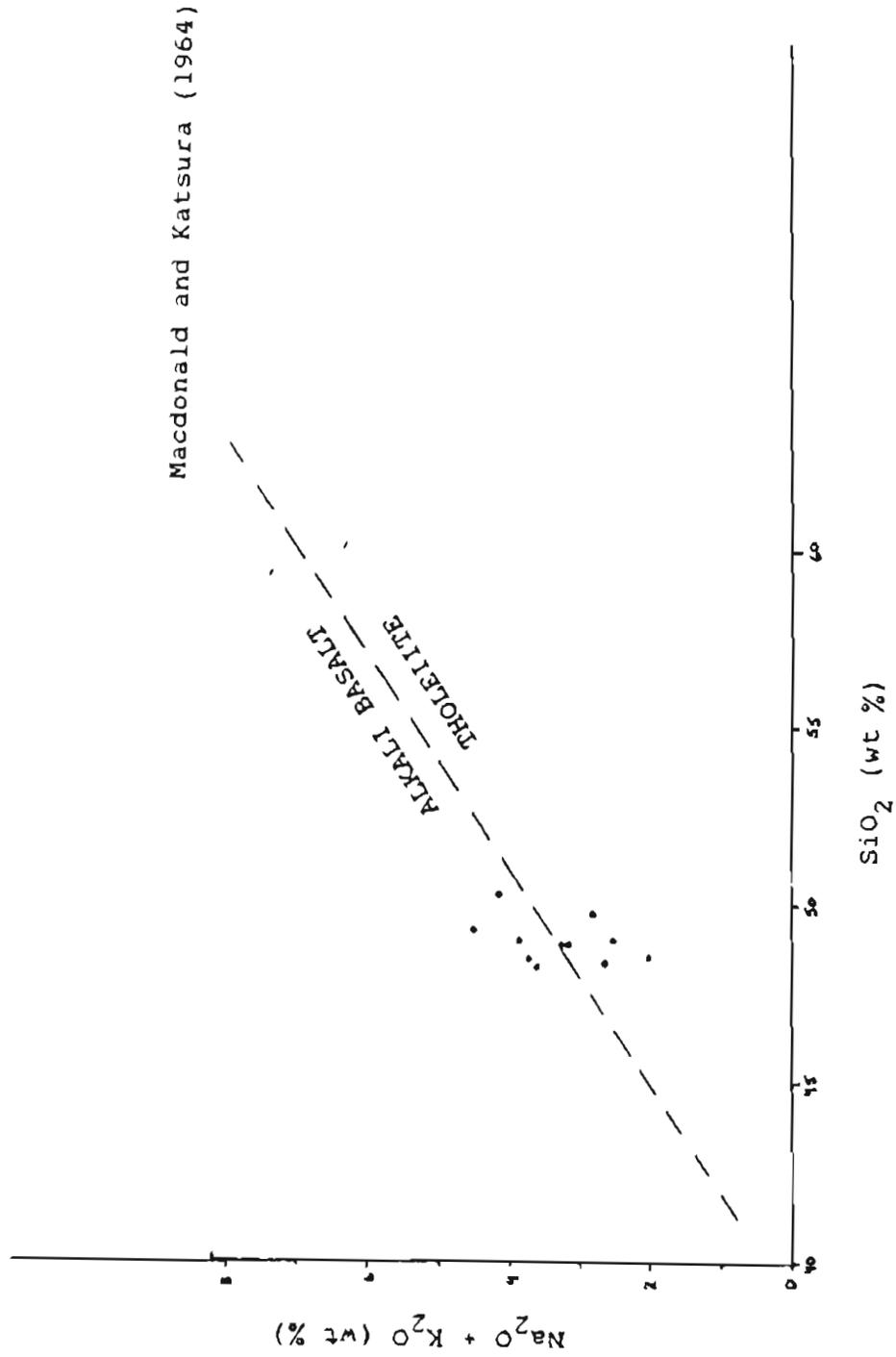


Figure 5. Variation diagram of Kikikat basalts, plotting Na<sub>2</sub>O+K<sub>2</sub>O vs. SiO<sub>2</sub>.

