

GEOLOGY OF THE SLEETMUTE A-5, A-6, B-5, AND B-6 QUADRANGLES, SOUTHWESTERN ALASKA

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John Decker, R.R. Reifentuhl, M.S. Robinson, C.F. Waythomas, and J.G. Clough



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Cover photo: *Northwest-vergent isocline of very fine-grained sandstone of the Lower(?) and Upper Cretaceous Kuskokwim Group. The outcrop of medium-bedded sandstone is 1.5 miles west of Kiokluk Lake in the Sleetmute B-6 Quadrangle (map unit Kkm). The hammer handle in the photo is 46 cm long. Photo by R.R. Reifentstahl.*

Fairbanks, Alaska
1995



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[in pocket]

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INTRODUCTION

The Sleetmute A-5, A-6, B-5, and B-6 Quadrangles are located 450 km (280 mi) west of Anchorage in the Kuskokwim Mountains of southwestern Alaska. Elevation in the map area ranges from about 1,300 m (4,265 ft) in the pluton-cored Kiokluk and Chuilnuk Mountains to 100 m (328 ft) along the Holitna River and Chineekluk Creek. Except for the glaciated peaks of the Chuilnuk and Kiokluk Mountains, the physiography of the area is dominated by rounded, tundra-covered hills of low to moderate relief.

Geologic mapping utilized 1:40,000-scale topographic base maps, 1:63,360-scale aerial photographs, and field notes and maps from unpublished U.S. Geological Survey field studies. A reconnaissance study by Cady and others (1955) included the quadrangles mapped in this report.

Preliminary 1:40,000-scale geologic maps of the Sleetmute A-6, B-5, and B-6 Quadrangles have been published, respectively, by Decker, Robinson, Murphy, Reifentstahl, and Albanese (1984), Reifentstahl, Robinson, Smith, Albanese, and Allegro (1984), and Robinson and others (1984). Rock, pan-concentrate, and stream-sediment geochemistry maps of the Sleetmute A-6, B-5, and B-6 Quadrangles were compiled by Robinson (1984a, b, c respectively). Compilations of previously unpublished data, mainly from U.S. Geological Survey field studies, are available in Alaska Division of Geological & Geophysical Surveys Public-data File maps (Decker, Reifentstahl, and Coonrad, 1984; Reifentstahl, Decker, Murphy, and Coonrad, 1984; and Murphy and others, 1984).

STRATIGRAPHY, AGE AND STRUCTURE

The oldest rock unit in the region is the Holitna Group of Cady and others (1955). These authors reported Silurian and Devonian fossils, but inferred that Ordovician strata may also be present. More recent field investigations have shown that strata as old as Cambrian and as young

as Triassic are also present within their Holitna Group. The oldest strata recognized in the unit, based on regional correlation with the Lone Mountain area of the McGrath Quadrangle, are of probable Late Proterozoic age (R.B. Blodgett, written commun., 1994). Adrain and others (in press) have suggested that the designation "Holitna Group" be ultimately abandoned, because it is too broadly defined. We agree with this, and use Holitna group as an informal unit designation within our report area.

Most of the area is underlain by the Kuskokwim Group, a >12-km-thick (39,370 ft) marine turbidite sequence of thick- to medium-bedded sandstone and shale. The age of the Kuskokwim Group has been generally considered Early Cretaceous and Late Cretaceous (Albian to Coniacian) (Moore and Wallace, 1985; Decker and others, 1994). The inclusion, by these authors, of the Early Cretaceous (Albian) age-range is based on the rare occurrence of one species of arenaceous foraminifera of probable Albian age from the Mulchatna River area. This occurrence was reported in an in-house ARCO Alaska fossil report (Moore and Wallace, 1985). Numerous published reports on Kuskokwim Group macrofossils indicate a Late Cretaceous (Cenomanian to Santonian) age and none reports ages other than Late Cretaceous (Cady and others, 1955; Elder and Box, 1992; Box and Elder, 1992; Elder and Miller, 1991). Consequently, we will refer to the Kuskokwim Group age as Early(?) and Late Cretaceous (Albian(?) to Santonian).

A far less extensive rock unit, superficially resembling the Kuskokwim Group, is the *Inoceramus*-bearing calcareous sandstone, siltstone, and shale of Cretaceous to Jurassic(?) age. Contacts between the calcareous deposits and other bedrock units are covered by surficial deposits.

The Holokuk Basalt (Cady and others, 1955) consists of calc-alkalic volcanic rocks (table 1) that unconformably overlie the Kuskokwim Group and range in composition from basaltic andesite to rhyolite. K-Ar ages range from 64.3 to 74.5 Ma. This formation is >1 km (3,280 ft) thick and consists of subaerial flows and interlayered agglomerate, tuff breccia, basalt-matrix conglomerate, and lahar deposits; it is spatially and chronologically associated with nearby plutonic rocks. Low initial Sr⁸⁷/Sr⁸⁶ ratios in the rhyolitic rocks preclude involvement of radiogenic continental crust in their genesis (table 2).

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Two granodiorite stocks of latest Cretaceous age (mean of K-Ar ages = 68.2 Ma; table 3) occur in the Chuilnuk and Kiokluk Mountains. The Holokuk Basalt is locally hornfelsed in the western Kiokluk Mountains by one of these plutons. Latest Cretaceous to earliest Tertiary volcano-plutonic complexes of similar size, age, and composition are common throughout southwestern Alaska (Bundtzen and Laird, 1982; Bundtzen and Laird, 1983; Bundtzen and Laird, 1992; Bundtzen and others, 1992; Decker, 1994; Miller and Bundtzen, 1994; Moll-Stalcup, 1994).

Hypabyssal rocks include mafic to felsic dikes and small stocks of Tertiary and Cretaceous age. The youngest volcanic rock is columnar-jointed olivine basalt, dated at 38 Ma (sample 20, table 3).

Quaternary surficial deposits, subdivided into 24 units, define four glacial advances: (1) Oskawalik Creek drift (early Pleistocene); (2) Chuilnuk drift (middle late Pleistocene); (3) Buckstock drift (late Pleistocene); and (4) Kiokluk peak drift (Holocene). The chronology of these surficial deposits was determined on the basis of relative dating.

The Kuskokwim Group is variably deformed by overturned and isoclinal folds, broad open folds, and local polyphase deformation. Layered Holokuk Basalt deposits were gently folded after 65 Ma. The Boss Creek and Holitna faults, which offset late Pleistocene deposits, are splays of the Denali-Farewell fault system and have uncertain Cretaceous and Tertiary histories.

MINERAL RESOURCES

Dickite-bearing altered dikes such as those in the andesite dikes map unit (Kim) are commonly associated with lode occurrences of cinnabar and antimony in the lower Kuskokwim region (Sainsbury and MacKevett, 1965).

Rocks of the lower Kuskokwim River region host placer deposits of gold and tungsten, vein deposits of mercury, antimony and other base and precious metals, and greisen deposits containing tin and other base and precious metals.

Gold and scheelite occur in placer deposits along Fortyseven Creek (lat 61°02'40" N., long 158°09' W.) and originate from a series of mineralized quartz veins that cut hornfelsed Kuskokwim Group where Fortyseven Creek heads. The hornfelsed sedimentary rocks weather a distinctive brown and contain secondary biotite; they were probably produced by intrusion of a small rhyolite plug near the head of Fortyseven Creek. The quartz veins are typical hydrothermal vein deposits, ranging in width from a few centimeters to more than a meter, and can be traced discontinuously for tens of meters along strike; they contain minor calcite and white mica, which yielded a K-Ar age of 60.9 ± 1.8 Ma. (sample 21, table 3). Chlorite

and sericite selvages border most of the veins. Ore minerals in the veins include gold, scheelite, wolframite, arsenopyrite, pyrite, jamesonite, stibnite, argentite, and rare gold-silver tellurides. Gangue minerals include quartz, carbonate, white mica, tourmaline, and sericite. Gold and scheelite have been recovered from placers (Cobb, 1976), but there is no record of production.

Several mercury and antimony deposits are known in the map area. The Mountain Top deposit (lat 61°23'47" N., long 157°58'23" W.), the best exposed of the mercury deposits, is located about 16 km (9.9 mi) north of Kiokluk Lake and has produced an unknown amount of mercury; visible reserves are estimated at 200 flasks (Cobb, 1976). The deposit is hosted in highly sheared and altered sedimentary and volcanic rocks. Most of the known mercury and antimony mineralization is localized in probable tectonic breccias near the contacts between brecciated and unbrecciated basaltic andesite. The source of mercury is not known, but some mineralized breccias at Mountain Top may represent vent breccias, and mineralization may originate from hot springs.

