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SEDIMENTOLOGY AND STRATIGRAPHY OF THE
KANAYUT CONGLOMERATE, CENTRAL BROOKS RANGE,
ALASKA—REPORT OF 1980 FIELD SEASON

By

T. H. Nilsen, T. E. Moore, D. F. Balin, and S. Y. Johnson

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This report is preliminary and
has not been reviewed for conformity
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standards and stratigraphic nomenclature.

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ABSTRACT

The Upper Devonian Kanayut Conglomerate crops out along the crest of the Brooks Range of northern Alaska for a distance of almost 1000 km. It ranges in thickness from 2600 m in the Atigun River area to 700 m south of Anaktuvuk Pass and has been subdivided into four regionally persistent members: (a) the basal sandstone member, consisting of marine sandstone and shale with some conglomerate; (b) the lower shale member, consisting of nonmarine quartzite, conglomerate and shale; (c) the middle conglomerate member, consisting of nonmarine pebble and cobble conglomerate and quartzite; and (d) the Stuver Member, consisting of nonmarine sandstone and shale. The Kanayut conformably overlies the Upper Devonian marine Hunt Fork Shale and is conformably overlain by the Mississippian marine Kayak Shale. The Kanayut is wholly allochthonous and has probably been transported northward on a series of thrust plates.

The basal sandstone member of the Kanayut Conglomerate, which overlies prodelta turbidites of the Hunt Fork Shale, contains marginal-marine coarsening-upward channel-mouth bar sequences. It is conformably overlain by the lower shale member.

Measured sections of the nonmarine members of the Kanayut show that the lower shale member ranges in thickness from 120 m to 1115 m and consists of fining-upward cycles interpreted to have been deposited by meandering streams on a broad floodplain. These cycles contain, in ascending order, channelized basal conglomerate, trough cross-stratified sandstone, and ripple-marked siltstone. The cycles are interpreted to be channel and point-bar deposits. Individual cycles average about 10 m in thickness and are separated by intervals of black, brown or maroon floodplain shale deposits. These typically contain thin coarsening-upward units that probably represent prograding levee sequences and irregular and ungraded sandstone bodies interpreted to be crevasse-splay deposits. In the Okokmilaga River area, the lower shale member contains a distinctive coarse-grained unit which is burrowed and interpreted to represent a widespread marine incursion.

The middle conglomerate member, which ranges in thickness from 155 m to 525 m, consists of braidplain deposits. It contains fining-upward couplets of conglomerate and parallel-stratified or cross-stratified sandstone that average 2-7 m in thickness. The couplets record deposition in channels and on bars of braided streams. The middle conglomerate member contains the largest clasts, little or no shale, and represents the maximum progradation of nonmarine sedimentation in the Kanayut depositional system.

The Stuver Member consists of fining-upward cycles that resemble those of the lower shale member. It ranges in thickness from 160 m to 1400 m and grades upward into tidal and marginal-marine deposits of the Kayak Shale.

Conglomerate in the Kanayut is compositionally very mature, averaging 82 percent white, gray, black or red chert clasts, 14 percent vein quartz clasts, 3 percent quartzite clasts, and less than 1 percent other lithologies, mainly argillite. Although red chert is locally abundant in the Shainin Lake-Galbraith Lake area, there is little variation in conglomerate composition in the Kanayut, suggesting derivation from a single major source terrane.

Range during the 1978 field season have been summarized by Nilsen and others (1980a). That report also contains a summary of the regional lithostratigraphy and biostratigraphy of the Endicott Group, and the reader is referred to that report for stratigraphic background information. Studies of the Kanayut and related units in the eastern and western Brooks Range during the 1979 field season have been summarized by Nilsen and others (1981a).

Four stratigraphic sections were measured in the Endicott Group during the 1978 field season (Nilsen and others, 1980a) and four during the 1979 field season (Nilsen and others, 1981a). In this report, we describe twelve additional measured sections, all from the Kanayut Conglomerate of the central Brooks Range.