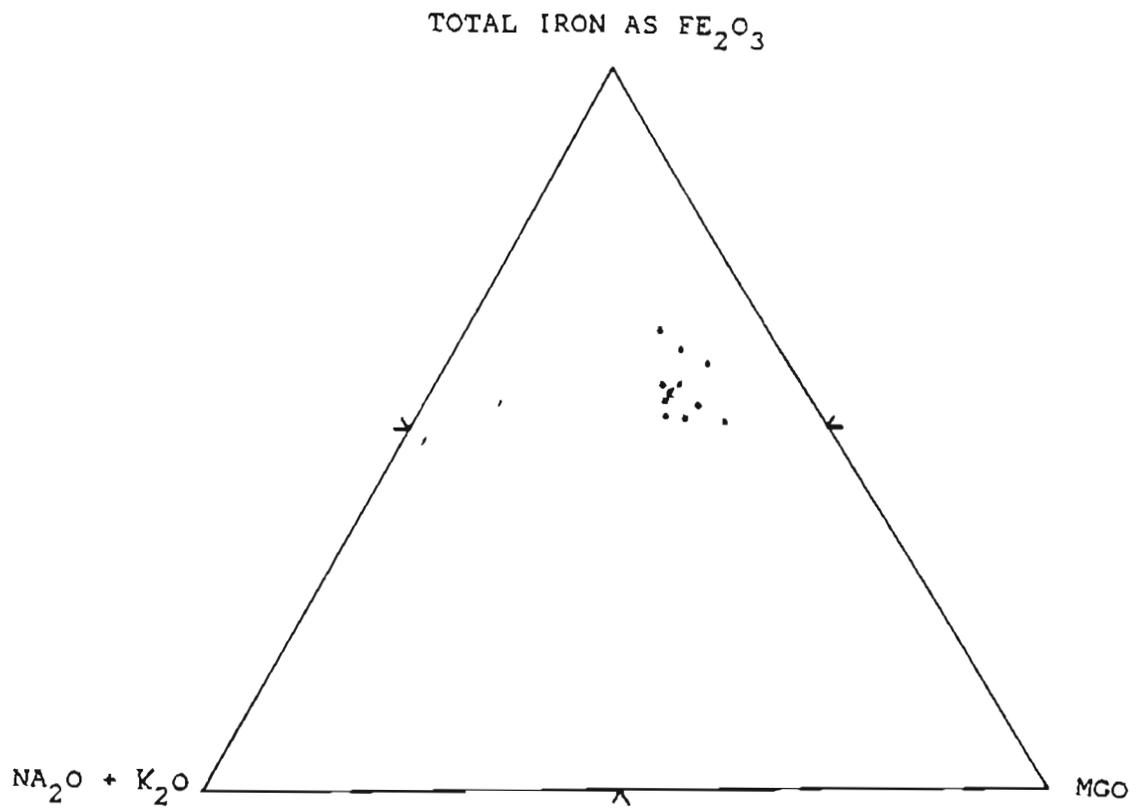


Figure 6. AFM diagram of Kikiktat basalts.