Tin occurs in quartz-tourmaline veins and stockworks in the Chuilnuk and Kiokluk plutons and in hornfelsed Kuskokwim Group sedimentary rocks near pluton contacts.

All known mineralized zones in the area appear to have downstream or downslope geochemical signatures. Stream-sediment geochemistry may therefore be useful in delineating new mineralized areas.

DESCRIPTION OF MAP UNITS

UNCONSOLIDATED DEPOSITS

- Qa MODERN FLOOD-PLAIN ALLUVIUM—Moderately to well-sorted cobble and pebble gravel, sand, and overbank silt along the channels of modern streams and rivers. Abundant organic debris locally and good primary stratification. May be mantled by thin (<1 cm) colluvium or loess or both
- Qaf ALLUVIAL-FAN DEPOSITS—Variously sorted, weakly stratified cobble and pebble gravel in a sand and silt matrix. Associated with cone- or fan-shaped alluvial deposits along the bases of mountain fronts and valley slopes. May contain lenses of angular colluvial gravel
- Qp PEAT, PEATY SILT, AND OTHER PALUDAL DEPOSITS—Generally thick accumulations of peat and organic silt in fens and swamps; locally interbedded with loess; usually capping fluvial sediments. Includes palsas, fens, and string fens

- Qc UNDIFFERENTIATED COLLUVIAL DEPOSITS—Angular, poorly sorted, unstratified gravel in a silty matrix on slopes of 45° or less. Derived principally from frost-shattered bedrock. May contain solifluction deposits and minor alluvium
- Qcs SOLIFLUCTION DEPOSITS—Poorly sorted, angular gravel in a silty matrix. Associated with lobate or tongue-like microrelief forms on colluvial slopes
- Qcx FINE-GRAINED COLLUVIUM—Retransported silt or muck, locally interbedded with sand and gravel lenses, mainly where it overlies alluvium on lower slopes and valley bottoms
- Qct TALUS—Frost-rived, blocky rock rubble forming cones, fans, and aprons along base of steep mountain slopes. Includes avalanche-boulder tongues, boulder trains and runs, and rockfall debris
- Qcl LANDSLIDE DEPOSITS—Hummocky, heterogeneous to chaotic debris deposited by slope failures, including slump, slides, and slips
- Qcd RETRANSPORTED DRIFT—Unsorted, unstratified silty gravel with variable amounts of sand and boulders on glaciated valley slopes; gradational with undisturbed drift
- Qrg ROCK-GLACIER DEPOSITS—Lobate or tongue-shaped accumulations of bouldery rubble below talus and scree slopes, showing evidence such as multiple, crescentic ridges, of downslope movement
- Qd UNDIFFERENTIATED GLACIAL DEPOSITS—Mixtures of till, outwash, and minor colluviated drift or other glacial diamicton not readily separated. Poorly sorted, generally unstratified, gravelly sediment with a sand to silt matrix
- Qd₄ GLACIAL DRIFT OF KIOKLUK PEAK ADVANCE—Poorly sorted bouldery drift in highest, north-facing cirque of Kiokluk Mountains
- Qd₃ GLACIAL DRIFT OF BUCKSTOCK GLACIATION—Diamicton associated with unmodified knob-and-kettle moraines. Poorly sorted bouldery gravels to well-sorted sandy gravel. May include some outwash gravel, flow till, and loess
- Qd₂ GLACIAL DRIFT OF CHUILNUK GLACIATION—Diamicton of subdued morainal deposits and outwash, forming isolated remnants close to modern valley floors. Poorly to moderately sorted silty gravel with variable amounts of sand and boulders; moraines exhibit local hummocky topography with kettles
- Qd₁ GLACIAL DRIFT OF OSKAWALIK CREEK GLACIATION—Diamicton of very subdued drift bodies partially capping ridges and knolls; includes minor outwash, retransported drift and solifluction deposits. Composed of poorly sorted silty gravel, locally frozen at depths of 2-3 m
- Qcc BLANKET COLLUVIUM—Blocky rock rubble with variable amounts of fine-grained material as blankets. Includes stone stripes, solifluction debris, stone-banked terraces, and protalus lobes
- Qld LACUSTRINE DELTAIC DEPOSITS—Well-sorted sands and silts deposited in standing bodies of water
- Qf UNDIFFERENTIATED FLUVIAL DEPOSITS—Moderately to well-sorted cobble and pebble gravel, sand, and overbank silt. Commonly with well-developed primary fluvial sedimentary structures, and associated with inactive flood plains and stream terraces
- Qcf UNDIFFERENTIATED COLLUVIUM AND ALLUVIUM—Complexly mixed colluvial and fluvial deposits forming debris-rich fans, on most mountain slopes or at mouth of steep bedrock walls, couliers, and valleys. Consists of poorly stratified, unsorted gravel to silt, often with crude clast imbrication and bouldery zones; includes deposits of mudflows, debris flows, slushflows, snow avalanches, and floods
- Qdo UNDIFFERENTIATED OUTWASH—Well-sorted cobble to small-boulder gravel, usually exhibiting obvious primary stratification. Contains rounded to subrounded clasts. Locally capped by 1-2 m of eolian silt or sand or both. Age uncertain
- Qdo₃ OUTWASH OF BUCKSTOCK GLACIATION—Well-sorted, stratified, cobble and pebble gravel and sand; locally mantled by 1-2 m of eolian silt. Outwash valley trains or outwash

- heads related to Buckstock terminal moraines and ice limits
- Qdo₂ **OUTWASH OF CHUILNUK GLACIATION**—Moderately to well-sorted, stratified, cobble and small-boulder gravel; locally mantled by 1-2 m of eolian sediment, colluvium or both. Most outwash valley trains or outwash heads related to Chuilnuk terminal moraines and ice limits
- Qdo₁ **OUTWASH OF OSKAWALIK CREEK GLACIATION**—Moderately well sorted cobble gravel forming terrace remnants not directly traceable to equivalent terminal moraines. Probable remnants of extensive valley-fill deposits that have been buried, redistributed, and dissected
- Qat **UNDIFFERENTIATED STREAM-TERRACE DEPOSITS**—Cobble to pebble gravel and sand, typically well sorted and stratified in stream terraces >2 m high. Commonly mantled by eolian silt

VOLCANIC ROCKS

- Tvb **COLUMNAR BASALT**—Rubble of fresh olivine basalt exposed locally on densely vegetated hillside 10 km northeast of Fortyseven Creek. One whole-rock sample yielded K-Ar age of 38.2 ± 1.1 Ma. (table 3)
- Tvr₃ **GRAY RHYOLITE**—Light- to medium-gray and buff, fine-grained, well-layered to massive rhyolite and rhyolite tuff breccia. Rhyolite contains irregular clots of hematite to 5 mm diam; weathered rhyolite is distinctively hematite-stained and spotted; exact nature of contact with underlying black rhyolite unit (Tvr₂) is unclear; one sample yielded K-Ar age of 43.8 Ma (table 3). Intermixed with hematite-spotted, gray rhyolite are zones of light-gray rhyolite tuff breccia composed of angular and subangular clasts of shale, sandstone, and volcanic rocks in a very fine grained scoriaceous matrix. Tuff breccia occurs in well-layered sequence near core of volcanic complex
- Tvr₂ **BLACK RHYOLITE**—Very dark gray to black, fine-grained, porphyritic rhyolite; includes minor basaltic andesite. Rhyolite contains alkali feldspar (sanidine) phenocrysts to 10 mm long and beta-quartz phenocrysts to 3 mm diam in a very fine grained, glassy groundmass. Quartz

and feldspar fragments are common. Locally contains exotic clasts of shale, sandstone, and rhyolitic to basaltic volcanic rock. Unit locally underlain and cut by thin (15 m) zones of dark-gray to black fragmental volcanic rock composed of shale, sandstone, rhyolite, and basalt clasts in obsidian-like matrix. Unit appears to be locally discordant; may represent dikes or sills

- Tvr₁ **RHYOLITE**—Tan to olive-green, fine-grained, porphyritic rhyolite with abundant quartz phenocrysts to 3 mm diam and alkali feldspar (sanidine) phenocrysts to 5 mm long; locally, phenocrysts constitute as much as 20 percent of rock. Rhyolite occurs as several thick (~30 m) layers which exhibit no perceptible change in composition. The contact between this unit and the overlying black rhyolite (Tvr₂) appears locally conformable. Exact nature of lower contact with basalt is unknown

HYPABYSSAL ROCKS

Holokuk Basalt

The Holokuk Basalt was named by Cady and others (1955). It is a gently dipping sequence of predominantly basaltic volcanic rock, containing minor rhyolite, vitric tuff, and breccia, in the western Sleetmute Quadrangle. The formation is at least 1 km thick and unconformably overlies the Kuskokwim Group. Cady and others (1955) assigned a Tertiary (late Eocene to early Miocene) age, mainly on the basis of structural arguments. New radiometric age data, however, suggest that Holokuk volcanism took place between 64.3 and 74.5 Ma, or during latest Cretaceous to very earliest Tertiary time. Whole-rock major-oxide data (table 1) indicate a basaltic andesite to rhyolite compositional range (Peccerillo and Taylor, 1976). Seven samples from the Holokuk Basalt yielded K-Ar ages ranging from 64.3 to 74.5 Ma (mean of K-Ar ages = 70.0 Ma; table 3); the Holokuk Basalt is therefore probably correlative with the Upper Cretaceous Iditarod Basalt (Cady and others, 1955), and possibly the undated Getmuna Rhyolite Group (Cady and others, 1955) in the northern Sleetmute and southern Iditarod Quadrangles. In the Kiokluk and Chuilnuk Mountains, rocks previously mapped by Cady and others (1955) as the Holokuk Basalt are here subdivided into 10 informal units.