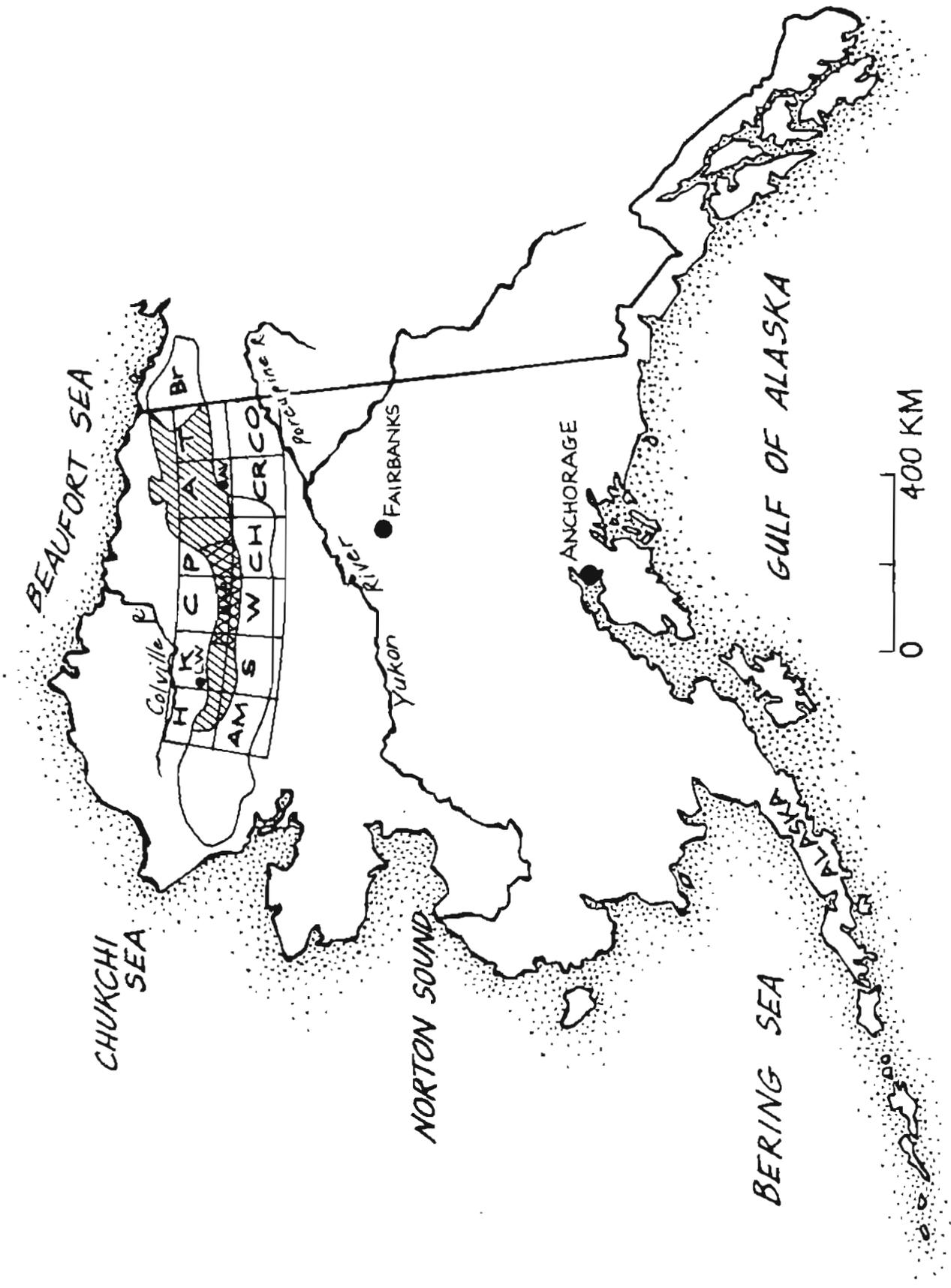
Stratigraphy

The Upper Devonian and Mississippian strata of the Brooks Range were thought to form a major offlap-onlap megacycle in which the Upper Devonian Kanayut Conglomerate and the Lower Mississippian Kekiktuk Conglomerate are the middle nonmarine part of the cycle (fig. 2). The Devonian Hunt Fork Shale comprises the lower marine part of the megacycle and the Mississippian Kayak Shale, or its local equivalent, the Itkilyariak Formation, comprises the upper marine part. The middle Devonian Beaucoup Formation (Dutro and others, 1979) underlies the Hunt Fork Shale and may represent an earlier depositional cycle.

Prior reconnaissance mapping had outlined the distribution of units of the Endicott Group. However, detailed mapping and division of the Kanayut Conglomerate into informal members were limited to areas near Shainin Lake and Anaktuvuk Pass (Bowsher and Dutro, 1957; Porter, 1966). Members of the Kanayut Conglomerate and of the Hunt Fork Shale were mapped in the Philip Smith Mountains quadrangle on the basis of field work done in 1975 and 1976 (Brosge and others, 1979a) and in most of the Chandler Lake and Killik River quadrangles on the basis of field work done in 1977 and 1978 (Brosge and others, 1979b). The mapped boundaries between members, based primarily on the percentage of resistant beds in the respective members, do not, however, necessarily coincide with boundaries determined from the detailed measurement of stratigraphic sections presented in this report.

The Hunt Fork Shale was defined by Chapman and others (1964) in its type area just west of the Okokmilaga Valley. The Kanayut Conglomerate and Kayak Shale were described by Bowsher and Dutro (1957) in the Shainin Lake area and by Porter (1966) near Anaktuvuk Pass. Bowsher and Dutro described and named the upper three members of the Kanayut, and Porter described the lowest member, which is absent at Shainin Lake. All four members were again described, with revised names, in the Philip Smith Mountains area (Brosge and others, 1979a).

The Kekiktuk Conglomerate, named by Brosge and others (1962) for outcrops west of Lake Peters in the northeastern Brooks Range, and the Itkilyariak Formation, named by Mull and Mangus (1972) for outcrops in the Sadlerookhit Mountains of the northeastern Brooks Range, form part of the autochthonous sequence of the Endicott Group. These units are not discussed further in this report.



Distribution and structure

The Upper Devonian Hunt Fork Shale and Kanayut Conglomerate crop out in an east-west-trending belt adjacent to the 68th parallel and are absent from areas to the north and south, where the Kekiktuk Conglomerate in outcrop rests unconformably primarily on deformed pre-Upper Devonian rocks. Outcrops of the Upper Devonian clastic rocks comprise a series of thrust sheets and are thus allochthonous (Brosge and others, 1979a, b), whereas outcrops of the Kekiktuk rest primarily on deformed basement and are generally thought to be autochthonous or parautochthonous (Nilsen, in press).

The Mississippian Kayak Shale stratigraphically overlies both the Kanayut Conglomerate and the Kekiktuk Conglomerate and is present throughout most of the Brooks Range. In a small area about 50 to 75 km west-northwest of Arctic Village, both the Kanayut and Kekiktuk are absent and the Kayak Shale rests directly on the Hunt Fork Shale (Nilsen and others, 1981).

