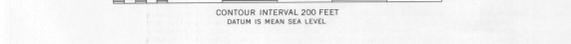
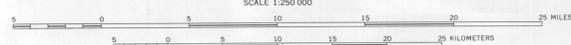
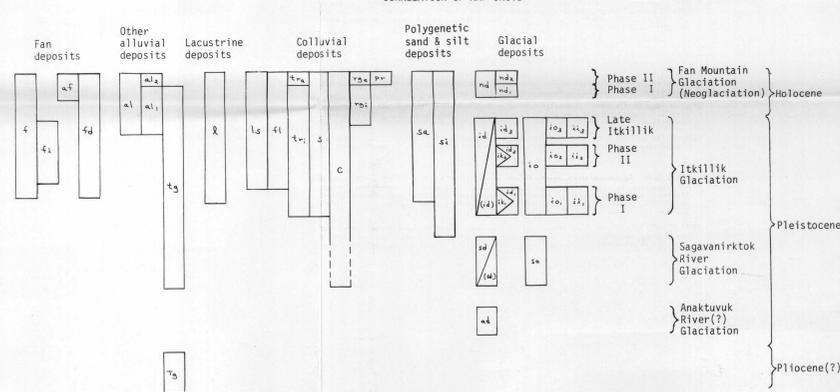


Base from U.S. Geological Survey, 1956

Geology by Reuben Kachadorian, 1969-1970; R.D. Reger and Roy Kreis, 1971-1974; and T.D. Hamilton, 1969-1977