- Khap **BASALTIC ANDESITE**—Dark-gray to brown and green, very fine grained, porphyritic basaltic andesite. These rocks contain plagioclase

phenocrysts to 3 mm long, clinopyroxene (augite?) to 2 mm diam, olivine in anhedral masses to 2 mm diam, and orthopyroxene as subhedral grains to 3 mm long. Olivine and orthopyroxene are altered to serpentine minerals. The basaltic andesite occurs as thick flow(?) units which unconformably overlie rocks of Kuskokwim Group. Six basaltic andesites in the map area yielded K-Ar ages of 64.3, 68.3, 68.0, 70.7, 70.8, and 74.5 Ma (mean of K-Ar ages = 69.4 Ma; table 3)

Kh1b INTERLAYERED LAHAR DEPOSITS AND BASALTIC ANDESITE—Grayish-green, brownish-red, dark-green, and black interlayered lahar deposits and basaltic andesite. Basal sequence positionally overlies Upper Cretaceous Kuskokwim Group and is locally crosscut by Cretaceous andesite dikes (Kim). Basaltic andesite occurs as massive flows and dikes and contains plagioclase, olivine, or pyroxene phenocrysts. Olivine is moderately to highly altered to serpentine minerals and magnetite. One K-Ar whole-rock sample gave a minimum age of 73.2 ± 2.2 Ma (table 3). This unit is 426 m thick in a measured section 5 km east of Holokuk Mountain (Reifenstuhl, Robinson, Smith, Albanese, and Allegro, 1984; sheet 2)

Kh1 RED AND GRAY LAHAR DEPOSITS—Brownish-red, matrix supported deposit; locally contains angular to rounded pyroxene basalt clasts from 0.5 to 70 cm diam. Muddy volcanoclastic matrix locally forms up to 60 percent of the rock. Gray lahar deposit is distinctive grayish green, unstratified, and composed of angular, very poorly sorted mafic volcanic clasts in weakly indurated groundmass. Locally interbedded with graded muddy sandstone and mudstone; individual beds range in thickness from 10 to 80 cm; may be tightly folded; contacts with volcanic rocks are sharp and irregular

Kha RED BASALTIC ANDESITE—Massive, brownish-dark-red basaltic andesite with plagioclase phenocrysts to 4 mm long in matrix of plagioclase laths <0.2 mm long, brown volcanic glass, and chalcedony. Locally, subhedral orthopyroxene is 1.2 mm long and constitutes up to 8 percent of rock. Occurs above and below red and gray lahar deposits unit (Kh1)

Khr WHITE RHYOLITE—White aphanitic rhyolite

containing quartz phenocrysts to 4 mm diam, with laminated, swirled, and contorted flow banding

Khf INTERMEDIATE TO FELSIC FLOWS, TUFF, AND AGGLOMERATE—Textures and compositions vary greatly within this unit, and only thick, distinct layers can be correlated laterally. Prominent lithologies are: (1) dark-green to gray, platy-weathering, thinly laminated aphanitic andesite containing aligned plagioclase laths, relict hornblende(?), and minor secondary chlorite with local, small-scale isoclinal flow(?) folds; (2) light-green to medium-gray felsic- to intermediate aphanite; (3) tan-weathering intermediate agglomerate containing subangular, intermediate lithic clasts to 2.5 cm long, with epidote and calcite along fractures and to 3 percent disseminated pyrite; (4) finely laminated amygdaloidal basalt with secondary calcite and chlorite; and (5) finely laminated, light-green lithic tuff with clasts to 1.5 cm long, a black matrix, and locally graded bedding. Conformably overlies massive and amygdaloidal basaltic andesite unit (Khaa). Unit is hornfelsed by Kiokluk pluton in western Kiokluk Mountains

Khaa MASSIVE AND AMYGDALOIDAL BASALTIC ANDESITE—Sequence of basaltic andesite and andesite flows with thick massive bases and thin amygdaloidal tops. Massive flows are dark-gray, olivine-augite andesites with fine-grained trachytic groundmass. Olivine is subhedral, generally 0.2 mm diam, and constitutes 1 percent of rock. Augite is euhedral to subhedral, 1.5 mm long, and 2 percent of the rock. Amygdaloidal andesites are gray to black where fresh, and weather to maroon and green. Amygdules are 2.5 cm long, irregularly shaped, rounded or flattened, commonly lined with concentric chalcedony, and filled with one or more of the following: quartz, calcite, epidote, or chlorite. Amygdaloidal zones range from 0.6 to 8.0 mm thick; amygdule size increases from base to top. Basalt-matrix conglomerate occurs at the base of the unit and contains subrounded shale and sandstone fragments derived from the underlying Kuskokwim Group. These clastic fragments are up to 2.5 cm long with little or no evidence of thermal metamorphism. This unit is overlain by the intermediate to felsic flows, tuff, and agglomerate unit (Khf). Unit is hornfelsed by Kiokluk pluton in western Kiokluk Mountains. Unit is 258 m thick at Peak 3770 and

- consists of at least six basalt flows, each with thick massive base and thin amygdaloidal top (Decker, Robinson, Murphy, Reifenhohl, and Albanese, 1984; sheet 2)
- Khtb** TUFF BRECCIA—White to green, felsic to siliceous, aphanitic to sandy-matrix tuff breccia. Local flow banding may be swirled or contorted. Quartz phenocrysts to 4 mm diam occur locally. Pyrite blebs to 1 cm diam constitute to 3 percent of rock
- Khb** FELSITE BRECCIA—White to light-gray, vertically jointed, matrix-supported breccia containing subangular felsite clasts to 0.6 m diam in a felsite matrix. Locally, north of Peak 3770, breccia matrix contains to 20 percent green to brown pleochroic tourmaline as radiating clusters with individual crystals to several centimeters long. Contact relationships with adjacent units unclear
- Khc** BASALT-MATRIX CONGLOMERATE—Brown-weathering sandstone and shale, pebble to granule conglomerate in a light-green volcanic matrix. Contains subrounded shale and sandstone fragments derived from the underlying Kuskokwim Group. Clasts are up to 2.5 cm long and indicate only minor thermal metamorphism. In western Kiokluk Mountains, unit is thermally metamorphosed by Kiokluk stock. In the northwest part of map area, this basalt unit is mapped as tuff breccia, which contains basalt and shale clasts in a sandy matrix

SEDIMENTARY AND METAMORPHIC ROCKS

Kuskokwim Group

The Kuskokwim Group (Cady and others, 1955) consists primarily of marine turbidites and subordinate fluvial and shallow marine strata deposited in an elongate, southwest-trending basin covering over 70,000 km² in southwestern Alaska. The Kuskokwim Group is thickest (>12 km) in the central part of the basin between the villages of McGrath and Aniak. Mixed marine and nonmarine sections are relatively thin (<3 km) and are restricted to basin margins and local basement highs. The Kuskokwim Group is Early(?) and Late Cretaceous (Albian(?) to Santonian) age, on the basis of sparse age-diagnostic megafossil collections (<50) and a few collections of fossil plants, palynomorphs, and dinoflagellates (table 4; Decker and others, 1994).

The Kuskokwim Group in the Sleetmute A-5, A-6, B-5, and B-6 Quadrangles is poorly exposed; it typically forms low, rolling hills mantled with vegetation and frost-shattered bedrock. There are few outcrops where bedding can be determined with certainty and stratigraphic relations observed; some occur along cutbanks of major rivers, in steep gulleys, and in hornfelsed rock adjacent to the Kiokluk and Chuilnuk stocks.