Most of the Endicott Group in subsurface is only slightly deformed, whereas all outcropping strata are folded and faulted. The Upper Devonian clastic rocks of the Endicott Group form a series of stacked thrust plates in the central Brooks Range. The Kanayut Conglomerate in the central Brooks Range forms three plates that are bounded by south-dipping thrust faults 125 to 450 km long and 15 to 30 km apart, with local displacements of at least 5 to 15 km (Brosge and others, in press). Westward, the faults obliquely intersect the northern front of the Brooks Range and the thrust plates are folded into west-plunging anticlines that plunge out westward beneath the overlying plate. Eastward, most of the faults also die out within the range into overturned anticlines with little or no horizontal displacement. However, the Toyuk thrust of Porter (1966) extends across most, if not all, of the Kanayut outcrop belt. Because the plates north of the Toyuk thrust plunge out en echelon to the west, the eastern part of the plate overlying the Toyuk thrust forms the south edge of the Kanayut outcrop belt and its western part forms the north edge of the Kanayut outcrop belt at the front of the range.

A set of north-dipping thrust faults separates the Upper Devonian clastic rocks in the central Brooks Range from the southern belt of Kekiktuk Conglomerate outcrops. If these faults are continuous at depth with the faults that bound the Devonian rocks to the north, the entire outcrop belt of the allochthonous Devonian Endicott Group may be a giant klippe (Mull and others, 1976). However, the connection between these faults appears to be buried beneath subsequent thrust plates on the west and has not yet been found where it should be exposed at the east end of the presumed klippe (Brosge and others, in press).

The Upper Devonian part of the Endicott Group is not known for certain to have been reached in the subsurface, but equivalent rocks that have been defined by seismic profiles (Tetra Tech, Inc., 1979) to the northwest may be present locally in the deeper parts of the Umiat and Meade Basins of the western North Slope area. The Endicott Group may be as thick as 2,800 m in these deeply buried basins because, although seismic interpretations (Tetra Tech, Inc., 1979) indicate a thickness of less than 1,000 m of strata between the base of the Carboniferous Lisburne Group and argillite basement over most of the area, as much as 4,000 m is present in the Umiat Basin. At the Inigok test well near the edge of the Umiat basin, the Lisburne Group is about 1,200 m thick, the Kayak Shale 80 m thick, and the Kekiktuk Conglomerate at least

settling from storm-generated overflows or interflows within the water column rather than bottom-flowing turbidity currents.

Cycles of shallow-marine sandstone characterize the upper part or wacke member of the Hunt Fork Shale of Brosge and others (1979a, b) and the basal sandstone member of the Kanayut Conglomerate. These sandstone units are locally fossiliferous and contain assemblages characteristic of shelf, intertidal, lagoonal, and marginal-marine environments. The sandstone strata, deposited as a variety of depositional bodies, most typically form thickening- and coarsening-upward cycles, especially in the the basal sandstone member of the Kanayut Conglomerate. The thickening-upward cycles record progradation of the delta over marine slope and shelf deposits. Although the geometry of these deposits has not been ascertained, the vertical sections resemble those of channel-mouth bars described from many modern deltas.

Other bodies of shallow-marine sandstone, characterized throughout by medium- to large-scale cross-strata that have consistent orientations, may represent offshore bars or spits that formed on the margins of the delta. Variable paleocurrents directions from the shallow-marine sandstone bodies reflect currents generated by waves, winds, tides, and longshore drift.

Kanayut Conglomerate

The lower shale and Stuver members of the Kanayut Conglomerate are interpreted to have been deposited by meandering streams on a floodplain. The members consist of a series of thinning- and fining-upward cycles similar to those described from many modern meandering rivers. Detailed analysis of the cycles may permit determinations in the future of the sizes of the rivers, their discharges, and temporal or geographical variability of the streams within the depositional basin.

The meandering-stream cycles characteristically commence at their base with erosional truncation of underlying shale or paleosols by thick beds of conglomerate or sandstone. The amount of downcutting observed varies from several cm to as much as 5 m. However, if viewed on a large enough scale, each fluvial cycle probably downcuts approximately the thickness of the individual cycle, inasmuch as the cycle is a preserved record of filling of the individual river channel. The basal beds typically consist of massive or crudely parallel-stratified conglomerate or conglomeratic sandstone containing abundant rip-up clasts of shale, siltstone, and paleosol material.

Overlying the basal conglomeratic beds are parallel-stratified beds of sandstone that are in turn overlain by trough-cross-stratified beds of sandstone. Trough amplitudes gradually decrease upward in the cycles concomitantly with decreasing grain size of the sandstone. The trough axes have variable attitudes, but generally plunge toward the southwest. These deposits represent fill of the channel by transverse and longitudinal bars that migrate downchannel as the stream channel gradually shifts and migrates laterally by the meander process.

The upper part of the cycles consists of thinly bedded current-ripple-marked fine-grained sandstone with thin shale interbeds. These ripple-marked sandstones contain abundant mica, clay, and carbonaceous material. Climbing ripples are locally common in these deposits, as well as plant fossils and

Kayak Shale

In its type area, near Shainin Lake, the Kayak Shale is about 300 m thick, rests conformably on nonmarine facies of the Stuver Member, and has been subdivided into 5 members: 1. basal fine-grained sandstone, 40 m thick; 2. lower black shale, 180 m thick; 3. argillaceous limestone, 24 m thick; 4. upper black shale, 40 m thick; and 5. red limestone, 5 m thick. The three lower members, although their thickness varies, can be traced along the entire Brooks Range, despite marked thickening and thinning of the Kayak as a whole from the effects of thrust faulting. However, in the southern and eastern Brooks Range, the Kayak Shale in the allochthonous sequence of the Endicott Group is generally less than 75 m thick and consists mostly of black shale.