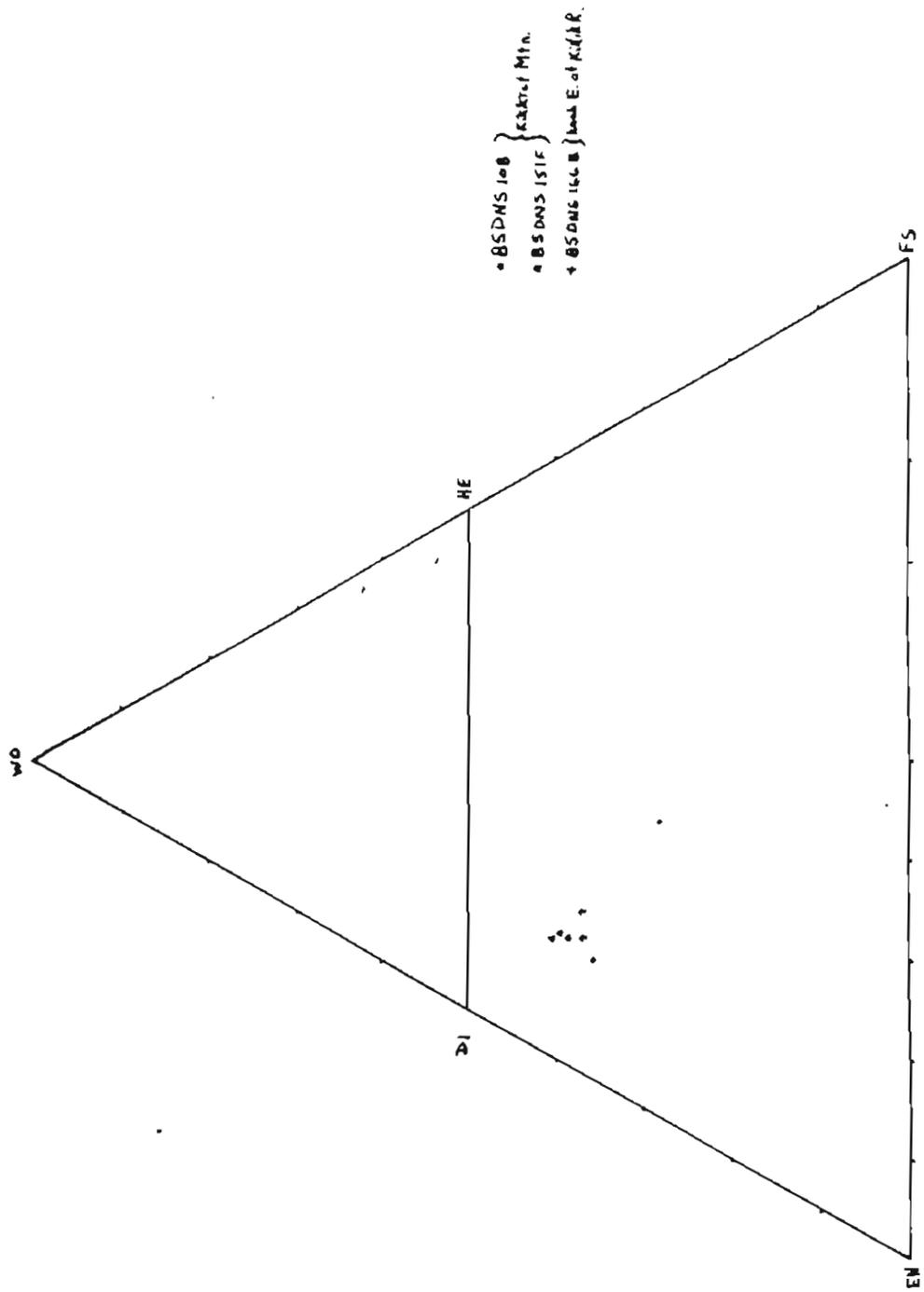


Figure 7. Pyroxene quadrilateral---analyses from Kikikat Mountain and Killik River basalts.

Table 1. Point counts of 5 Kikiktat basalts and 1 basalt (#166) from east of Killik River. 400 points counted per sample.

Sample#	POINT COUNTS--KIKIKTAT BASALTS					
	151	108	166	142	119	161
PLAG.	47.25	58.5	49.25	54.25	56.25	52
CPX.	41.75	30.5	38.75	36.75	34	38
OPAQUE	6	4.75	6.5	6.5	3.75	7
OLIVINE?	3	6.25	5.5	2.5	8	3
TOTAL %	100	100	100	100	100	100

Table 2. Major oxide analyses and corresponding CIPW norm calculations for 9 Kikiktat basalts and 2 basalts from east of Killik River (166A and 166B). Total iron was reported as Fe<sub>2</sub>O<sub>3</sub>, proportioned according to Irvine and Baragar, 1971 such that Fe<sub>2</sub>O<sub>3</sub>=TiO<sub>2</sub> + 1.5, with rest going to FeO.

