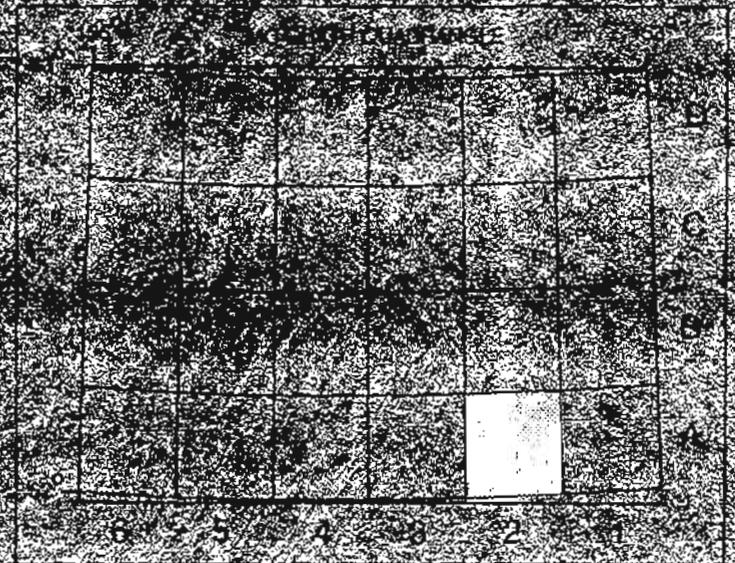
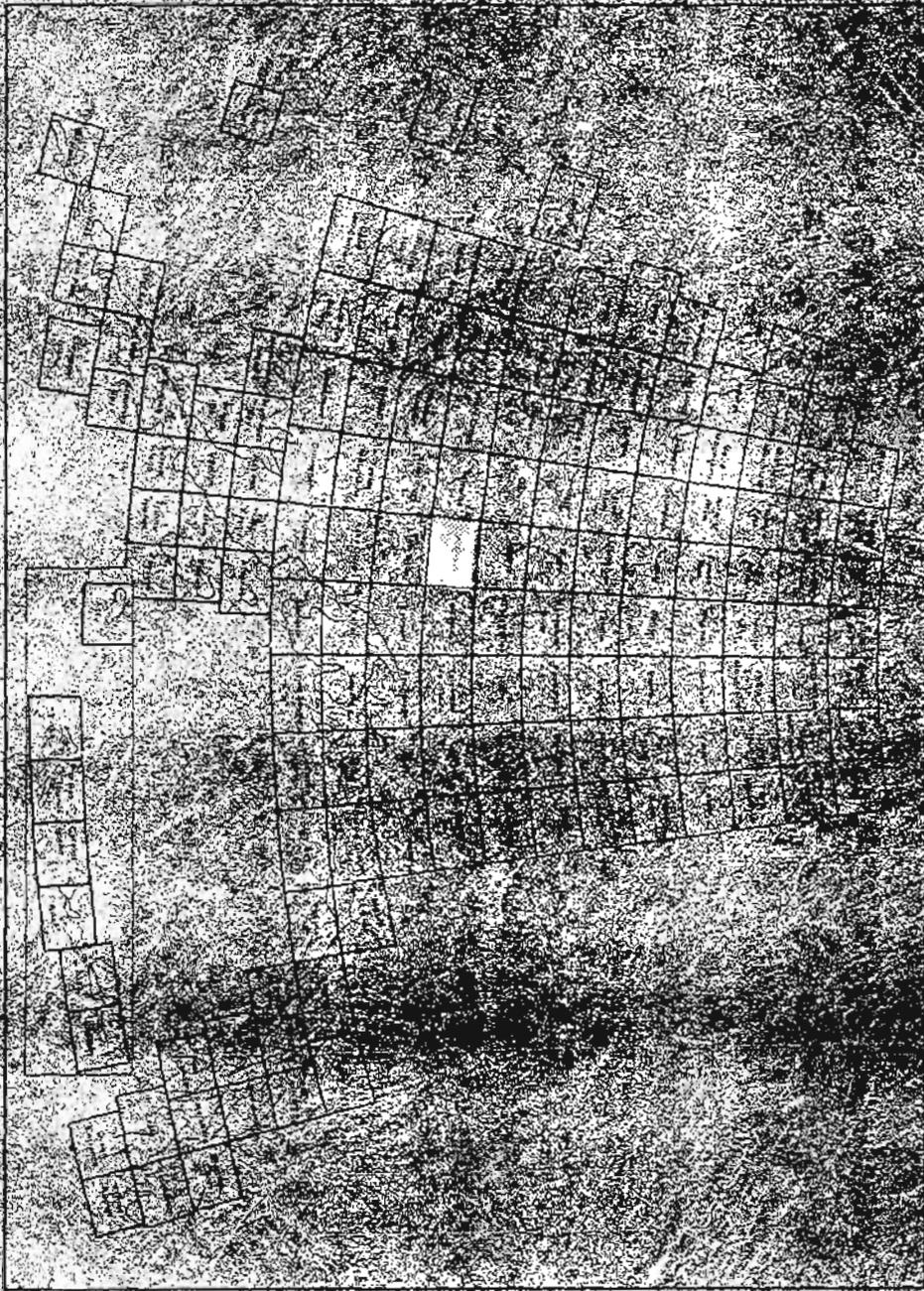