The basal sandstone member typically consists of thinly cross-stratified and ripple-marked fine-grained quartzose sandstone with abundant Scolithus burrows, a marine ichnofossil. In its type area and elsewhere, it contains both current- and oscillation-ripple markings, reactivation surfaces, flaser bedding, and herringbone cross laminae and cross strata suggestive of tidal current activity. In most localities within the allochthonous sequence of the Endicott Group, the sandstone rests directly on red or black shale or soils of the topmost fining-upward fluvial cycle of the Stuver Member; however, in contrast to coarse-grained deposits of the Kanayut, it consists wholly of fine-grained sandstone, may contain recumbent and syndepositionally folded cross strata, slump folds, and slurried layers, and may form a thickening- or coarsening-upward megasequence abruptly overlain by black shale of the next member. In a few rare sections, the basal sandstone member forms a very thick cycle of intermixed fine-grained sandstone typical of the basal member and conglomerate of the fluvial Kanayut Conglomerate.

The overlying black shale contains some thin graded beds of fine-grained sandstone that appear to be either turbidites or vertical accumulations of storm-generated sediment overflows. The argillaceous limestone, which appears to form a number of different beds of variable thickness in different areas, contains megafossil debris, including brachiopods, bryozoans, echinoderms, mollusks, and ostracodes. Preliminary examination of these beds suggest a debris-flow origin, because they rest abruptly on and are overlain abruptly by black shale, they are ungraded, have a partly argillaceous matrix, and contain unsorted and chaotic assemblages of calcareous fossil fragments.

The Kayak Shale in general represents a sequence that was deposited in progressively deeper water, except at its top, where it shoals upward into limestone of the Lisburne Group. The basal sandstone represents nearshore, probably tidal sand flat-deposition. Paleocurrent directions from it are highly variable and indicate flow toward the southwest, southeast, and northeast (Nilsen et al., 1980b). The overlying black shale represents deeper marine sedimentation, probably a prodelta slope setting, into which some massive fossiliferous debris flows of argillaceous limestone were resedimented.

Conglomerate. The stratigraphically highest marine strata are 590 m above the base of the section.

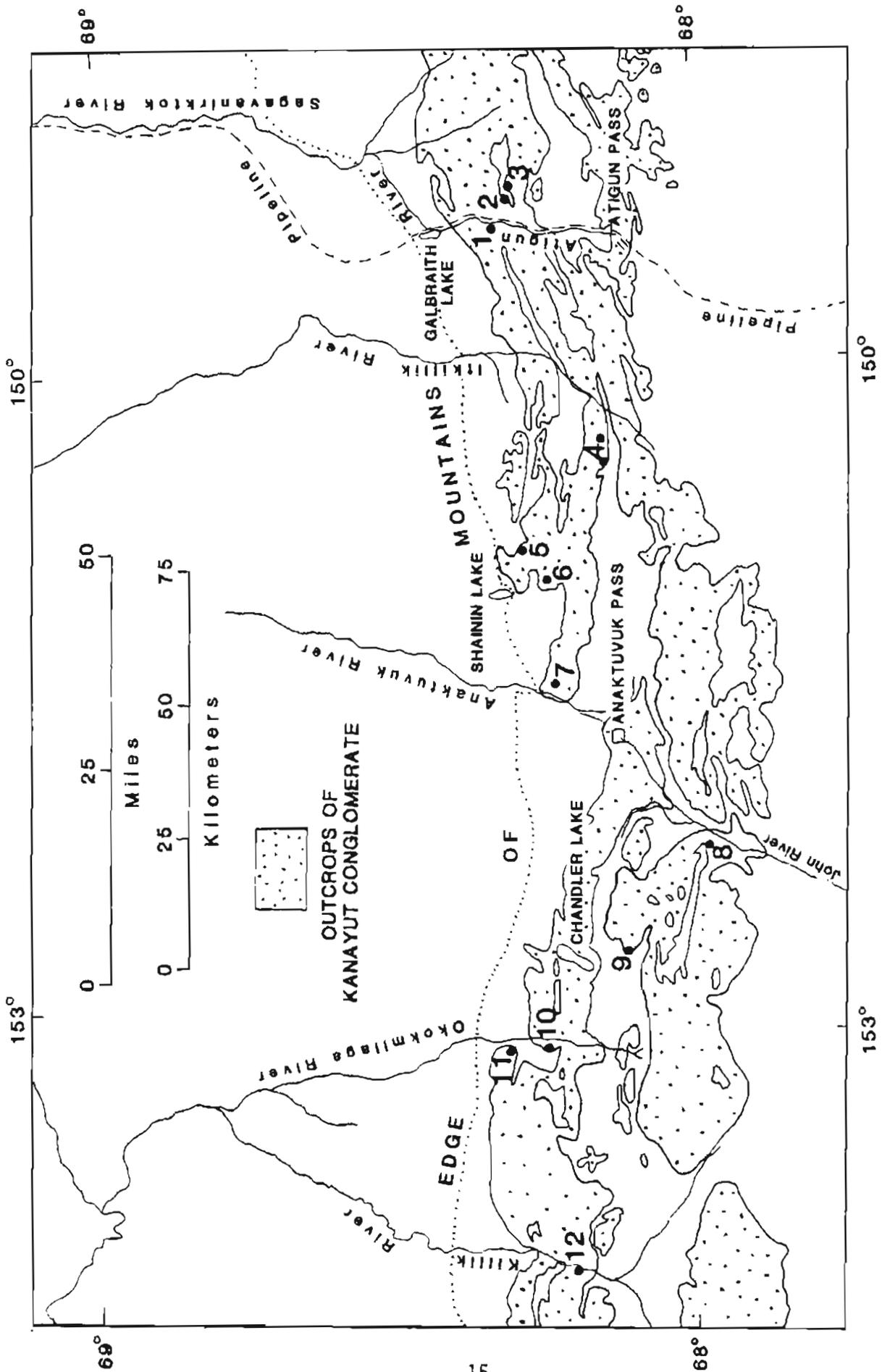
The lowest part of the lower shale member is marked by two prominent fining-upward cycles at 253 m and 262 m that contain conglomeratic sandstone at their bases. The cycles near the base of the member, between 253 m and 405 m, typically contain fine- to medium-grained trough cross-stratified sandstone at their base and are overlain by parallel-stratified, very fine-grained sandstone, siltstone and shale. The fine-grained tops of some cycles are thin and contain Scolithus-type burrows, indicating a marine origin. Two cycles between 300 m and 380 m contain thick accumulations of interstratified ripple-marked fine-grained sandstone and shale that lack burrows and may represent either fluvial floodplain or freshwater lake deposits. We infer that these interbedded marine and fluvial strata near the Hunt Fork-Kanayut contact record a gradual progradation of coarse-grained delta-plain meandering streams over fine-grained delta-front sediments. The complex interfingering probably reflects vertical and lateral accretion of meandering-fluvial-channel, interdistributary-bay, brackish-swamp, low-energy-shoreline, and freshwater lake deposits.