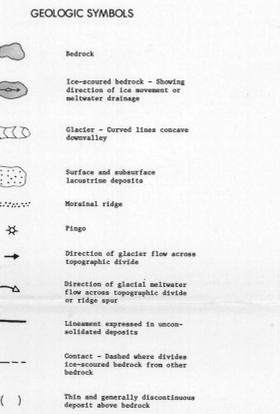


QUADRANGLE LOCATION



DESCRIPTION OF MAP UNITS

- INTRODUCTION**
- Surficial geologic mapping of the Wiseman quadrangle was based on (1) surface observations of morphology, composition, and distribution of unconsolidated deposits, (2) examination of test pits, roadcuts, placer workings, and other shallow exposures, (3) stratigraphic studies of representative river bluffs, and (4) correlation of geologic data prepared for the corridor occupied by the Alyeska Pipeline (for example, Kachadorian, 1971; Hamilton, 1979a). Map units are defined on the basis of lithology, stratigraphy, and geomorphology, and are related to units mapped previously within the neighboring Chandalar and Philip Smith Mountains quadrangles (Hamilton, 1979a, b).
- The basic stratigraphic framework for the surficial map units is provided by the central Brooks Range glacial sequence of Dettemann, Bowser, and Duto (1958), with modifications by Porter (1964), Hamilton and Porter (1970), and Hamilton (1979a and in press). Glaciers were more extensive at relatively high altitudes along the Continental Divide near the north margin of the Wiseman quadrangle and also in the Brooks Range and in localized glacial highlands of the John Valley. During the more extensive glaciations, ice flowed from these source areas southward through the major valley systems to footwalling basins beyond the south flank of the Brooks Range. Outwash trains were deposited along south- and west-flanking streams that issued from the ice fronts, and loess derived from outwash and other glacial deposits formed blankets between the flood-plain margins.
- Drifts of four major glaciations are recognized within the Wiseman quadrangle. The Anaktuvuk and Sagavirktok Glaciations of Dettemann, Bowser, and Duto (1958) are currently termed Anaktuvuk River and Sagavirktok River Glaciations to avoid confusion with previously named rock units (Bowser and others, 1966, p. 31 and 327). Although Sagavirktok River drift forms two units in its type locality (Hamilton, 1979b), only a single drift can be distinguished in the Wiseman quadrangle. Phases I and II of the Itkillik glaciation are as described by Hamilton and Porter (1970), but the Itkillik "III" moraines of those authors are here designated as drift of late Itkillik age. Fan Mountain I and II advances (Porter, 1964) are equated with the late Holocene Neoglaciation interval as described by Porter and Denton (1967) for the North American Cordillera. An additional glaciation, the Alaph Mountain Glaciation of Dettemann, Bowser, and Duto (1958), Porter (1964), and Hamilton (1979a, b), is not used here. Drift formerly assigned to Alaph Mountain is reinterpreted as alpine moraines of late Itkillik age.
- Radiocarbon dates from the neighboring Chandalar and Philip Smith Mountains quadrangles indicate that Itkillik I glaciation occurred more than 13,000 years B.P., that late Itkillik readvances took place about 12,500 to 11,500 yr B.P., and that post-Itkillik basin filling had commenced by at least 10,500 yr B.P. in some valleys (Hamilton, 1979b, c). The age of Itkillik II glaciation is uncertain but probably lies largely within the 23,000 to 16,000 yr B.P. time range of late Wisconsin maximum glaciation in the standard North American glacial succession (for example, Lindebaum and Goldswart, 1971). About 30 radiocarbon dates from the Wiseman quadrangle will be available by 1980.
- Lineaments are evident in unconsolidated deposits within the John, Wile, and Middle Fork valleys near the south flank of the ranges; they also occur south of the Koyukuk River near the confluence of its North and Middle Forks and are present within the mountain valleys of Glacier Creek and the North Fork. These features are expressed as aligned and unusually straight weathering profiles and vegetation lines, association of some places with ridges or pingo clusters, stream deflections, or abrupt linear boundaries of surficial map units. No vertical or horizontal offsets of unconsolidated deposits could be detected in the Fingus, Lower, and Upper Sagavirktok valleys. Linear elements are unproven. They may represent true faults, perhaps caused by tectonic readjustments at the time of deglaciation (Romer, 1970), alternatively, they may merely be surface expressions of underlying bedrock structures.
- Permafrost probably is present beneath all but the larger lakes and rivers of the Wiseman quadrangle. Its upper surface ranges in depth from more than 1 m to somewhat coarse-grained sediments at depths of a few centimeters. In poorly drained deposits beneath thick moss and sod cover. Although its general thickness is unknown, mining records (for example, Reed, 1958, p. 62 and 129) suggest that the base of permafrost may lie at a depth of slightly more than 100 m in some valleys.
- FAN DEPOSITS**
- FAN DEPOSITS, UNDIFFERENTIATED**—Range from poorly sorted, weakly stratified, subangular silty sandy coarse gravel at mouths of steep canyons to moderately sorted and stratified subrounded sandy gravel at mouths of large tributary streams with relatively gentle gradients. Locally subject to icings during winter (Sloan and others, 1976).