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GEOLOGY OF THE MC GRATH A-2  
QUADRANGLE, ALASKA



# GEOLOGY OF THE MCGRATH A-2 QUADRANGLE, ALASKA

By T.K. Bundtzen, J.T.Kline, T.E. Smith,  
and M.D. Albanese

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## PROFESSIONAL REPORT 91

*A summary of the stratigraphy, lithology, and  
mineral resources of a structurally deformed  
450 km<sup>2</sup> (175 mi<sup>2</sup>) region on the north flank  
of the western Alaska Range, Alaska*





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## INTRODUCTION

The McGrath A-2 Quadrangle is composed of rugged mountains with numerous glacier-scoured, U-shaped valleys on the northwest flank of the southern Alaska Range. Elevations range from 1,750 ft (534 m) at Post River to 7,880 ft (2,402 m) in the west-central part of the quadrangle. No roads exist in the area; however, a 1,700-m-long gravel airstrip is maintained at the Farewell Federal Aviation Administration station 35 km to the north. Fixed-wing airplanes can also land on an unmaintained airstrip at Bowser Creek, on gravel flood plains, and on Post Lake.

Sedimentary rocks of the McGrath area form irregular, rolling hillslopes; igneous plutons, dike swarms, and thermally altered rocks form prominent, rugged, circular massifs. Valley floors and hillslopes below 4,000-ft (1,220 m) elevation are mantled with surficial deposits and vegetation. Colluvium covers many hillslopes, and bedrock exposures are concentrated on ridge tops and cirque headwalls and in steep canyons. Bedrock and surficial control is generally excellent. Although the metric system is used in this report, elevations are also expressed in feet to be consistent with the published topographic maps.

## ACKNOWLEDGMENTS

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## BEDROCK GEOLOGY

Sixteen mappable rock units that range in age from Cambrian(?) to Cretaceous are exposed in the McGrath A-2 Quadrangle. Figure 1 (map inset) shows a simplified version of the bedrock geology. Most layered units were originally described by Brooks (1911) as the Tatina Group after exposures on the Tatina River in the McGrath B-1 Quadrangle 12 km northeast of Post Lake. Armstrong and others (1977) later referred to these rocks as the 'sedimentary rocks of the Dillinger River.' Jones and others (1982) assigned the term 'Dillinger terrane' to those deepwater rocks of lower to mid-Paleozoic age in the study area; Bundtzen and Gilbert (1983) used the term 'Dillinger group' for

these same rock units.<sup>1</sup> Gilbert and Bundtzen (1984) referred to most pre-Jurassic layered rocks south of the Farewell fault in the southern Alaska Range as the 'Dillinger-Mystic succession.' The Dillinger terrane consists of Cambrian and Ordovician limestone turbidite deposits (OEs) with minor greenstone sills (OEG); Ordovician and Lower Silurian interbedded shale, chert, and minor limestone (SOsh, ISI); Middle to Upper Silurian laminated limestone (mSl, uSl, uSsl), sandstone turbidite (Ss, mSs), and shale (mSa); and Upper Silurian to Lower Devonian limestone (DSl) and limestone breccia and calcareous sandstone (DSls). Bundtzen and Gilbert (1983) suggested that the Dillinger terrane reflects a shallowing-upwards marine regime that includes basinal, turbidite-fan, and foreslope depositional environments on a Paleozoic continental margin. Limited paleocurrent data (table 1) and probable facies relationships with the Nixon Fork terrane (Bundtzen and Gilbert, 1983) indicate that during the Paleozoic Era, a shoreline existed northwest of the study area. Bearings of the paleocurrents are dominantly southeasterly in lower Paleozoic units, but are dominantly northeasterly in the Silurian turbidite fan. The latter bearing was probably derived from a paleocurrent that was directed down the axes of a sedimentary trough parallel to the continental margin. The Dillinger group ranges in thickness from 2,750 to 3,000 m in the study area; the base is not exposed.

The Dillinger terrane is overlain by a 1,000-m-thick section of sublithic quartz sandstone, limestone, and fluvial(?) conglomerate (IPDs, IPDI, IPDe) of probable Middle Devonian through Late Pennsylvanian age. Gilbert and Bundtzen (1984) suggested that these units form a southwestern extension of the Mystic terrane as described by Jones and others (1982) in the Talkeetna Quadrangle. Gilbert and Bundtzen (1984) also suggested that the Mystic terrane is a continuation of the sedimentary basin that formed in early Paleozoic time (Dillinger terrane).

A deformed package of lithic sandstone, siltstone, and conglomerate (KJs) of Late Jurassic to Early Cretaceous age is in tectonic contact with Paleozoic rocks. The deformed unit is part of a 50,000-km<sup>2</sup> turbidite-dominated basin that trends from southcentral Alaska near Cantwell (Smith, 1981, p. 15-17) 300 km west-southwest to Lake Clark (Eakins and others, 1978). Jones and others (1982) referred to this flysch sequence as the Kahiltna terrane after exposures in the Talkeetna Quadrangle to the east.