The middle and upper parts of the lower shale member above 405 m consist of thinning- and fining-upward cycles inferred to be meandering river deposits. These cycles contain both coarser-grained deposits than those at the base of the member and thicker basal sandstone and conglomerate units that exhibit multiple scour horizons and local trough cross-stratification. The coarse basal facies of each cycle probably represent channel and lower point-bar deposits. The basal deposits are typically overlain by trough cross-stratified, fine- to medium-grained sandstone capped by red, maroon, or black siltstone and shale. These fine-grained sediments were probably deposited on the upper surfaces of point-bars, levees, and floodplains. Shale and siltstone intervals are thin or nonexistent in the middle of the lower shale member between 720 m and 890 m, but increase in thickness upward toward the top of the section. Near the top, the floodplain deposits are as thick as 20 m and commonly rest directly on conglomeratic strata without intervening sandstone of substantial thickness. One well-exposed floodplain sequence near the top, between 1350 m and 1365 m, contains two cycles of black shale, thin rippled-marked sandstone, and a paleosol. These thin coarsening-upward cycles are 2-3 m in thickness and probably record outbuilding of levees into adjacent swampy lowlands or lakes on the floodplains.

The contact with the overlying middle conglomerate member at the top of the section is abrupt and marked by the disappearance of shale and appearance of massive sandstone and conglomeratic strata. This contact reflects the influx of coarser-grained sediment and transition to deposition by braided streams.

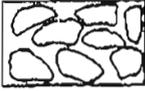
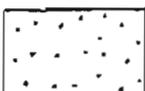
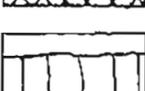
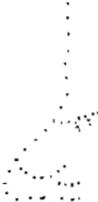
Cycles are generally thicker at the base of the lower shale member, thinnest in the middle part, and thickest at the top. The cycles average 20-30 m in thickness at the base of the member, 10-20 m in the middle, and 45 m at the top.

Clast size increases regularly upward. In the Hunt Fork Shale and at the base of the lower shale member, conglomerate is rare and the maximum clast size is 2 cm. At the top of the lower shale member, the maximum clast size is