The Kuskokwim Group in and around the Kiokluk and Chuilnuk Mountains consists entirely of marine turbidites, subdivided here into two informal units: medium-bedded sandstone and shale (Kkm), and thick-bedded sandstone and shale (Kkt). Bouma (1962) intervals for individual turbidite beds are indicated by subscript: (T_{abcde}). Groups of beds are assigned to facies according to the model of Mutti and Ricci Lucchi (1972): B, C, D, E, F, G. It was not possible to consistently distinguish between shale related to turbidity currents (T_e) and that of normal hemipelagic regime (facies G). Thin shale beds interbedded with sandstone and siltstone turbidite are referred to arbitrarily as T_e , and thick shaly intervals are assigned to facies G.

Kkm MEDIUM-BEDDED SANDSTONE AND SHALE—Light- to dark-gray, medium- to very fine grained, massive, graded, plane-laminated and crossbedded lithic sandstone turbidites interbedded with medium-gray to black, massive, plane-laminated and crossbedded siltstone, and black shale. Sandstone beds range in thickness from 1 cm to 4 m; mean values range from 15 to 20 cm. Beds are laterally continuous and have sharp bases and sharp or gradational tops. Amalgamated beds and shale rip-up layers are rare, but well-preserved flute casts are common. This unit is characterized by alternating sand- and shale-dominated sequences that range to 25 m and 3 m, respectively. The sand to shale ratio is as high as 50:1, averages about 4:1, and locally is less than 1:1. Both thickening- and coarsening-upward, and thinning- and fining-upward megacycles occur. Bouma (1962) intervals in sandstones include massive or graded T_{ab} , T_{ac} , T_{abcde} , T_{bcde} , T_{cde} , T_{ce} , and T_{de} . Bouma interval T_{ce} is dominant, and these rocks show internal cross-lamination, sharp bases, and sharp tops. Bouma intervals in shaly sections composed of fine- to very fine grained sands and shale include T_{bcde} , T_{cde} , and T_{de} ; T_{de} is dominant. Shaly sections are composed of repeated 1- to 3-cm-thick turbidite beds interlayered with black shale. Thin sands have either sharp or gradational tops and are interpreted to be facies D

and E deposits; medium-bedded sandstone sequences assigned to facies C, D, and E; suggests middle- and outer-fan deposition

Kkt THICK-BEDDED SANDSTONE AND SHALE—Light- to dark-gray, very coarse to very fine grained, massive, graded, plane-laminated and cross-laminated lithic sandstone turbidites interbedded with black, massive, plane-laminated and cross-laminated siltstone and silty shale. Sandy, pebbly, and chaotic mudstones also occur locally. Unit characterized by thick alternating sequences of sandstone and shale with sand to shale ratios of 100:1 to 1:1,000 (commonly about 10:1)

Sandstone beds from <1 cm to >5.5 m thick (mean thickness = 0.4 m) occur in massive shale sections, either as thin, laterally discontinuous beds with sharp bases and tops, or as thick, commonly amalgamated beds with scoured bases and gradational tops; contain diffuse laminations composed of shale-chip intraclasts (see Kuskokwim Group stratigraphic section, sheet 2). Other bedding features include load, flute, and groove casts. Paleocurrent indicators suggest a southwesterly transport direction (fig. 1). Shale-chip intraclasts range in long dimension from microscopic to 7 cm and may occur at any level in the enclosing sandstone bed. Bouma intervals in sandstone-dominated sections include T_a , T_{ac} , T_{ab} , and T_{abe} and subordinate T_{ce} and T_{cde} . Bouma intervals in shale-dominated sections include parallel-laminated T_{de} and subordinate, thin T_{bc} and T_{cde} ; sandstone and siltstone intervals all have sharp bases and tops. Some shale-dominated sections are composed of poorly sorted sandy and pebbly (sometimes chaotic) mudstones, which are matrix-supported and contain clasts of angular to nebulitic shale and sandstone intraclasts and *Inoceramus* sp. fragments. Pebbly mudstones correspond to facies F chaotic deposits, whereas sandstone-dominated sections correspond mainly to facies B, and the shale-dominated sections to facies D and G. Facies associations indicate inner-fan deposition.

Unit also characterized by moderately to well-developed foliation, structurally dismembered sandstone phacoids in deformed shale, moderate recrystallization, grain elongation up to 2:1, highly shattered and flaggy outcrops, and local white quartz veins. Swale and ridge topography

reflects alteration of sandstone and shale; however, structural deformation, including faulting, foliation, and isoclinal and open folding, masks lateral continuity of bedding

Kkq MINERALIZED QUARTZ VEINS AND QUARTZ-VEIN STOCKWORK AND HORNFELS—Near Fortyseven Creek: T. 11 N., R. 50 W.; (secs. 8, 17, and 18) brown-weathering hornfelsed clastic sediments with mineralized quartz veins and quartz-vein stockwork. Hornfelsed zone is adjacent to small rhyolite bodies. The brown-weathering character and fine-grained mineral laths distinguish this hornfels zone from nonhornfelsed rocks and from other nonmineralized hornfels. Southwest of Fortyseven Creek, a hornfelsed zone similar to the zone at Fortyseven Creek may represent an extension of the Fortyseven Creek trend. Quartz veins and their alteration selvages contain native gold, scheelite, and arsenopyrite. The Fortyseven Creek occurrence also contains wolframite, arsenopyrite, pyrite, jamesonite, stibnite, argentite, and traces of gold-silver tellurides; tourmaline and sericite accompany quartz gangue (Cady and others, 1955). On Fortyseven Creek downstream from the main mineralized shear zone, two distinct mineralized horizons occur in placer deposits: gold concentrations on bedrock and small amounts of gold distributed in gravels above bedrock, and gold and scheelite concentrations on a yellow clay false bedrock

Kkh HORNFELSED SANDSTONE AND SHALE—Hornfelsed clastic rocks of the Kuskokwim Group occur adjacent to plutons, stocks, and dikes. Intensely hornfelsed rock contains poikiloblastic biotite; very fine grained, subhedral, granular zoisite(?); sparse euhedral pyrite; black-tourmaline rosettes; and fibrous sillimanite. These inner-zone minerals are enclosed in completely recrystallized detrital grains and, rarely, by myrmekite; visible biotite found adjacent to large igneous bodies

KJs CALCAREOUS SANDSTONE, SILTSTONE, AND SHALE—Medium-brown and dark-gray, light-reddish-brown-weathering sandstone, siltstone, and shale; locally crosscut by quartz and carbonate veinlets; sandstone is generally medium grained and poorly sorted; contains abundant carbonate fossil fragments; larger fragments west of Kiokluk Mountains have been identified as *Inoceramus* sp. (of probable

Cretaceous age, table 4, map no. VI). Contact relationships with units in report area obscured by Quaternary deposits

Holitna Group

DSH LIMESTONE AND DOLOMITE—Very light to medium-gray, thin- to massive-bedded, finely laminated, limy and dolomitic mudstone to wackestone with interbeds of grainstone. Represents shallow, subtidal to peritidal cyclic carbonate sedimentation (Clough and others, 1984). Commonly fossiliferous in subtidal wackestone to grainstone facies. Locally recrystallized and brecciated in areas adjacent to faults. Estimated thickness of Holitna group in the central Kuskokwim region is 1,500 to 3,000 m (Cady and others, 1955). Age in map area is Silurian to Devonian, on the basis of our lithologic correlation with equivalent, megafossil-bearing strata in the Kutukbuk Hills, approximately 14 km east of map area

INTRUSIVE ROCKS

Tit/
Tird/
Tirg ALBITE RHYOLITE AND ALBITE RHYOLITE PORPHYRY—Medium-light-gray to buff, fine- to medium-grained, porphyritic, biotite-albite rhyolite and white mica-biotite-albite rhyolite porphyry. Tertiary age assigned on the basis of intrusive relationship with Upper Cretaceous Kuskokwim Group and three K-Ar ages on white mica (61.7, 63.6, and 61.5 Ma, table 3). Biotite-albite rhyolite is characterized by euhedral and subhedral grains of quartz, albite, biotite and sanidine in a matrix of very fine grained quartz, feldspar, and altered glass. Geochemistry indicates a rhyolite to rhyodacite range (table 1). Quartz phenocrysts contain resorbed rims and resorption channels. Sanidine and albite are altered to fine-grained white mica and clay minerals locally. Subhedral grains of biotite to 8 mm across are fresh and appear to be primary phase of rhyolite. Minor white mica (muscovite?) occurring as subhedral grains to 5 mm across appears also to be primary. Scattered subhedral to euhedral grains of dark-blue tourmaline occur throughout rhyolite. Near western flank of Holokuk Mountain, a breccia body contains angular to rounded fragments of tourmalinized rhyolite and rhyolite porphyry set in matrix of quartz and tourmaline. Where phenocrysts of quartz and feldspar become dominant constituent, unit is rhyolite porphyry. No stratigraphic evidence for age exists other than

apparent intrusive relationship into Kuskokwim Group. Similar albite rhyolite on Barometer Mountain (near Sleetmute, northeast of map area) yielded one minimum whole-rock K-Ar age of 67.9 ± 2.0 Ma and one K-Ar age on biotite of 70.5 ± 2.1 Ma (DGGS unpublished data, 1984); relationship with this unit is unclear

Tird—Greenish medium-gray, fine-grained, altered, porphyritic quartz-plagioclase feldspar rhyolite phase. These non-biotite rocks are deeply weathered, with green clots of dickite(?)