Seven units of hypabyssal rocks crop out in the study area (Tim, Tif, Tids, Tqm, Tqp, Tia, Td). They include the Bowser and Post River intrusive complexes and an extensive east-west-trending dike swarm in the westcentral

<sup>1</sup>For descriptive convenience only, the following text uses the term 'Dillinger terrane.'

Table 1. Paleocurrent data<sup>a</sup> from Paleozoic sedimentary rocks in the McGrath A-2 Quadrangle, Alaska.

Map no.	Field no.	Rock unit	Flow regime <sup>b</sup>	Azimuth in ° (corrected for tilt)	Grand mean
1	82BF233	DSls	upper	25	25.0
2	82BF166	SOsh	upper	24	
				355	08.0
3	82BF163	mSs	upper	352	
				24	26.5
4	82BF188	mSs	upper	29	
				183	192.5
5	82BF191	mSs	upper	202	
				322	321.5
6	83BF300	uSsl	upper	321	
				25	36.2
				40	
				50	
				26	
				40	
7	83BF304	Ss	upper	25	28.5
				32	
8	83BF303	mSs	upper	45	45.0
9	83BF311	mSs	upper	25	
				15	20.0
10	83BF311	uSsl	upper	0	
				25	
				40	21.7
11	83BF148	OEs	lower	120	
				130	
				119	133.0
12	83BF147	OEs	lower	125	
				118	
				109	117.3
13	82BF264	OEs	lower	225	
				250	
				235	237.5
				240	
14	82BF281	IPDe	lower	220	
				210	215.0
15	83BF123	mSs	upper	332	
				342	
				355	313.0

<sup>a</sup>Measurements taken and interpreted by T.K. Bundtzen. Churkin and others (1977) also present paleocurrent data from study area.

<sup>b</sup>Upper flow-regime structures include flutes, striation casts, and longitudinal ripples. Lower flow-regime structures include cross-beds.

part of the map area. The Bowser complex consists of igneous breccia with a quartz monzonite core and quartz porphyry intrusive rocks that cut the igneous lithologies. The Post River pluton is a porphyritic quartz monzonite. Both plutons have sharp high-angle contacts with host rock units.

Prominent dike swarms and scattered dikes of augite basalt, andesite, and rhyolite cut layered rocks in the area; basaltic and andesitic dikes are most abundant. Rhyolite dikes consistently cut the more mafic varieties.

Two K-Ar analyses from intrusive rocks in the study area (table 2) yield early Tertiary ages consistent with those reported by Reed and Lanphere (1972). However, some dikes in adjacent map areas yield ages as young as 20 m.y. (Bundtzen and others, in preparation). Solie and others (1982) believed that all igneous rocks are part of at least two calc-alkaline subduction-related suites active during Tertiary time (table 3).

Table 2. Analytical data for  $^{40}\text{K}$ - $^{40}\text{Ar}$  age determinations.<sup>a</sup>

Sample	1 <sup>b</sup>	2 <sup>c</sup>
Rock type	Quartz porphyry	Quartz monzonite
Mineral dated	Biotite	Biotite
K <sub>2</sub> O (wt %)	5.287	5.412
Sample wt (g)	0.1600	0.1277
$^{40}\text{Ar}$ (rad) moles/g) $\times 10^{-11}$	46.76	49.00
$\frac{^{40}\text{Ar}(\text{rad})}{^{40}\text{K} \times 10^{-3}}$	3.57	3.65
$\frac{^{40}\text{Ar}(\text{rad})}{^{40}\text{Ar}(\text{total})}$	0.534	0.477
Age (m.y.)	60.4 $\pm$ 1.8	61.8 $\pm$ 1.9

<sup>a</sup> Constants used:  $\lambda = 0.581 \times 10^{-10} \text{ yr}^{-1}$ ,  $\lambda = 4.962 \times 10^{-10} \text{ yr}^{-1}$ .

<sup>b</sup> Analyzed by Daniel Krummenacher, San Diego State University, San Diego, California.

<sup>c</sup> Analyzed by J.D. Blum, DGGG-University of Alaska Cooperative Geochronology Laboratory, Fairbanks, Alaska.

## QUATERNARY GEOLOGY

Eleven Quaternary units were mapped in the McGrath A-2 Quadrangle; these units include Wisconsin and Holocene drift, ice-contact, outwash, and till deposits and Holocene colluvium and alluvium. Multistoried cross-valley profiles, perched side-glacial notches, faceted spurs, high level erratic boulders, and multiple lateral and end moraines demonstrate that the study area underwent several major episodes of glaciation. A more complete glacial chronology of the general Farewell area is summarized in Kline and Bundtzen (1986). Before late Wisconsin time, most of the McGrath A-2 Quadrangle was overlain by an extensive ice field with only a few nunataks protruding above ice. Two major glacial episodes of Wisconsin age are primarily responsible for the current morphology and unconsolidated surficial deposits.