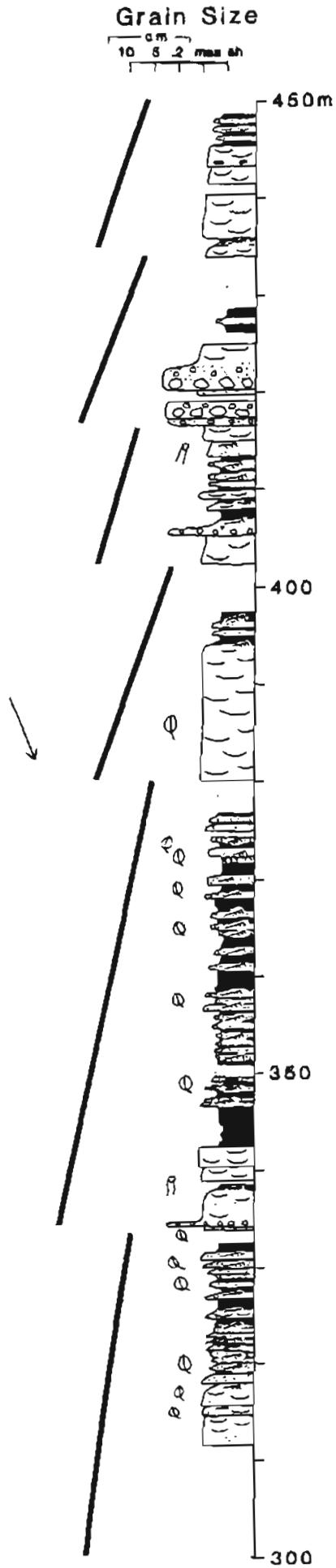


SYMBOLS

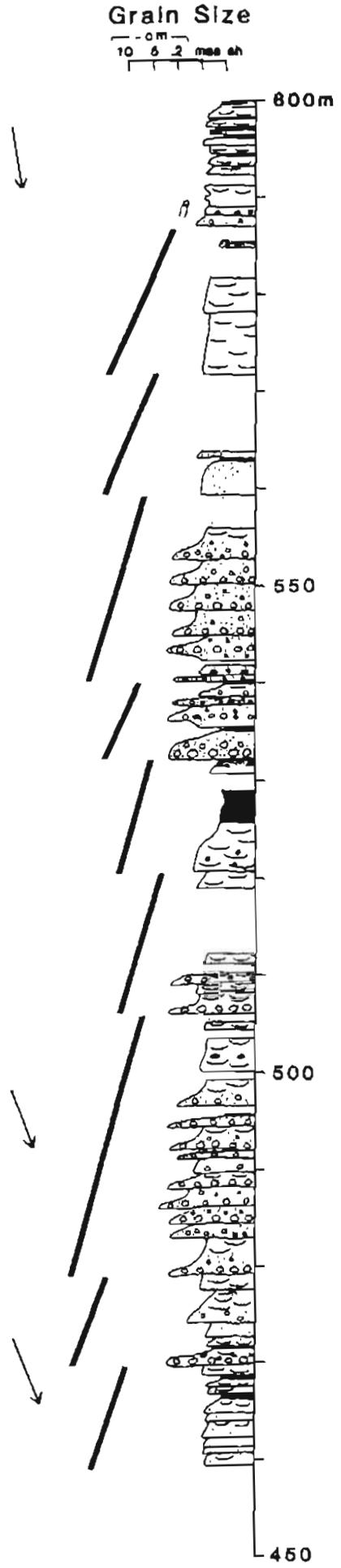
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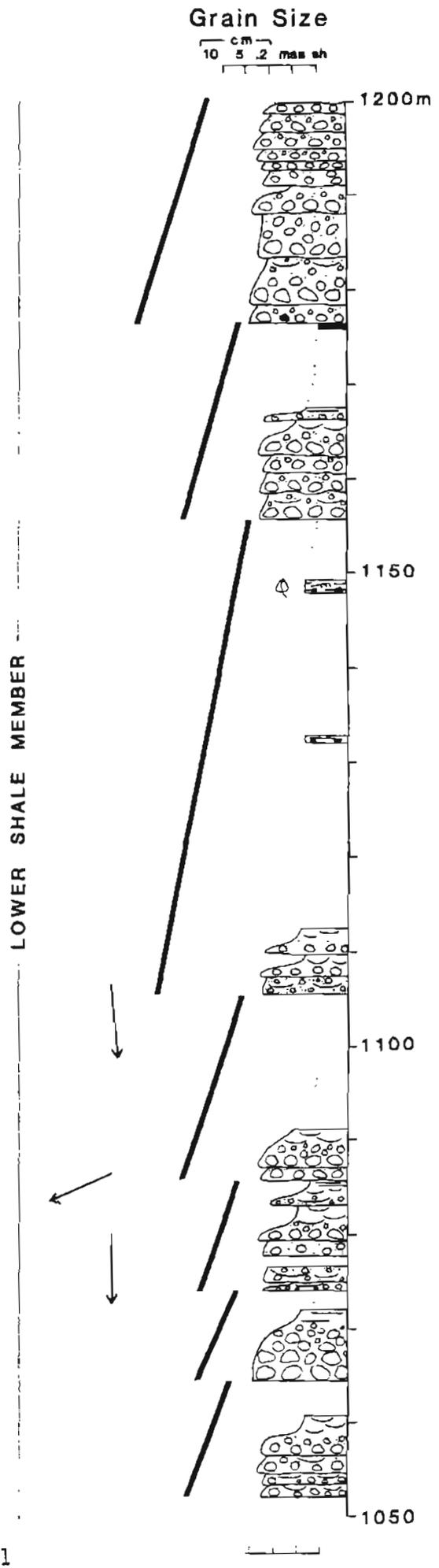
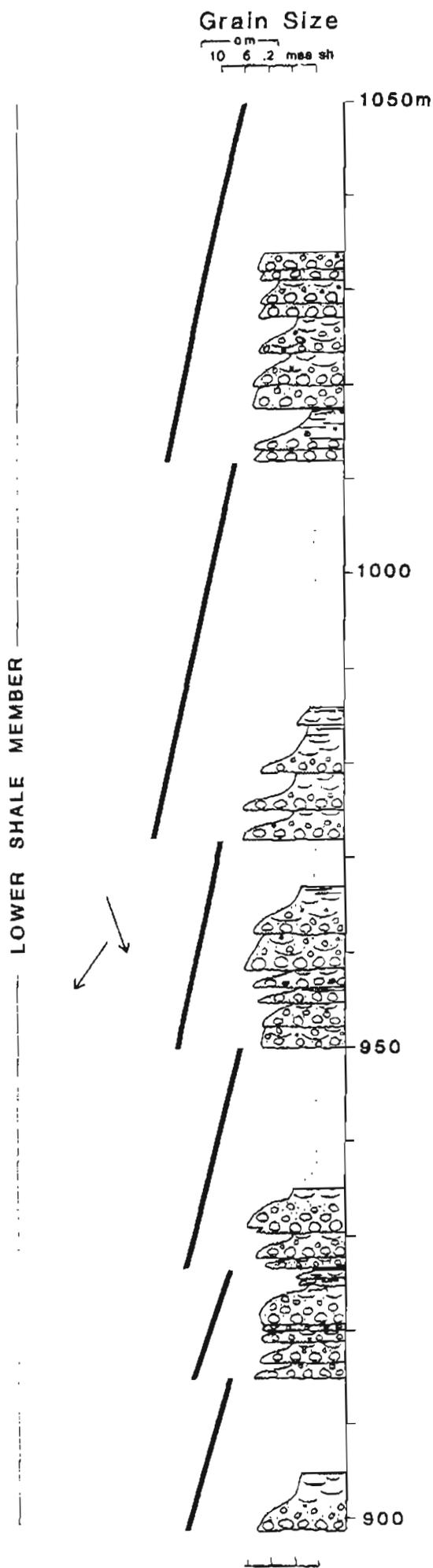
	Conglomerate		
	Conglomeratic sandstone		
	Shale or siltstone rip-up clasts		
	Sandstone, massive		
	Sandstone, parallel-stratified		
	Sandstone, trough cross-stratified		
	Sandstone, tabular cross-stratified		
	Epsilon cross-stratification		
	Siltstone and shale		
	Coal		
	Paleosol		
	Current ripple marks		
	Oscillation ripple marks		
	Sand waves		
	Mudcracks		
	Syndepositional slump fold		
			Float lithology
			Fining-upward cycle
			Coarsening-upward cycle
			Paleocurrent azimuth
			Burrow, invertebrate
			Invertebrate fossil
			Plant fossil
			Root cast