	KIKIKTAT										
	105A	107A	119E	122B	123	126D	142	151F	161	166A	166B
SiO <sub>2</sub>	48.38	48.32	48.45	49.77	48.90	48.97	48.96	50.29	48.52	49.32	49.02
Al <sub>2</sub> O <sub>3</sub>	14.84	14.32	14.15	12.83	13.26	14.48	14.49	14.74	13.70	14.33	13.76
Fe <sub>2</sub> O <sub>3</sub>	3.47	3.24	2.94	3.77	3.11	3.12	3.31	2.96	3.12	3.18	3.45
FeO	0.14	0.28	0.62	10.13	0.26	7.23	8.66	8.15	8.51	9.54	9.02
MgO	6.08	6.81	7.80	5.36	7.74	8.55	6.71	6.93	6.62	7.11	6.97
CaO	10.95	9.16	9.29	10.19	10.84	10.69	10.29	9.15	11.11	7.17	9.72
Na <sub>2</sub> O	2.47	3.26	3.61	2.52	3.01	2.37	3.04	3.55	1.97	4.36	3.53
K <sub>2</sub> O	0.19	0.38	0.14	0.24	0.25	0.15	0.19	0.67	0.86	0.17	0.33
TiO <sub>2</sub>	1.97	1.74	1.44	2.27	1.61	1.62	1.81	1.46	1.62	1.68	1.95
P <sub>2</sub> O <sub>5</sub>	0.18	0.16	0.12	0.21	0.15	0.13	0.15	0.15	0.15	0.14	0.17
MnO	0.23	0.23	0.20	0.24	0.20	0.16	0.22	0.25	0.20	0.23	0.24
TOTAL	97.88	97.50	96.74	97.55	97.31	97.27	97.63	98.24	97.56	97.23	98.06
Q	1.518	2.308	0.855	5.197	1.518	0.992	1.150	6.010	2.156	1.033	1.989
OR	1.167	1.454	1.574	1.454	26.174	0.917	0.917	0.150	0.363	0.363	0.461
AB	21.353	21.866	22.876	21.866	22.538	28.617	26.348	38.685	17.086	37.944	30.461
AN	29.469	23.965	22.876	23.569	15.300	29.226	23.945	22.786	29.071	19.570	21.135
MO	18.325	10.323	10.885	11.214	15.300	18.197	18.580	9.414	15.334	6.711	11.235
EH	15.478	11.064	9.522	13.687	15.013	21.379	16.207	12.013	16.900	7.949	11.378
FS	11.332	8.018	5.585	12.494	8.031	8.555	9.912	7.335	11.015	5.631	7.275
FO	4.103	4.103	7.399	3.361	3.361	8.555	8.555	3.893	11.015	7.192	4.254
FA	3.274	4.783	4.783	1.981	1.981	4.651	0.638	2.620	4.637	5.615	2.998
MY	5.140	4.828	4.406	5.605	4.634	4.651	4.916	4.349	4.637	4.742	5.101
IL	3.822	3.396	2.827	4.420	3.142	3.163	3.521	2.823	3.154	3.282	3.777
AP	0.436	0.389	0.294	0.518	0.316	0.317	0.364	0.311	0.316	0.341	0.411
TOTAL	100.012	100.011	100.009	100.014	100.009	100.009	100.011	100.010	100.009	100.010	100.012
SALIC	53.487	54.623	55.108	52.804	50.230	51.747	53.443	57.232	48.655	56.567	53.585
FEMIC	46.526	45.388	44.901	47.210	49.780	48.262	46.568	42.778	51.354	43.443	46.427
D.I.	24.018	30.658	32.431	28.515	27.692	22.521	27.498	34.435	19.584	38.977	32.449

Table 3. Pyroxene analyses, Kikiktat Mountain basalts, Killik River Quadrangle, Alaska. Electron microprobe, Va. Tech, analyzed by D.N. Solie.