Evidence for pre-Wisconsin glaciation in the McGrath A-2 Quadrangle consists of scattered erratics that are found at elevations above 4,500 ft (1,372 m) and relict valley morphology; however, most older deposits have been eroded from the rugged terrain. The oldest extensive glacial deposits in the McGrath A-2 Quadrangle (Odt in high-level terraces) are inferred to be roughly correlative with early Wisconsin deposits of Knik age in the upper Cook Inlet basin (Karlstrom, 1964) and Talkeetna Quadrangle (Nelson and Reed, 1978) based on morphology, degree of weathering, and extent. These glacial deposits may also correlate with deposits of Selatna age along the Big River to the northwest and of Farewell I age along the South Fork Kuskokwim River (originally described by Fernald, 1960). The apparent discrepancy in the latter correlation with previously mapped drift units is due to a recent reevaluation of Fernald's earlier valley-to-valley correlations of drift limits (Kline and Bundtzen, 1986).

In the McGrath A-2 Quadrangle, late Wisconsin drift is roughly correlative with Naptowne-age deposits of the upper Cook Inlet basin (Karlstrom, 1964) and Talkeetna Quadrangle (Nelson and Reed, 1978) and with Farewell II age drift (Fernald, 1960) that occurs along the mountain front in adjacent quadrangles.

Multiple recessional moraines that occur in major southerly tributaries to the Post River resemble the latest Wisconsin and early Holocene deposits of the upper Cook Inlet basin and Kenai Peninsula that Karlstrom (1964) referred to as Skilak and Tanya readvances of late Naptowne time. Late Holocene moraines that extend a few kilometers down-valley from cirque headwalls throughout the study area were deposited well after the main trunk-glacier systems collapsed. Nested neoglacial moraines indicate at least two, and possibly more, late Holocene pulses of glacier activity.

Modern alluvial units (Qa, Qat, Qaf) include terrace, fan, and modern-stream deposits. Colluvial processes and vigorous stream erosion caused by

Table 3. Major-oxide analyses and CIPW norms of igneous rocks.

Map no. Field no. Rock type	1 82BT184 Olivine gabbro	2 82BT165 Gabbro dike	3 82BT183 Rhyolite	4 82BT300 Andesite dike	5 82BT323 Dacite	6 82BT275 Quartz monzonite	7 82BT181 Rhyolite
SiO <sub>2</sub>	45.53	52.02	77.31	56.01	67.00	68.98	77.58
Al <sub>2</sub> O <sub>3</sub>	14.06	17.04	12.84	15.18	16.81	14.39	14.16
Fe <sub>2</sub> O <sub>3</sub>	3.45	4.09	0.13	3.19	2.04	0.96	0.14
FeO	5.94	3.61	1.33	7.32	0.71	1.80	1.26
MnO	0.16	0.16	0.05	0.19	0.03	0.05	0.02
MgO	6.15	4.77	0.12	1.53	1.71	0.69	0.15
CaO	8.77	6.69	1.34	4.80	5.05	2.55	0.36
Na <sub>2</sub> O	3.31	3.78	3.43	4.28	3.61	3.07	0.15
K <sub>2</sub> O	4.28	0.89	2.06	3.38	0.93	4.18	2.80
TiO <sub>2</sub>	1.95	0.95	0.10	1.69	0.54	0.46	0.03
P <sub>2</sub> O <sub>5</sub>	0.70	0.29	0.10	0.85	0.33	0.17	0.03
H <sub>2</sub> O <sup>+</sup>	0.34	0.16	0.65	0.30	0.40	0.32	0.00
H <sub>2</sub> O <sup>-</sup>	4.77	4.81	1.32	0.94	1.04	0.88	3.16
TOTAL	99.41	99.26	100.78	99.66	100.21	98.50	99.84
				Norms			
Quartz	0.000	4.201	46.710	5.503	30.006	28.275	67.317
Orthoclase	25.291	5.259	12.173	19.973	5.495	24.700	17.114
Albite	6.975	31.983	29.022	36.214	30.545	25.976	1.313
Anorthite	10.866	26.899	5.995	12.226	22.898	11.540	1.645
Diopside	22.497	3.561	0.000	5.128	0.000	0.000	0.000
Hypersthene	0.000	16.197	0.299	9.579	4.259	3.564	0.386
Olivine	6.951	0.000	0.000	0.000	0.000	0.000	0.000
Magnetite	5.002	3.552	1.260	4.625	0.824	1.392	1.448
Ilmenite	3.704	1.804	0.151	3.210	1.026	0.874	0.044
Corundum	0.000	0.000	2.771	0.000	1.472	0.586	10.654
Apatite	1.621	0.672	0.232	1.968	0.764	0.394	0.072
Nepheline	11.394	0.000	0.000	0.000	0.000	0.000	0.000
Differentiation index	43.65	41.44	87.90	61.68	66.05	78.95	85.74
Irvine and Baragar (1971) classification	Alkalic trachybasalt	Subalkalic andesite	Subalkalic rhyolite	Tholeiitic andesite	Subalkalic andesite	Subalkalic adamellite (quartz monzonite)	Low-calcium rhyolite

uplift and lowering of local base level by glacial scour continue to modify surficial and bedrock units of the landscape.