LOWER SHALE MEMBER



LOWER SHALE MEMBER





5 cm. Two pebble counts from the lower shale member show that the clast composition is typical of the Kanayut, averaging 85 percent chert and 15 percent vein quartz. Quartzite is absent and red chert comprises 2 percent of the conglomerate clasts.

Paleocurrent measurements from the section are consistent in orientation and indicate southward sediment transport. There is little difference between directions from the Hunt Fork Shale and the lower shale member of the Kanayut Conglomerate. The mean and standard deviation of 8 measurements from the Hunt Fork Shale are $186^{\circ} \pm 21^{\circ}$ and of 17 measurements from the lower shale member are $171^{\circ} \pm 38^{\circ}$ p.

Middle conglomerate member

A complete section of the middle conglomerate member of the Kanayut Conglomerate was measured on a steep east-facing valley wall approximately 15 km southeast of Galbraith Lake, (NE 1/4, T.13S., R.12E., Philip Smith Mountains Quadrangle). The measured section (fig. 5) is very well exposed along a small northeast-trending creek. The middle conglomerate member at this location is 155 m thick and is conformably underlain and overlain by the lower shale and Stuver members, respectively. The contacts between the members are abrupt and marked by the absence of shale in the middle conglomerate member. The basal 12 m (1 cycle) of the Stuver Member were also measured at this location, resulting in a total thickness of 167 m for the entire section. The middle conglomerate member at this location was previously mapped by Brosge and others (1979a) as the massive marker bed within a thicker and lithologically more variable middle conglomerate member.

The measured section consists almost exclusively of interstratified conglomerate, conglomeratic sandstone, and coarse-grained sandstone. Fine-grained sandstone, siltstone and shale are present only in the basal part of the Stuver Member at the top of the section. Beds of conglomerate are lenticular, as thick as 7.3 m, and contain clasts as large as 10 cm. Beds of sandstone are as thick as 1.4 m and are commonly conglomeratic, containing pebbles as large as 3 cm.

The conglomerate and sandstone beds form fining-upward couplets in which the conglomeratic base of each couplet is channeled into the underlying sandstone or conglomeratic sandstone. The conglomerate is clast-supported, typically massive to crudely parallel-stratified, and has a well-defined fabric characterized by upstream dip of imbricated pebbles and orientation of long axes parallel to flow. The matrix consists of sandstone and finer conglomeratic sandstone. Beds of sandstone are generally parallel-stratified to trough cross-stratified but are locally massive. They form thin, flat lenses in outcrops that are truncated above by beds of conglomerate. In the interval between 100 m and 150 m above the base of the section, conglomerate beds are commonly overlain by a scoured surface and another conglomerate bed, resulting in a sequence of amalgamated conglomerate. This interval contains the coarsest conglomerate. The section above 150 m changes abruptly in character. Two fining- and thinning-upward cycles can be recognized, although partly covered, by the fine-grained sandstone, siltstone, and shale which is present in both float and outcrop in their upper parts. The overlying conglomerate is scoured into these fine-grained deposits and locally contains well-developed flute casts at its base.

The lower 155 m of the section was probably deposited by braided streams. Fining-upward conglomerate-sandstone couplets probably were formed by lateral and vertical accretion of migrating longitudinal bars. Migrating dunes and sandwaves superimposed on larger bars resulted in crossbedded sandstone. The transition to the thicker fining-upward cycles of the Stuver Member, which include siltstone and shale in their upper parts, may mark a change to deposition by a meandering river system. The coarse basal parts of these thicker cycles are interpreted to be river-channel and point-bar deposits, and the upper fine-grained strata upper-point-bar, levee, and floodplain deposits.