Tirg—Medium-gray, fine-grained, porphyritic garnet-biotite rhyolite. Medium-grained garnet and biotite in fine-grained, potassium-feldspar-rich matrix

TKii HYPABYSSAL INTERMEDIATE INTRUSIVE ROCKS—Medium-gray and light-gray porphyritic quartz diorite to granodiorite with medium-grained hornblende and plagioclase phenocrysts (mean length = 3 mm) in fine-grained matrix of randomly to slightly oriented plagioclase laths. Intrudes Upper Cretaceous Kuskokwim Group. Forms dikes and small stocks. Locally, quartz monzonite to granite occur, containing potassium feldspar phenocrysts to 3.5 cm long and medium-grained quartz and biotite phenocrysts. Extensive hydrothermal alteration products include calcite, chlorite, epidote, opaque minerals, and fine-grained white mica

TKi APLITE—Light- to dark-gray and buff, fine- to medium-grained aplite dikes composed of quartz, potassium feldspar, and plagioclase feldspar, with minor white mica, biotite, and hornblende

TKif FELSITE DIKES—Light-gray to buff, fine-grained felsite dikes containing disseminated, fresh or oxidized pyrite cubes, or euhedral garnet

Kkgd GRANODIORITE TO GRANITE OF THE KIOKLUK MOUNTAINS—Medium- to dark-gray, nonfoliate hypidiomorphic-granular, medium- and coarse-grained hornblende-biotite and biotite-hornblende granodiorite to granite. Geochemistry indicates a granite to quartz monzonite composition (table 1). The color index is approximately 18. The Cretaceous age of Kiokluk pluton is based on the K-Ar age of one biotite separate: 68.7 ± 2.1 Ma (table 3).

Plagioclase (An_{35}) is subhedral to euhedral and is slightly altered to fine-grained white mica, locally to epidote. Quartz is anhedral and medium to fine grained. Potassium feldspar is typically medium grained, occurs as subhedral and euhedral grains to 7 mm long, and locally forms rosettes. Biotite occurs as 0.5- to 3.0-mm subhedral to anhedral grains that constitute 5 to 8 percent of the rock and may surround hornblende, producing mafic mineral clots. Hornblende occurs as 2- to 4-mm-long subhedral grains that form 6 to 12 percent of the rock and is moderately altered to white mica, chlorite, epidote, and opaque minerals; accessory apatite is common. Quartz-tourmaline veins and minor greisens occur locally

Kcgd/
Kcg GRANODIORITE AND GRANITE OF THE CHUILNUK MOUNTAINS—Medium- to dark-gray, nonfoliate hypidiomorphic-granular, medium-grained, hornblende- biotite granodiorite. Two geochemical analyses indicate a granodiorite composition (table 1). Textures range from equigranular to seriate and porphyritic. Plagioclase feldspar (An_{38}) is the predominant feldspar and occurs as subhedral to euhedral grains to 10 mm long. Potassium feldspar usually occurs as interstitial, anhedral grains. Biotite occurs as subhedral grains from 1 to 3 mm across and constitutes 11 percent of rock. Hornblende occurs as subhedral grains and anhedral masses to 4 mm long and constitutes 2 to 10 percent of rock. Quartz occurs as anhedral masses that constitute >8 percent of rock. Accessory minerals include zircon, apatite, and opaques. Locally, granodiorite is hydrothermally altered to chlorite, white mica, epidote, and minor calcite. Crosscutting quartz-tourmaline veins occur locally and are characterized by thin selvages of tourmaline surrounding quartz-dominant vein centers. Granodiorite yielded K-Ar ages on biotite of 67.5 ± 2.0 Ma and 68.4 ± 2.1 Ma (table 3). The border phase, near the eastern and northern boundaries of Chuilnuk pluton (Kcg), is light- to medium-gray, nonfoliate, seriate to porphyritic, medium- to coarse-grained granite. One geochemical analysis indicates a granite composition (table 1). Potassium feldspar occurs as euhedral and subhedral phenocrysts to 10 mm long and as subhedral masses in groundmass. Biotite occurs as subhedral grains that constitute to 10 percent of rock; minor hornblende is also present. Accessory zircon and apatite are common, and epidote, sericite, and chlorite occur

where granite is hydrothermally altered. Two K-Ar ages on biotite separates from the granite phase (Kcg) are 68.7 ± 2.1 Ma and 67.5 ± 2.0 Ma (table 3)

Kgi GRANODIORITE—Small granodiorite intrusion in northeastern part of Sleetmute B-5 Quadrangle; age and relationship to other granodiorite intrusions uncertain

Kii INTERMEDIATE TO MAFIC STOCKS AND DIKES—Medium- to dark-gray, medium-grained, granodiorite and biotite, quartz-bearing gabbro. Intrudes Upper Cretaceous Kuskokwim Group. A minimum K-Ar age on one hornblende separate is 63.8 ± 1.9 Ma (table 3); unit cross-cut by andesite dike unit (Kim), which is dated at 69.8 ± 2.1 Ma. Textures range from equigranular in stocks to porphyritic in dikes. Plagioclase (An_{50}) occurs as grains to 3 mm long in stocks and to 4.5 mm long in porphyritic rocks. Augite occurs locally, forming to 15 percent of rock in a mafic border phase of stocks and to 5 percent in one dike. Alteration is moderate; products include fine-grained white mica, chlorite, carbonate and opaque minerals

Kim ANDESITE DIKES—Medium- to dark-gray, fine- to medium-grained andesite containing subhedral plagioclase as dominant feldspar and locally minor quartz and potassium feldspar. Geochemical analysis yields an andesitic composition (table 1). Andesite dikes crosscut lower(?) part of interlayered lahar deposits and basaltic andesite unit (Kh1b) which has been radiometrically dated at 73.2 ± 2.2 Ma. One K-Ar age is 69.8 ± 2.1 Ma (table 3). Andesite dikes also crosscut the intermediate to felsic flows, tuff, and agglomerate unit (Khf), the felsite breccia unit (Khb), and the intermediate to mafic stocks and dikes unit (Kii). Mafic-mineral content variable; includes biotite and hornblende. Andesite dikes altered locally; plagioclase altered to clay minerals, some of which are green (dickite?), biotite to chlorite, and hornblende to biotite and epidote

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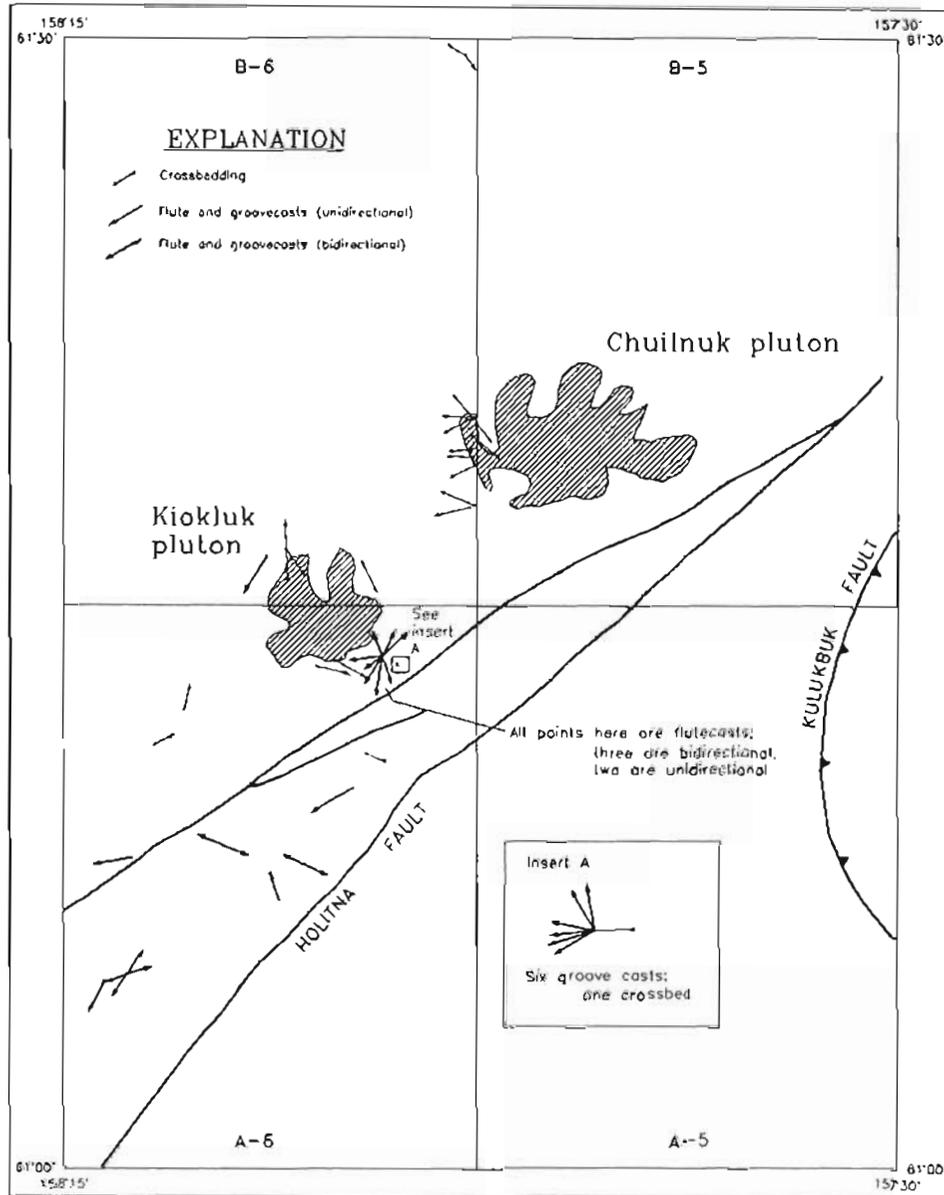