## STRUCTURE

Layered rocks of the study area have undergone isoclinal to subisoclinal folding and synkinematic thrusting (see cross sections A, B, C, D). The Terra Cotta and Sheep Creek anticlines can be traced south and north of the study area for at least 100 km and 70 km along strike, respectively (fig. 1, map inset). Most subisoclinal folds trend N. 10°-40° E., plunge to the northeast, and are locally overturned to the northwest. However, west of the Sheep Creek fault (fig. 1), fold asymmetry is reversed, and axes plunge to the southwest and are overturned to the southeast. The trend of fold axes sharply turns 20° E. near Bowser Creek (fig. 1). Three major overturned folds in the McGrath A-2 Quadrangle are responsible for an estimated 20 to 25 km of crustal shortening. Locally, minor slip has occurred along thrust faults. Isoclinal folding and thrust faulting occurred after Lower Cretaceous units (KJs) were deposited and before rocks were intruded by plutons and dikes during Paleocene time. These plutonic rocks apparently were intruded during extensional processes in early Tertiary time. Later, east-west to northeast broad warping locally steepened thrust faults and isoclinal folds. The Tatina fault (after Reed and Lanphere, 1972) separates the Mesozoic flyschoid rocks from the Dillinger terrane; the fault trace is obscured by Quaternary deposits throughout the map area.

Northwest-trending, high-angle faults, such as the Bowser Creek fault, cut bedrock units and older structures; however, lateral and vertical displacement along these youthful faults only amounts to a few hundred meters.

## ECONOMIC GEOLOGY

Mineralization in the McGrath A-2 Quadrangle consists of lead-zinc-silver-copper skarn replacement bodies in limestone; veins and stockwork that contain zinc, tungsten, tin, and silver in igneous rocks; cobalt- and nickel-bearing massive pyrrhotite replacement deposits; and anomalous lead, zinc, and silver in shale (SOsh). The assay data in table 4 are not derived from channel samples and hence do not necessarily show average metal values of mineral deposits in the area.

Base-metal skarns at Bowser Creek (table 4) are typical of low-temperature, fracture-controlled lead-zinc skarns described by Einaudi and Burt (1982). Reed and Elliott (1968a,b) described the Bowser Creek occurrences and presented assay and prospect map information in addition to that reported here. The main Bowser Creek deposit (table 4, map no. 2) consists of replacement deposits of pyrrhotite, sphalerite, galena, and

chalcopryrite in a hedenbergite-johannsenite skarn next to a felsite dike. The main pyrrhotite-sphalerite occurrences are in a johannsenite-rich exoskarn in the marble front; however, important fissure-controlled, silver-rich, galena-tetrahedrite-pyrrhotite-calcite mineralization occurs in the marble 5 to 25 m from the marble front. The fissure deposits have been mined for silver, and modest amounts of ore were marketed before 1973. Some metasomatic fluids have replaced felsic dikes with anhedral garnet and pyroxene and produced a small area of endoskarn. Skarn deposits northeast of Bowser Creek (table 4, map no. 3) are apparently associated with quartz monzonite porphyry. A sphalerite-pyrrhotite-johannsenite skarn zone 3 to 10 m thick continues at distances of 14 to 40 m within the marble away from the intrusive contact. Assay data (table 4) for both deposits show positive correlation of silver with lead content.

Fracture-controlled, cobalt- and nickel-bearing massive pyrrhotite deposits cut hornfelsed shale (SOsh) in a bedrock canyon near Post Lake (table 4, map no. 5). Several massive pyrrhotite zones 0.5 to 2 m thick and 20 to 80 m long trend N. 15°-25° W. and parallel the joint-cleavage set in the host rock. A small quartz monzonite plug is exposed 1 km north of the occurrences. Skarn mineralogy is notably absent. Pyrrhotite samples contain up to 570 ppm copper, 156 ppm cobalt, and 637 ppm nickel, well below commercial assay levels. However, the occurrences are similar to the Chip-Loy nickel-cobalt deposit (McGrath A-3 Quadrangle) that contains up to 2.25 percent nickel and 0.18 percent cobalt (Herreid, 1968; Gilbert and Solie, 1983).

Veins and stockwork in igneous rocks occur at the head of Bowser Creek and in the Post River pluton. The Bowser mineralization (table 4, map no. 4, fig. 2, and map inset) occurs as 3- to 20-cm-thick veins of pyrrhotite, sphalerite, and minor chalcopryrite in a felsic intrusion (Tqp) that is similar in composition to the felsic dike exposed in the skarn described above. Sulfide-bearing veins in the intrusion are localized along steeply dipping joints that trend N. 65-70° W. Skarn mineralogy is notably absent. Occurrences in the Post River pluton (table 4, map no. 6a,b) consist of pyrite ± scheelite-chalcopryrite-quartz-calcite veinlets in hornfelsed roof pendants, in fracture zones in the intrusion, or in the contact zones between intrusions and country rock. The mineralized areas are poorly exposed and of unknown extent. A sample of mineralized vein found in the rubble yielded 1.1 percent copper, 0.089 percent tin, 98 ppm silver, and 0.028 percent tungsten. Scheelite occurs in pan concentrates from ravines that drain the south side of the intrusion. The Post River intrusion is similar to a tungsten-bearing quartz monzonite plug in the McGrath B-2 Quadrangle (Bundtzen and others, in preparation) and is also similar to tungsten-bearing intrusions that cut identical Paleozoic rock units (Rabbit Kettle Formation) in the Selwyn Basin, Canada (Anderson, 1982).