- INACTIVE FAN DEPOSITS**—Sandy gravel, as described above, usually with thick (more than 0.5 m) cover of organic silt to stony silt beneath unbroken sod and vegetation. Best developed in valleys that remained unglaciated during Itkillik I glaciation (for example, valleys of Metzenberg and Florence Creeks, Malamaute Fork of John River, and South Fork of Koyukuk River).
- DEPOSITS OF STEEP ALPINE FANS**—Coarse, very poorly sorted, nonstratified to weakly stratified, subangular to subrounded muddy sandy gravel at mouths of avalanche slopes and steep canyons near heads of mountain streams. Upper segments generally channelled, with levees of angular to subangular coarse debris. Subject to snow-avalanches during winter (Kachadorian, 1970). A thick layer of loess is present in some places with pingo clusters (p. 164-165), and debris flows during summer (Reger, 1975). Surface gradients (generally 12°-25°) intermediate between those of alluvial fans and talus cones.
- FAN-DELTA DEPOSITS**—Aluvial-fan facies (poorly sorted, weakly stratified, subangular sandy gravel near valley walls, grading into deltaic and lacustrine facies (well sorted and stratified silt, sand, and fine gravel) near valley centers. Occur in mountain valleys of John, Wile, and Middle Forks of Koyukuk, and Malamaute Fork of Alaina River, and North Fork of the Koyukuk. Locally subject to icings.
- OTHER ALLUVIAL DEPOSITS**
- ALLUVIUM, UNDIFFERENTIATED**—Ranges from poorly sorted, moderately well stratified, subangular to subrounded sandy gravel near heads of mountain valleys to moderately well sorted, well stratified fine to medium sand and lenses of sand and sandy silt. Includes fan, flood-plain, and low terrace deposits too small to be designated separately.
- MODERN ALLUVIUM**—Sand and gravel, as described above; generally unvegetated and commonly subject to icings (see Sloan and others, 1976). Differentiated only along Koyukuk River and its North, Middle, and South Forks, and along John River.
- LOW ALLUVIAL TERRACE DEPOSITS**—Sand and gravel, as described above; generally mantled with up to 0.5 m of organic silt, very fine sand, peat, and sod, and generally vegetated. Form terraces generally 0.5 to 3.0 m above upper limit of modern alluvium.
- TERRACE GRAVEL, UNDIFFERENTIATED**—Gravel and sandy gravel, as described above, in terrace remnants of uncertain origin along parts of Wolverine, Shukluk, Chitken, and Florence Creeks and near heads of Clear and John Rivers.
- TERTIARY(?) GRAVEL DEPOSITS**—Rounded to well rounded, well to moderately sorted pebbles and small cobbles in matrix of strongly oxidized coarse sand. Stones are better rounded, better sorted, smaller, and more quartzite (50 percent of total) than in any other local outwash, interstitial, or postglacial deposits. Forms residual loess near Ninimite Hills that stand 50 to 100 m high. Probably present beneath glacial drifts of Sagavirktok River(?) near southeast corner of map area.
- LACUSTRINE DEPOSITS**
- LACUSTRINE DEPOSITS, UNDIFFERENTIATED**—Well stratified silt and clay, containing sparse dropstones, grading into generally well stratified sand near former shorelines and sandy fine gravel near former stream mouths. Extensive thick deposits beneath Itkillik I-age moraines in valleys of John, Wile, Alaina, and Florence Rivers, Alaina and Sixymite Creeks, Malamaute Fork of Alaina River, and Koyukuk River's North Fork. Shown only by stippled map pattern where buried beneath or beneath Holocene alluvium, interstitial, or postglacial deposits, and fan deposits. Include beach and deltaic deposits too small to designate individually.
- COLLUVIAL DEPOSITS**
- LANDSLIDE DEPOSITS**—Unsorted, nonstratified, coarse to fine angular rubble, commonly with matrix of finer debris, forming lobes below detachment scars and slide tracks on high, steep rock walls. Subject to episodes of rapid downslope motion and long periods of relative stability. Most common in schist, phyllite, the Hunt Fork Shale, and other shale and siltstone units as mapped by Brose and Reiser (1971), and in very shaly and argillaceous units underlying the Anaktuvuk Conglomerate and by limestone. Only 18 of the 54 landfills recognized in the Wiseman quadrangle lie on or close to faults mapped by Brose and Reiser (1971).
- FLOW-SLIDE DEPOSITS**—Very poorly sorted, angular, tabular rubble in abundant matrix of generally micaceous sand and silt containing little clay. Forms lobes subject to slow and probably continuous downslope motion. Most numerous in phyllite, siltstone, and the Hunt Fork Shale as mapped by Brose and Reiser (1971); also present in quartz mica schist and in limestone with siltstone or phyllite interbeds.
- ACTIVE TALUS RUBBLE**—Angular, unsorted, nonstratified rock debris forming cones and aprons along lower walls of mountain valleys and in cirques at valley heads. Also forms thin (less than 1-2 m) and generally discontinuous sheets over bedrock. Generally unvegetated, unweathered to slightly weathered, and with lichen cover sparse to absent. Subject to rockfalls, especially during spring snowmelt season.