Figure 1. Paleocurrent data for the Sleemute A-5, A-6, B-5, and B-6 Quadrangles. Data have been corrected for tectonic dip.

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Table 1. Major oxide geochemical analysis of rock samples^a

Map no.	Rock type ^b	Map unit	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MgO	CaO	Na ₂ O	K ₂ O	TiO ₂	P ₂ O ₅	MnO	LOI ^c	Total
1	Basaltic andesite	Khap	54.81	16.64	3.41	5.62	4.61	6.62	3.33	1.95	1.16	0.22	0.13	0.95	99.45
2	Basaltic andesite	Khap	53.84	17.23	4.33	5.14	2.32	6.69	3.50	1.35	1.64	0.24	0.13	3.54	99.95
3	Rhyolite	Tir	70.98	14.82	2.44	1.18	0.23	1.33	3.81	3.72	0.11	0.03	0.02	1.00	99.67
4	Rhyolite	Tir	68.36	15.06	3.31	2.29	0.92	1.95	3.71	3.41	0.56	0.13	0.07	1.23	101.00
5	Dacite	Khap	57.68	16.77	2.41	6.47	4.39	6.05	3.48	1.76	1.03	0.17	0.14	0.01	100.36
6	Basaltic andesite	Khap	55.35	16.45	2.38	6.56	4.92	6.95	3.55	1.63	0.91	0.13	0.13	0.19	99.15
7	Basaltic andesite	Tvr ₂	54.73	14.79	4.20	6.12	2.53	5.44	3.80	3.03	1.97	0.27	0.15	2.49	99.52
8	Rhyolite	Tvr ₃	75.74	11.99	2.59	0.56	0.08	0.44	3.39	4.54	0.17	0.02	0.02	0.52	100.06
9	Rhyolite	Tvr ₃	76.02	11.71	2.52	0.59	0.05	0.36	3.15	4.97	0.17	0.02	0.02	0.31	99.89
10	Rhyolite	Tvr ₃	75.00	13.22	2.67	0.09	0.22	0.19	2.56	4.64	0.20	0.02	0.00	1.07	99.88
11	Basaltic andesite	Khap	53.77	16.51	4.37	4.04	4.45	6.13	2.57	1.85	1.12	0.18	0.08	5.18	100.25
12	Basaltic andesite	Khap	54.01	15.52	3.50	5.77	6.95	7.50	2.86	1.20	0.85	0.12	0.14	1.63	100.05
13	Quartz latite	Kh1b	70.12	14.75	2.85	0.90	0.59	1.71	2.78	3.94	0.19	0.04	0.02	1.63	99.52
14	Quartz latite	Kh1	70.29	15.00	2.84	0.87	0.56	1.54	2.93	3.92	0.20	0.04	0.02	1.32	99.53
15	Basaltic andesite	Kh1b	51.36	16.43	6.37	3.69	5.72	9.45	3.20	0.72	1.22	0.13	0.14	1.18	99.61
16	Basaltic andesite	Kh1b	54.46	14.51	4.15	5.16	7.53	7.57	2.97	1.19	0.85	0.11	0.14	0.70	99.34
17	Andesite	Kh1b	54.86	13.21	4.07	4.99	9.77	7.04	2.80	1.35	0.78	0.11	0.13	0.39	99.50
18	Rhyolite	Tir	74.06	14.58	2.49	0.56	0.05	0.32	3.09	3.47	0.04	0.73	0.06	1.71	101.16
19	Rhyolite	Tir	71.56	14.93	2.58	0.64	0.01	0.26	4.72	3.41	0.01	0.75	0.06	0.98	99.91
20	Rhyolite	Tir	71.65	15.02	2.50	0.72	0.01	0.23	4.61	3.47	0.01	0.81	0.06	1.03	100.12
21	Rhyolite	Tir	71.96	14.81	2.58	0.64	0.01	0.24	4.74	3.41	0.01	0.77	0.06	0.98	100.21
22	Rhyolite	Tir	70.72	15.67	2.26	0.99	0.01	0.27	4.93	3.39	0.01	0.45	0.12	0.79	99.61
23	Rhyolite	Tir	70.82	15.60	2.72	0.55	0.01	0.24	4.89	3.49	0.01	0.43	0.06	0.73	99.55
24	Rhyodacite	Tir	61.99	15.77	3.46	4.03	2.15	3.47	3.79	2.90	0.95	0.21	0.12	1.48	100.32
25	Rhyolite	Tir	71.30	15.78	2.53	0.46	0.01	0.32	4.40	3.67	0.01	0.36	0.07	1.10	100.01
26	Basaltic andesite	Kha	54.38	14.60	3.84	5.49	8.38	6.69	2.94	1.25	1.01	0.14	0.14	0.26	99.12
27	Basaltic andesite	Kha	54.65	14.85	3.91	5.31	8.01	6.72	2.95	1.18	1.02	0.15	0.15	0.41	99.31
28	Basaltic andesite	Khap	53.93	14.39	3.84	5.59	6.93	7.62	3.02	1.21	1.19	0.17	0.15	1.75	99.79
29	Basaltic andesite	Khap	54.19	16.27	4.51	5.17	5.95	7.95	2.99	1.36	1.02	0.16	0.14	0.77	100.48
30	Granite	Kcg	72.62	13.23	2.78	1.02	0.31	1.16	3.33	4.45	0.26	0.03	0.03	0.31	99.53
31	Granodiorite	Kcgd	63.66	15.62	3.18	3.73	2.02	3.79	3.77	2.92	0.91	0.14	0.10	0.71	100.55
32	Granodiorite	Kcgd	65.18	16.09	1.52	3.61	1.77	3.78	3.54	3.26	0.86	0.15	0.10	0.28	100.14
33	Basaltic andesite	Khaa	53.34	16.10	4.26	4.75	5.80	8.13	2.99	0.90	0.82	0.11	0.13	3.02	100.35
34	Andesite	Kim	63.24	15.66	3.11	2.97	2.75	2.87	4.45	3.37	0.60	0.16	0.05	--	99.23
35	Andesite	Khaa	58.45	14.89	4.14	3.45	5.49	5.92	3.03	1.49	0.79	0.12	0.09	2.67	100.53
36	Andesite	Khf	59.45	13.37	2.82	4.49	6.03	5.26	2.83	3.21	0.56	0.21	0.10	1.59	99.92
37	Quartz monzonite	Kkgd	63.66	15.62	3.18	3.73	2.02	3.79	3.77	2.72	0.91	0.14	0.10	0.71	100.35
38	Quartz monzonite	Kkgd	64.56	14.94	2.65	3.33	2.81	3.63	3.15	3.83	0.66	0.13	0.08	0.58	100.35
39	Quartz monzonite	Kkgd	63.77	14.78	2.81	3.40	2.89	3.78	3.14	3.85	0.66	0.14	0.08	--	99.30
40	Hornfels	Kch	66.97	8.55	2.27	3.23	3.51	10.48	1.53	1.17	0.48	0.13	0.07	--	98.39
41	Basaltic andesite	Khaa	54.90	15.29	3.52	5.68	8.11	6.56	3.12	1.22	0.99	0.13	0.14	0.43	100.09
42	Dacite	TKif	58.44	15.08	3.11	3.31	2.87	3.83	5.62	1.11	0.71	0.13	0.08	0.97	95.26
43	Aplite	TKif	72.32	15.08	2.52	0.46	0.02	0.29	4.88	3.57	0.03	0.03	0.03	0.95	100.18
44	Dacite	Tkif	69.63	14.36	2.64	3.01	0.92	1.61	3.33	3.96	0.55	0.19	0.08	0.58	100.86

^aX-ray fluorescence analyses by the DGGs Geochemical Lab, Fairbanks, Alaska. Major oxide values are in percent.^bPeccerillo and Taylor (1976) classification system.^cLoss on ignition.