Table 4. Descriptions and analytical results of selected mines and minerals prospects and occurrences in the McGrath A-2 Quadrangle, Alaska (sample data in ppm).

Map no.	Field no.	Au	Ag	Cu	Pb	Zn	Mo	Sb	W	Sn	As	Co	Ni	Cr	Remarks
1	82MK92	ND	0.4	83	38	453	10	ND	2	ND	ND	ND	52	64	Random chip channel; geochemical samples of sulfurous (SOsh) shale.
	82MK93	ND	0.6	38	19	335	17	28	2	ND	ND	ND	24	74	
Main Bowser Creek															
2	82BT237	0.02	140.0	3,060	3,300	93,560	ND	ND	1	1	209	157	13	13	Skarn with disseminated sulfides. Average of six high-grade samples of massive pyrrhotite-galena ores. Sphalerite-rich ores; contain 0.10 and 0.11% cadmium, respectively. (See map inset, fig. 2).
	82BT238	0.11	2,510.0	1,510	237,600	34,430	ND	ND	1	10	816	3	7	9	
	82BT235 2527	0.30 ND	3.7 10.6	652 990	69 377	35,500 110,000	13 -	ND -	1 -	6 -	17 ND	57 42	14 36	7 64	
Northeast Bowser Creek															
3	83BT117a	ND	391.0	200	240,000	1,600	-	-	-	-	-	-	-	-	Selected sample of galena-pyrrhotite ore. Sphalerite-rich vein in marble.
	83BT117b	0.7	98.0	5,500	2,600	221,000	-	-	-	-	-	-	-	-	
Headwaters of Bowser Creek															
4	82BT236	ND	83.0	550	2,780	34,800	-	ND	1	1	469	55	102	108	Selected samples from pyrrhotite-sphalerite-chalcopyrite quartz vein in quartz porphyry body; sphalerite-rich ores contain up to 0.12% cadmium and 0.2% barium.
	2507	ND	6.2	78	358	1,070	-	ND	2	1	23	ND	ND	39	
	2508	ND	9.5	110	202	1,210	-	ND	3	ND	44	10	10	50	
	2514	ND	50.0	10,900	2,060	123,000	-	ND	3	ND	83	169	30	52	
Post Lake prospect															
5	83BT158	-	-	244	-	ND	10	ND	1	1	240	156	154	-	Grab sample of pyrrhotite 'ore.' Selected 0.5-m-long chip channel. 1-m-long chip channel; massive pyrrhotite contains 33.5% Fe. 0.5-m-long chip channel; contains 13.1% Fe. Grab sample of massive pyrrhotite. (See map inset, fig. 3).
	83BT158a	0.7	0.1	182	23	28	2	20	2	1	378	60	104	61	
	83BT158b	-	-	570	10	170	21	ND	1	1	ND	139	167	10	
	83BT158c 83BT158d	- ND	- ND	429 230	- 840	294 302	20 14	ND ND	1 2	1 1	25 40	38 118	183 637	137 -	
Post River pluton															
6a	82BT269	ND	0.8	759	10	41	ND	ND	4	ND	ND	45	76	77	*1-m-wide pyrite zone in skarn. Chalcopyrite-quartz-scheelite vein.
	83MSL62	0.9	98.0	11,500	67	413	49	ND	275	890	ND	ND	66	134	
6b	83BT119	ND	13.7	5	6	14	3	ND	3	3	91	10	20	105	Sulfide vein in intrusive rock (Tqm). Contact zone between Tqm and SOsh. Quartz vein in Tqm; 0.1-m-wide chip sample. Concentrate sample below Tqm-SOsh contact in creek ravine.
	83BT120	ND	ND	35	16	76	18	ND	3	26	56	ND	77	90	
	83BT122a	ND	8.1	47	1,820	191	19	117	125	5	38	11	51	137	
	83BT122b	ND	59.0	29	24	114	9	41	3	8	ND	ND	31	80	
7a	82BT316	ND	ND	155	6	709	5	ND	2	ND	-	37	196	62	Random chip samples of sulfur-rich shale (contains SOsh).
	82BT486	0.3	13.6	1,220	-	13,700	3	1	1	1	ND	14	11	25	
8	83BT303	0.3	1.2	28	41	700	31	ND	2	1	52	4	30	56	Three 10-cm-thick stratiform pyritic layers with disseminated sphalerite.
9	82MK203a	ND	4.5	33	186	8,660	2	ND	3	13	ND	13	52	65	Chip samples of sulfur-rich shale (SOsh). Same as above.
	82BT321	ND	2.0	104	61	172	5	ND	3	ND	ND	ND	-	145	

Analyses by M.R. Ashwell, T.A. Benjamin, and M.K. Polly using X-ray fluorescence and atomic-absorption spectrophotometry, DGGs Minerals Laboratory.