- INACTIVE TALUS RUBBLE**—Angular rock debris, as described above, generally weathered and lichen-covered, and commonly with partial sod cover. In places (less than 1-2 m) blankets of stratified talus occur on upland masses as bedrock beneath limits of Itkillik I glaciation in southwestern and south-central parts of map area.
- SOLIFLUCTION DEPOSITS**—Very poorly sorted, nonstratified to weakly stratified stony sandy silt and organic silt. In sheets and aprons more than 1 m thick on lower slopes of hills and on drifts and bedrock beyond outer limits of Itkillik I glaciation but also form generally smaller, more widely scattered deposits on shales, schist, and phyllite farther north (see Brose and Reiser, 1971). Several active, unweathered to weathered mountain valleys (for example, North Fork Koyukuk near Inok Creek and Moving Mountain; Wile Valley near Matheo Dene) appear to be associated with these deposits. They are mapped as talus cones (less than 3 m) sheets and aprons over many areas mapped as Itkillik I and older drift.
- COLLUVIUM, UNDIFFERENTIATED**—Mixed talus rubble and siltification deposits, as described above, forming aprons more than 1-2 m thick on slopes beyond 100 m of Itkillik I glaciation in valleys of Clear River, Mashochathuk Creek and Kevak Creek and near heads of Wolverine Creek, Shukluk Creek, and Malamaute Fork of the John River. Most common on phyllite, shale, and slate as mapped by Brose and Reiser (1971).
- ROCK-GLACIER DEPOSITS, ACTIVE**—Very poorly sorted, nonstratified, coarse angular rock debris, commonly with matrix of silt and fine rubble. Contains abundant interstitial ice. Upper surfaces generally unvegetated, unweathered to moderately weathered, and with sparse to moderate lichen cover. Frontal slopes barren, steep, and unstable, meeting upper surfaces at abrupt angle. Form (1) lobate deposits at bases of talus cones along valley walls and (2) tongue-shaped deposits within cirques (Lee White, 1976). The tongue-shaped variety commonly overlies stagnant glacier ice (Ellis and Galkin, in press). Subject to slow downslope motion.
- ROCK-GLACIER DEPOSITS, INACTIVE**—Coarse angular rock debris, as described above, but generally lacking abundant interstitial ice and underlying stagnant glacier ice. Upper surfaces and frontal slopes weathered, lichen covered, and commonly bear partial sod or vegetation cover. Frontal slopes grade into upper surfaces without abrupt angles.
- PROTALUS BARRIAGE DEPOSITS**—Unsorted, nonstratified, coarse, angular rock debris forming arcuate low ridges at bases of cirque headwalls. Associated with persisting snowbanks in shaded sites. Subject to rockfalls during spring snowmelt season.
- POLYGENETIC SAND AND SILT DEPOSITS**
- SAND**—Moderately sorted fine to medium sand, horizontally bedded to crossbedded, commonly with thin interbeds of sandy peat or organic silty fine sand. Deposited by slow-moving streams within basins partly dammed by end moraines in valleys of John and Alaina Rivers, Sixymite and Michtam Creeks, and Malamaute Fork of Alaina River. Over 0.5 to 5 m locally reworked by wind into sand sheets and dunes. Commonly grade downward into lacustrine deposits (see stippled map pattern).
- ORGANIC SILT**—Poorly stratified organic silt and silty peat, containing abundant ice as lenses, wedges and interstitial grains. Occurs beneath sparsum and black spruce in poorly drained depressions (muskeg) beyond Itkillik I glaciation. Most common near south flank of Brooks Range. Formed largely of loess, with amixed organic and siltification deposits.
- GLACIAL DEPOSITS**
- FAN MOUNTAIN DRIFT, UNDIFFERENTIATED**—Unsorted, nonstratified, coarse to fine angular rubble in cirque near north margin of map area. Designated (1) remnants of eroded fan Mountain drift recognizable in either Phase I or Phase II and (2) composite fan Mountain drift lobes too small for subdivision.
- DRIFT OF FAN MOUNTAIN PHASE II**—Angular rubble, as described above, forming lobes and arcuate ridges of ice-covered drifts with steep, weathered frontal slopes. Unvegetated to moderately weathered, and with lichens sparse to absent. Restricted to cirques and generally associated with active glaciers.
- DRIFT OF FAN MOUNTAIN PHASE I**—Angular rubble, as described above, forming dissected lobes and ridges with stable frontal slopes. Generally weathered and lichen covered, with partial sod cover in some localities.
- ITKILLIK DRIFT**—UNDIFFERENTIATED—Poorly to very poorly sorted nonstratified to weakly stratified till, ranging in composition from muddy sandy gravel to clayey stony silt, with local stratified ice-contact deposits consisting of moderately to poorly sorted sand, sandy gravel, and silty sandy gravel. Contains bedded and stratified stones up to boulder size. Designates thick (greater than 3 m) drift deposits within mountain valleys, that cannot be assigned to a specific Itkillik I moraine system.
- DRIFT OF LATE ITKILLIK AGE**—II—Stratified ice-contact deposits, as described above. Forms sharp-pointed end moraines and very irregular ground moraine and ice-stagnation deposits. In upper John and Itayvak valleys and along Hunt Fork and Fish Creek. Loess cover generally absent, and exposed tones very slightly weathered.



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SURFICIAL GEOLOGIC MAP OF THE WISEMAN QUADRANGLE, ALASKA

By
Thomas D. Hamilton
1979