-- = Not analyzed.

Table 1a. Normative mineralogy (calculated using U.S. Geological Survey Petcal program)

Map no.	CIPW Norms												Total
	Quartz	Corundum	Orthoclase	Albite	Anorthite	Hypersthene	Wollastonite	Diopside	Magnetite	Hematite	Ilmenite	Apatite	
1	6.44	--	11.70	28.61	25.07	14.90	--	5.52	5.02	--	2.24	0.52	100.02
2	10.98	--	8.27	30.72	28.33	7.75	--	3.63	6.51	--	3.23	0.58	100.00
3	31.90	2.21	22.28	32.67	6.49	0.59	--	--	3.59	--	0.21	0.07	100.01
4	28.31	2.04	20.20	31.46	8.85	2.97	--	--	4.81	--	1.07	0.30	100.01
5	8.66	--	10.36	29.34	24.85	17.72	--	3.24	3.48	--	1.95	0.39	99.99
6	3.43	--	9.54	29.75	29.31	19.27	--	3.27	3.42	--	1.71	0.30	100.00
7	6.73	--	18.45	33.14	14.79	6.92	--	9.19	6.28	--	3.86	0.64	100.00
8	--	--	--	--	--	--	--	--	--	--	--	--	--
9	38.05	0.54	29.49	26.77	1.66	0.13	--	--	1.48	1.51	0.32	0.05	100.00
10	--	--	--	--	--	--	--	--	--	--	--	--	--
11	12.39	--	11.50	22.87	29.50	13.41	--	0.99	6.67	--	2.24	0.44	100.01
12	6.06	--	7.21	24.59	26.38	20.15	--	8.54	5.16	--	1.64	0.28	100.01
13	35.19	2.96	23.78	24.03	8.40	1.50	--	--	2.47	1.21	0.37	0.10	100.01
14	--	--	--	--	--	--	--	--	--	--	--	--	--
15	5.49	--	4.32	27.51	28.79	7.99	--	13.99	8.95	0.30	2.35	0.31	100.00
16	6.40	--	7.13	25.48	23.06	18.56	--	11.39	6.10	--	1.64	0.26	100.00
17	5.13	--	8.05	23.90	19.66	23.77	--	11.78	5.95	--	1.50	0.26	100.00
18	42.97	5.77	20.62	26.29	--	0.13	--	--	1.90	1.20	0.08	--	98.96
19	31.38	3.51	20.37	40.37	--	0.03	--	--	2.25	1.05	0.02	1.54	100.52
20	31.04	3.72	20.71	39.40	--	1.85	--	--	2.21	--	0.02	1.36	100.31
21	31.57	3.35	20.31	40.42	--	0.03	--	--	2.25	1.05	0.02	1.42	100.42
22	29.33	3.94	20.27	42.21	--	0.18	--	--	3.32	--	--	1.60	100.85
23	29.36	3.82	20.87	41.87	--	0.03	--	--	1.96	1.40	0.02	1.42	100.75
24	17.55	0.60	17.34	32.44	16.03	8.65	--	--	5.08	--	--	0.49	98.18
25	32.00	4.62	21.93	37.64	--	0.03	--	--	1.70	--	0.02	1.90	99.84
26	5.65	--	7.47	25.16	23.21	23.09	--	7.52	5.63	--	1.94	0.33	100.00
27	6.66	--	7.05	25.24	24.06	22.06	--	6.90	5.73	--	1.96	0.35	100.01
28	6.37	--	7.29	26.06	22.58	17.44	--	11.88	5.68	--	2.31	0.40	100.01
29	6.66	--	8.06	25.37	27.03	14.84	--	9.17	6.56	--	1.94	0.37	100.00
30	33.62	0.91	26.50	28.40	5.60	0.78	--	--	2.65	0.97	0.50	0.07	100.00
31	19.38	--	16.13	32.01	17.73	7.88	--	0.18	4.63	--	1.74	0.33	100.01
32	19.95	0.23	19.29	29.99	17.80	8.56	--	--	2.21	--	1.64	0.35	100.02
33	7.63	--	5.46	25.99	28.61	14.44	--	9.65	6.35	--	1.60	0.26	99.99
34	13.95	--	20.68	37.94	12.90	8.75	--	0.32	4.54	--	1.15	0.37	100.60
35	15.40	--	9.00	26.20	23.12	13.53	--	4.80	6.13	--	1.53	0.28	99.99
36	11.12	--	19.29	24.35	14.54	16.46	--	8.50	4.16	--	1.08	0.50	100.00
37	15.98	--	25.16	25.39	13.73	12.91	--	1.22	3.73	--	1.50	0.38	100.00
38	19.07	--	22.68	26.71	15.35	9.29	--	1.49	3.85	--	1.26	0.30	100.00
39	18.18	--	22.91	26.76	14.97	9.09	--	2.41	4.10	--	1.26	0.33	100.01
40	33.50	--	7.03	13.16	13.22	--	2.87	25.65	3.35	--	0.93	0.31	100.02
41	5.16	--	7.23	26.49	24.19	23.53	--	6.08	5.12	--	1.89	0.30	99.99
42	9.64	--	6.92	50.15	14.53	8.50	--	3.46	4.76	--	1.42	0.32	99.70
43	29.94	2.75	21.26	41.61	1.25	0.05	--	--	1.51	1.50	0.06	0.07	100.00
44	29.56	2.12	23.34	28.10	6.73	4.86	--	--	3.82	--	1.04	0.44	100.01

-- = Not present.

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Table 2. Trace element geochemical analyses of rock samples