ND = Not detected.

- = Not analyzed.

Composite grab samples of sulfur- and iron-rich shale (SOsh) locally contain anomalous lead, zinc, silver, and molybdenum (table 4, map no. 1, 7a,b, 9). The occurrence at map number 9 (table 4) is coincident with a significant barium anomaly reported in stream-sediment sample data by Aamondt and others (1979).

Selected shale samples were analyzed for Rock-Eval pyrolysis geochemistry. Four samples of shale (SOsh) contain 1.20 to 1.93 percent total organic carbon and 0.05 mill/g free hydrocarbons. Three samples of Silurian shale (mSs) average only 0.37 percent total organic carbon and 0.03 mill/g free hydrocarbon. Hence, the shale (SOsh) has fair to good potential as a hydrocarbon source. However, all Paleozoic units are overmature and contain conodonts with maturation indexes that range from 3 to 5.

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Table 5. Fossil identifications from the McGrath A-2 Quadrangle, Alaska.

Map no.	Field no.	University of Alaska Museum no.	Description
1	82BT152	A-1841	Graptolites in calcareous sandstone of mSl. <i>Pristiograptus</i> cf. <i>P. dubius</i> , Wenlockian (Middle Silurian) age.
2	82BT190	--	Poorly preserved, ribbed atrypaccan brachiopod in mSl; suggests Silurian to early Late Devonian range.
3	82BT175	A-1864	Graptolites midway through SOsh; mixed collection of <i>D. decoratus</i> and <i>C. bicornis</i> zones, Middle Ordovician age; <i>D. decoratus</i> zone contains <i>Pterograptus</i> cf. <i>P. elegans</i> (f), <i>Cryptograptus schaeferi</i> (r), <i>Didymograptus</i> sp. (r), <i>Glossograptus</i> cf. <i>G. acanthus</i> (f), <i>G. ciliatus</i> (c), <i>G. ciliatus</i> cf. subsp. <i>antennatus</i> (r), <i>Phyllograptus</i> cf. <i>P. angustifolius elongatus</i> (r). <i>C. bicornis</i> zone contains <i>Climacograptus bicornis</i> (c), <i>Dicellograptus</i> cf. <i>D. divaricatus divaricatus</i> (r), <i>D. cf. D. divaricatus salopiensis</i> (c).
4	82BT176A	A-1874	Graptolite collection made at top of SOsh 6 m stratigraphically below ISl; contains <i>Monograptus</i> cf. <i>M. planus</i> (r), late Llandoveryan (late Early Silurian) age.
5	82BT174	A-1851	Collection made midway through SOsh section contains graptolites of <i>P. tentaculatus</i> zone, late Early Ordovician age. Includes <i>Isograptus forcipiformis laus</i> (r), <i>Brachiograptus etaformis</i> (f), <i>Cryptograptus schaeferi</i> (r), <i>Phyllograptus nobilis</i> (c), <i>Glyptograptus</i> cf. <i>G. austrodentatus</i> (r), <i>Glyptograptus</i> sp. (f), <i>Didymograptus</i> sp. (f).
6	82BT173	A-1846	Collection made near core of Sheep Creek anteline on west side of SOsh; contains graptolites of <i>Oncograptus</i> zone, Early Ordovician age. Includes <i>Isograptus forcipiformis</i> (r), <i>Didymograptus</i> cf. <i>D. v-deflexus</i> (r), <i>Didymograptus</i> sp. (r).

Table 5. (con.)

Map no.	Field no.	University of Alaska Museum no.	Description
7	82BT170	A-1970	Collection made in axis of Sheep Creek anticline in SOsh; graptolites are of <i>T. approximatus</i> and <i>T. fruticosus</i> zones, earliest Ordovician age, includes <i>Tetragraptus approximatus</i> (r).
8	82BT172	A-1862	<i>Phyllograptus?</i> sp. (r) in SOsh east of axis of Sheep Creek anticline.
9	82BT169	A-1948	Graptolite collection from SOsh near ISI contact; contains poorly preserved diplograptids including <i>Climacograptus?</i> and <i>Glyptograptus?</i> , Middle or Late Ordovician age.
10	82BT168	A-1848	Graptolite collection from mSI contains <i>Pristiograptus</i> cf. <i>P. dubius</i> (a), retiolitid (r), Wenlockian (Middle Silurian) age.
11	82BT186	--	Graptolite collection from ISI contains <i>Cryptograptus cennifugus</i> , late Early Silurian age.
12	82BT185	A-1861	Scree slope in SOsh contains <i>Cyrtograptus</i> sp. (r), <i>Dicellograptus?</i> sp., Middle Ordovician age.
13	82BT301	--	<i>Pristiograptus dubius</i> , Wenlockian (Middle Silurian) age.
14	82BT362	A-1845	Scree slope in SOsh contains undiagnostic diplograptid graptolites, Middle or Late Ordovician age.
15	82MK209	A-1938	SOsh contains graptolites <i>Climacograptus caudatus</i> (c), <i>Glyptograptus?</i> sp. (r), <i>Orthograptus</i> aff. <i>O. calcareus</i> s.l. (r), <i>Didymograptus?</i> sp. (r) of <i>C. nubiliferus</i> zone, late Middle Ordovician age.
16	82BT258	A-1907	Small limestone unit 20 m above ISI contains <i>Cryptograptus</i> sp. (r), <i>Monograptus</i> sp. aff. <i>M. fumus</i> (c), <i>M.</i> sp. ( <i>priodon</i> type) (c), <i>Monoclimacis</i> sp., earliest Wenlockian (Middle Silurian) age.