PYROXENE ANALYSES									
	BSM5100								
SiO2	51.23	53.22	52.15	53.37	50.19	51.67	52.12		
TiO2	0.98	0.49	0.71	0.59	0.96	0.64	0.53		
Al2O3	1.75	2.04	3.55	2.34	3.68	3.47	2.52		
FeO	18.03	7.61	0.00	8.66	9.68	7.72	7.25		
MnO	0.48	0.20	0.24	0.22	0.23	0.22	0.21		
MgO	14.72	18.90	17.23	17.76	15.96	16.54	16.67		
CaO	13.75	18.00	19.13	18.30	17.53	19.34	20.04		
NaO	0.09	0.04	0.04	0.05	0.09	0.0	0.04		
Na2O	0.23	0.15	0.22	0.23	0.26	0.24	0.22		
K2O	0.02	0.01	0.01	0.01	0.01	0.01	0.01		
F	0.01	0.02	0.05	0.07	0.04	0.0	0.06		
CL	0.0	0.0	0.0	0.0	0.0	0.0	0.0		
SUM	101.28	100.66	101.08	101.53	98.59	99.85	99.59		
-O=F+CL	0.0	0.0	0.0	0.0	0.0	0.0	0.0		
SUM	101.28	100.66	101.08	101.53	98.59	99.85	99.59		
Si	1.927	x	1.900	x	1.933	x	1.926	x	
Al	0.073	2.000	0.064	2.000	0.100	2.000	0.074	2.000	
Al	0.005	x	0.023	x	0.044	x	0.056	x	
Ti	0.020	x	0.013	x	0.019	x	0.018	x	
Fe	0.567	x	0.231	x	0.262	x	0.258	x	
Mn	0.015	x	0.006	x	0.007	x	0.007	x	
Mg	0.825	1.441	1.025	1.299	0.936	1.250	0.909	1.228	
Ca	0.554	x	0.701	x	0.747	x	0.764	x	
Na	0.001	x	0.001	x	0.001	x	0.0	x	
Na	0.017	x	0.011	x	0.016	x	0.017	x	
K	0.001	0.573	0.000	0.713	0.000	0.763	0.000	0.782	
O	6.000	x	6.000	x	6.000	x	6.000	x	
H2O	28.25		35.72		38.62		36.93		40.85
EN	42.07		52.18		48.39		46.77		47.27
FS	29.69		12.10		12.99		13.88		11.87
F/M	0.706		0.232		0.268		0.281		0.253
F/FM	0.414		0.188		0.212		0.258		0.201

Table 4. Plagioclase analyses, Kikiktat Mountain basalts, Killik River Quad-  
 range, Alaska. Electron microprobe analyses, analysed by D.N. Solie.

		FELDSPAR ANALYSES									
		85DMS10g		85DMS166b		85DMS166b		85DMS166b		85DMS166b	
SiO2		50.71		62.83		68.44		53.36		52.80	
TiO2		0.07		0.08		0.01		0.11		0.13	
Al2O3		31.67		24.34		21.12		29.47		29.24	
FeO		0.74		0.53		0.10		1.04		1.03	
MnO		-0.01		-0.01		-0.04		0.0		0.01	
MgO		0.23		0.06		-0.02		0.18		0.19	
CaO		14.82		5.99		1.24		13.04		12.91	
BAO		0.03		-0.01		0.0		-0.03		0.0	
NA2O		3.40		8.58		11.03		4.24		4.18	
K2O		0.03		0.06		0.03		0.06		0.06	
SUM		101.69		102.45		101.91		101.47		100.55	
SI		2.283	*	2.730	*	2.939	*	2.395	*	2.392	*
AL		1.680	*	1.246	*	1.069	*	1.559	*	1.561	*
TI		0.002	3.966	0.003	3.979	0.000	4.008	0.004	3.957	0.004	3.958
FE		0.028	*	0.019	*	0.004	*	0.039	*	0.039	*
MN		-0.000	*	-0.000	*	-0.001	*	0.0	*	0.000	*
MG		0.015	*	0.004	*	-0.001	*	0.012	*	0.013	*
CA		0.715	*	0.279	*	0.057	*	0.627	*	0.627	*
BA		0.001	*	-0.000	*	0.0	*	-0.001	*	0.0	*
NA		0.297	*	0.723	*	0.918	*	0.369	*	0.367	*
K		0.002	1.057	0.003	1.028	0.002	0.978	0.003	1.050	0.003	1.050
O		8.000	*	8.000	*	8.000	*	8.000	*	8.000	*
AlI		70.51		27.75		5.84		62.77		62.84	
AB		29.27		71.93		93.99		36.94		36.82	
OR		0.17		0.33		0.17		0.34		0.35	
CIH		0.05		-0.02		0.0		-0.05		0.0	
F/M		1.780		4.861		0.0		3.242		3.072	
F/PM		0.640		0.829		0.0		0.764		0.754	