Map no.	Rb ^a (ppm)	Rb ^b (ppm)	Sr ^b (ppm)	$\frac{87\text{Sr}}{86\text{Sr}}^c$ (ppm)	$\frac{87\text{Sr}}{86\text{Sr}}$ (ppm)	K ^b (%)	Na ^a (%)	Ba ^a (ppm)	La ^a (ppm)	Hf ^a (ppm)	Ta ^a (ppm)	Ce ^a (ppm)	Nd ^a (ppm)	Sm ^a (ppm)	Eu ^a (ppm)	Tb ^a (ppm)	Yb ^a (ppm)	Th ^a (ppm)	Tb ^b (ppm)	U ^a (ppm)	Y ^b (ppm)	Nb ^b (ppm)	Zr (ppm)	Sc ^a (ppm)	Cr ^a (ppm)	As ^a (ppm)	Sb ^a (ppm)	W ^a (ppm)	Fe ^a (%)	
1	59.00	64.00	439.00	0.70536	0.70574	1.58	2.21	910.00	19.10	3.71	0.80	41.80	19.40	4.78	1.35	0.72	2.31	4.60	5.00	2.43	26.00	9.70	168.00	19.60	83.40	5.25	0.65	137.00	5.06	
2	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
3	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
4	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
5	62.90	63.00	324.00	--	--	1.53	2.36	747.50	18.00	4.24	0.92	40.00	20.20	4.63	1.05	0.69	2.81	5.51	5.00	2.80	26.00	9.60	166.00	18.90	151.00	57.10	2.23	175.00	2.36	
6	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
7	69.30	79.00	389.00	0.70317	0.70369	2.57	2.52	445.80	51.10	8.14	3.85	103.00	44.00	9.22	1.52	1.30	3.87	9.21	9.00	3.32	43.00	49.30	318.00	15.50	14.00	3.22	0.37	185.00	6.05	
8	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
9	229.00	224.00	25.00	--	--	4.06	2.02	156.90	62.60	7.40	5.75	131.00	49.40	9.95	0.20	1.71	6.59	25.80	29.00	9.85	75.00	48.60	196.00	1.67	2.77	12.00	1.68	346.00	0.58	
10	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
11	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
12	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
13	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
14	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
15	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
16	28.90	36.00	309.00	--	--	1.09	2.01	522.20	12.80	2.82	0.56	29.50	15.70	3.43	0.96	0.34	1.93	3.56	3	1.67	20.00	7.70	109.00	23.70	525.00	5.53	0.52	84.20	5.43	
17	35.20	41.00	307.00	0.70601	0.70636	1.03	1.95	349.10	14.40	2.42	0.76	32.50	13.70	3.40	0.93	0.57	1.43	3.60	3.00	1.69	15.00	8.90	109.00	19.60	895.00	8.84	0.76	117.00	5.08	
18	477.00	446.00	124.00	0.70604	0.71496	3.02	1.99	77.77	0.12	0.76	7.67	1.42	2.36	0.63	0.13	0.31	0.25	0.49	3	8.82	12.00	22.50	16.00	1.01	2.26	9.06	1.85	4.93	0.54	
19	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
20	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
21	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
22	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
23	483.00	460.00	1.00	--	--	2.93	3.17	61.01	0.454	0.67	8.20	1.55	3.03	0.36	0.11	0.20	0.28	0.29	3	9.02	14.00	24.70	18.00	0.94	2.58	65.00	3.39	246.00	0.67	
24	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
25	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
26	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
27	32.60	38.00	329.00	--	--	1.23	2.03	371.30	13.10	3.26	0.79	30.80	15.60	3.57	1.10	0.58	1.90	3.05	3.00	1.76	21.00	7.70	119.00	23.10	509.00	3.55	0.42	185.00	5.32	
28	40.10	38.00	371.00	0.70403	0.70432	1.04	2.07	484.60	14.00	3.25	1.06	36.80	14.80	3.90	1.10	0.41	1.72	3.18	3.00	1.59	20.00	10.70	139.00	22.20	498.00	3.79	0.45	199.00	5.71	
29	40.80	43.00	334.00	--	--	1.03	1.99	683.20	13.60	2.86	0.43	31.30	15.10	4.09	1.06	0.86	2.29	3.19	3.00	1.87	26.00	8.40	120.00	27.00	255.00	6.48	0.61	1.23	5.55	
30	161.00	158.00	77.0	0.70488	0.71040	3.8	2.18	812.90	33.00	3.38	2.35	68.90	28.90	5.95	0.50	1.09	3.63	13.90	15.00	3.29	41.00	7.90	110.00	4.04	5.45	6.91	1.32	721.00	1.08	
31	88.90	90.00	256.00	--	--	2.17	2.42	882.80	25.00	5.29	1.19	56.10	26.90	5.25	1.13	0.78	2.64	8.07	7.00	2.29	30.00	10.30	184.00	12.80	32.10	2.42	2.04	233.00	3.56	
32	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
33	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
34	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
35	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
36	80.00	94.00	434.00	--	--	2.59	1.81	1448.00	28.80	3.35	0.65	66.10	28.40	5.71	1.28	0.45	1.81	6.53	7.00	3.07	17.00	9.00	144.00	16.00	449.00	59.70	128.00	95.50	3.79	
37	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
38	160.00	165.00	345.00	0.70626	0.70752	3.37	2.05	977.50	31.30	5.02	1.51	66.50	26.80	4.61	0.98	0.52	2.02	16.00	17.00	6.49	22.00	11.10	193.00	10.40	151.00	25.00	3.79	319.00	2.73	
39	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
40	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
41	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
42	32.50	27.00	547.00	--	--	1.05	3.70	493.80	17.10	2.79	0.78	35.50	13.60	3.21	0.88	0.26	1.39	3.68	3	1.78	13.00	7.30	116.00	10.80	54.20	1.74	0.65	153.00	2.92	
43	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--
44	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--	--

*Instrumental neutron activation analysis by Becquerel Labs, Mississauga, Ontario, Canada.

^bAnalysis by X-ray fluorescence by Becquerel Labs, Mississauga, Ontario, Canada.^cCalculated initial ratio by Dr. Terry Davis, University of California, Davis, California.^dMeasured initial ratio by Dr. Terry Davis, University of California, Davis, California.

-- = Not analyzed.

Table 3. *K-Ar age analyses*^a

Map no.	Rock unit	Mineral dated	K ₂ O wt(%)	Sample wt(g)	⁴⁰ Ar _{rad} × 10 ⁻¹¹	$\frac{^{40}\text{Ar}_{\text{rad}}}{^{40}\text{K} \times 10^{-11}}$	$\frac{^{40}\text{Ar}_{\text{rad}}}{^{40}\text{Ar}_{\text{total}}}$	Age ^b (Ma)
1	Khap	Plagioclase	1.609	7.813	16.7	4.19	0.702	70.7 ± 2.1 ^c
2	Khap	Plagioclase	1.779	1.2043	18.5	4.20	0.834	70.8 ± 2.1
3	Tirg	Biotite	5.574	0.1598	56.44	4.09	0.877	69.0 ± 1.8
4	Tir	Biotite	8.548	0.206	79.6	3.76	0.469	63.6 ± 1.9
5	Khap	Whole rock	1.61	7.813	17.6	4.42	0.669	74.5 ± 2.2 ^c
6	Khap	Plagioclase	1.202	1.9044	12.0	4.03	0.856	68.0 ± 2.0
7	Trv ₁	Whole rock	4.849	2.5464	30.9476	2.5763	0.5061	43.8 ± 1.3
8	Khap	Plagioclase	0.28	3.0186	2.64	3.80	0.666	64.3 ± 1.9
9	Khlb	Whole rock	1.35	8.5938	14.5	4.34	0.607	73.2 ± 2.2 ^c
10	Tir	White mica	9.732	0.1189	10.4	3.64	0.833	61.7 ± 1.8
11	Tir	White mica	9.685	0.2106	87.2	3.63	0.865	61.5 ± 1.8
12	Khap	Plagioclase	0.51	2.8638	5.11	4.04	0.404	68.3 ± 1.9
13	Kcg	Biotite	7.991	0.1139	80.5	4.07	0.859	68.7 ± 2.1
14	Kcgd	Biotite	7.609	0.2957	76.3	4.05	0.74	68.4 ± 2.1
15	Kcgd	Biotite	8.423	0.5323	83.3	3.99	0.755	67.5 ± 2.0
16	Kkgd	Biotite	7.097	0.1757	71.5	4.07	0.889	68.7 ± 2.1
17	Kkgd	Hornblende	0.358	1.0787	3.62	4.08	0.578	68.9 ± 2.1
18	Kii	Hornblende	0.683	0.8824	6.39	3.78	0.272	63.8 ± 1.9
19	Kim	Biotite	1.45	0.124	14.85	4.13	0.657	69.8 ± 2.1
20	Tvb	Whole rock	1.532	8.7988	8.51	2.41	0.806	38.2 ± 1.1
21	Kkq	White mica	0.755	0.1442	87.03	3.6	0.866	60.9 ± 1.8

^aAnalyses by J.D. Blum and D.L. Turner, DGGs-University of Alaska Fairbanks Cooperative Geochronology Laboratory.

^bConstants used in age calculations: $\lambda_e = 0.581 \times 10^{-11}$; $\lambda_g = 4.962 \times 10^{-10} \text{ yr}^{-1}$; and $^{40}\text{K}/\text{K}_{\text{total}} = 1.167 \times 10^{-4} \text{ mol/mol}$.

^cMinimum age (sample does not meet petrographic criteria for a reliable date).

rad = radiogenic.

Table 4. *Fossil information*^a

Map no.	Rock unit	Latitude	Longitude	Specimen ^b no.	Fossil type	Age	Remarks
1	Kkm	61°26'12"	157°32'55"	M8089	Pelecypod <i>Inoceramus</i> sp.	Cretaceous(?)	
2	Kkm	61°21'12"	157°40'40"	M8082	Plant	Indeterminate	
3	Kkh	61°17'53"	157°53'19"	M8088	Pelecypod <i>Inoceramus</i> sp.	Cretaceous(?)	Fragments only
4	Kkh	61°17'19"	157°50'00"	M8085	Pelecypod <i>Inoceramus</i> sp. <i>Inoceramus hobetsensis</i>	Late Cretaceous (Turonian)	
5	Kkm	61°16'30"	157°57'18"	M8096	Pelecypod <i>Inoceramus tenuistraiatius</i>	Late Cretaceous (Turonian)	
6	Kkh	61°16'22"	157°57'11"	M8094	Pelecypod <i>Inoceramus</i> sp.	Cretaceous(?)	In deformed hornfels
7	Kkh	61°16'13"	158°00'09"	M8083	Pelecypod <i>Inoceramus</i> sp. <i>Inoceramus hobetsensis</i>	Cretaceous (Turonian?)	
8	KJs	61°17'27"	158°11'45"	M8090	Pelecypod <i>Inoceramus</i> sp.	Cretaceous(?)	Fragments only
9	Kkt	61°11'14"	157°57'59"	M8093	Pelecypod <i>Inoceramus</i> sp.	Cretaceous(?)	Cast of fragments

^aFossils identified by John W. Miller, U.S. Geological Survey, Menlo Park, California. Unpublished fossil report dated May 2, 1985.

^bUSGS Museum number.