Table 5. (con.)

Map no.	Field no.	University of Alaska Museum no.	Description
17	82BT335	A-1935	Collection in SOsh contains <i>Climacograptus bicornis</i> (c), <i>Orthograptus</i> aff. <i>O. calcareus</i> s.l. (r), <i>Glyptograptus</i> sp. (f), <i>Dicellograptus</i> cf. <i>D. divaricus salopiensis</i> (c), <i>Dicranograptus?</i> sp. (r) of <i>C. bicornis</i> zone, Middle Ordovician age.
18	82MKTC1	A-1856	Base of SOsh contains <i>Tetragraptus</i> cf. <i>T. quadribrachians</i> (r), Early Ordovician age.
19	82MKTC2	A-1857	Upsection from map no. 18 in SOsh; contains <i>Glossograptus</i> sp. (r), <i>Phyllograptus</i> sp. (r), <i>Climacograptus</i> sp. (c), <i>Glyptograptus</i> sp. (f), <i>Didymograptus?</i> sp. (r) of <i>D. decoratus</i> zone, early Middle Ordovician age.
20	82MKTC3	A-1858	Fragments of <i>Didymograptus?</i> sp. indicates Early to Middle Ordovician age.
21	82MKTC4	A-1859	Upper portion of SOsh section yields <i>Pristiograptus</i> sp. (r), <i>Monograptus</i> aff. <i>M. exiguus</i> (r), <i>Monograptus</i> sp. (f), Llandoveryan (Early Silurian) age.
22	82MKTC5	A-1860	Collection in carbonate layer above ISI yields <i>Monograptus</i> aff. <i>M. priodon</i> , Early to Middle Silurian age.
23	83BT132	--	mSI contains compressed orthocone nautiloid shells, probable Silurian age.
24	83BT147	--	OEs contains abundant branching trace fossils on shale partings similar to <i>Thalassinoides</i> and <i>Tonowangea</i> from basinal deposits in Australia, Europe, and North America.
25	83BT148	--	Lowest carbonaceous shale horizon in SOsh yields <i>Tetragraptus</i> cf. <i>T. approximatus?</i> and <i>Didymograptus laus acquilis</i> of <i>T. fruticosus</i> zone, Early Ordovician age.

Table 5. (con.)

Map no.	Field no.	University of Alaska Museum no.	Description
26	82MK182B	A-1876a-b	SOsh yields <i>Glossograptus</i> cf. <i>G. ciliatus</i> (f), <i>Amplexograptus fallax</i> (r), <i>Orthograptus calcaratus acutus</i> (r), <i>Climacograptus</i> sp. (r), <i>Orthograptus calcaratus acutus</i> (r), <i>Climacograptus</i> sp. (r), and <i>Glyptograptus</i> sp. <i>Didymograptus?</i> sp. (r) of <i>C. bicornis</i> zone, Middle Ordovician age.
27	82MK193	A-1910	Carbonate unit in mSs yields <i>Monograptus</i> of <i>prionon</i> type; Middle Silurian(?) age.
28	82BF272	--	<i>Climacograptus bicornis</i> , Middle Ordovician age.
29	82BF192	A-1964	SOsh yields <i>Tetragraptus quadribachiatas</i> (c), Early Ordovician age.
30	82MA320	A-1843	SOsh yields <i>Climacograptus bicornis</i> , Middle Ordovician age.
31	82BF321	--	Uppermost part of SOsh below ISI yields <i>Monograptus spiralis</i> , Early Silurian age.
32	82MSL61	--	Kls yields <i>Inoceramus</i> sp. with ribbing similar to <i>Inoceramus peliformis pochialaynen</i> of Early Cretaceous (Hauterivian) age.
33	82MA300	--	<i>Thalassinoides?</i> trace fossil; very similar to map no. 24.

Fossils collected by T.K. Bundtzen (BF), K.M. MacDonald (MK), M.S. Lockwood (MSL), and M.D. Albanese (MA). Graptolite identifications by Claire Carter, U.S. Geological Survey. Pelecypod (map no. 32) identified by J.W. Miller, U.S. Geological Survey. Other megafossil examinations by R.B. Blodgett, Oregon State University. Trace-fossil identifications by Bundtzen after Hantzschel (1975). Most fossil collections are archived at the University of Alaska Museum in Fairbanks.

(a) = abundant

(f) = frequent

(c) = common

(r) = rare