Two pebble counts indicate an average composition of 88 percent chert, 11 percent vein quartz, and 1 percent quartzite. Red chert clasts are notably abundant, composing approximately 40 percent of the total number of clasts.

Four paleocurrent measurements of trough cross-strata axes, clast imbrication, and clast long axes indicate that sediment transport was consistently southward during deposition of the middle conglomerate member. Flute casts at the top of the section are oriented toward the west, however, possibly suggesting a change to more westerly sediment transport during deposition of the basal strata of the Stuver Member. The mean and standard deviation for all seven paleocurrent measurements from this section are $188^{\circ} \pm 43^{\circ}$.

Stuver Member

A thick section of the Stuver Member of the Kanayut Conglomerate was measured on a north-trending ridge about 6 km east of the Atigun River Valley (sec. 13, T.13S., R.13E., Philip Smith Mountains Quadrangle). Contacts with the underlying middle conglomerate member and overlying Kayak Shale are well exposed (Fig. 6).

Nearly 1,300 m of section were measured within the Stuver Member, with less than 1 percent covered. The contact with the middle conglomerate member is marked by the termination of massive conglomerate beds and beginning of distinct fining-upward cycles that contain shale intervals that are 1 to 2 m thick at the base of the Stuver Member. The cycles commence with beds of conglomerate or conglomeratic sandstone that have a maximum clast size of 7 cm, compared to 8.5 cm at the top of the middle conglomerate member. These generally massive beds have an erosional contact with underlying strata and may contain shale rip-up clasts as large as 1 m. The coarse-grained beds grade upward into trough-cross-stratified and current-rippled-marked sandstone that is overlain by extensive sequences of interbedded shale, siltstone, and very fine-grained sandstone as thick as 55 m.

The fining-upward cycles and the overall fining-upward pattern of the Stuver Member are indicative of meandering stream deposits and gradual retrogradation of the Kanayut delta. The conglomerate and conglomeratic sandstone at the base of the cycles represent channel deposits, grading upward into point-bar deposits represented by the trough-cross-stratified and ripple-marked sandstone. Floodplain, crevasse-splay, and levee deposits are represented by the thick sequences of shale, siltstone, and fine-grained sandstone at the tops of the cycles.

cycles at additional places on the east-west trending ridge, but bad weather permitted us time to measure only sections A and B.

Section A is about 36 m thick and contains three major individually distinctive major fining-upward cycles. The lower cycle, about 13 m thick, is most clearly fining- and thinning-upward in character. It contains about 7 m of conglomerate and sandstone with a maximum clast size of 3 cm in its lower part and 6 m of maroon and black shale in its upper part. Two prominent beds of fine-grained sandstone are present in the floodplain deposits and probably represent crevasse-splay deposits.

The middle cycle of section A, about 11 m thick, contains a thick basal amalgamated bed of conglomerate that is about 2 m thick and has a maximum clast size of 2 cm. The bed coarsens upward in its lower 30 cm and then fines abruptly upward into 8 m of maroon and black shale that contains three thin ripple-marked and laminated beds of very fine-grained sandstone.

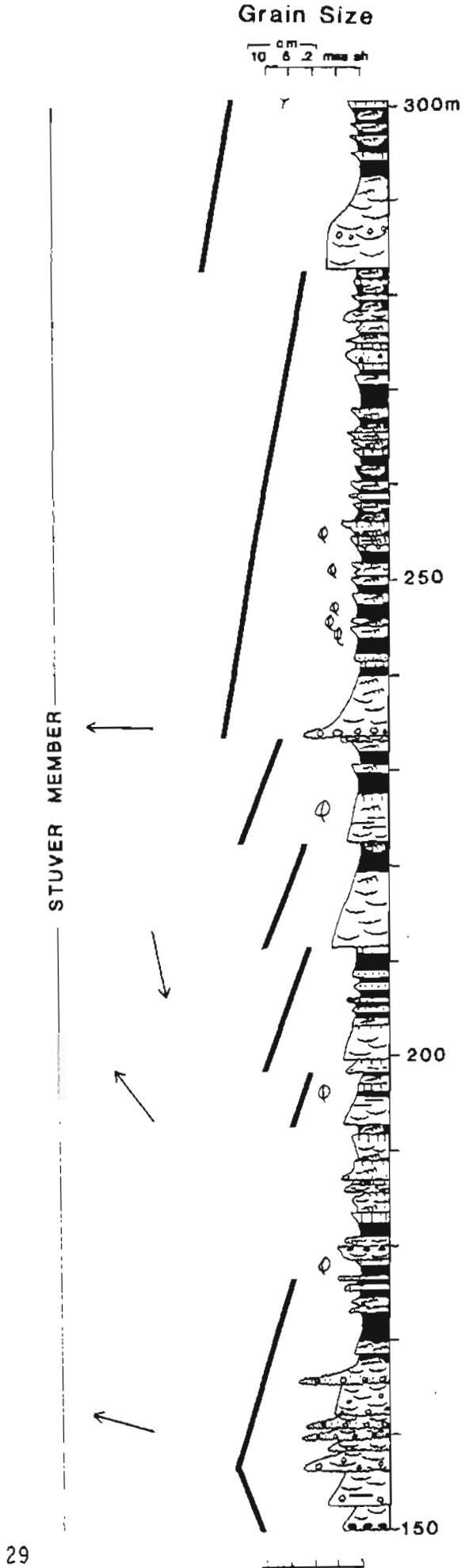
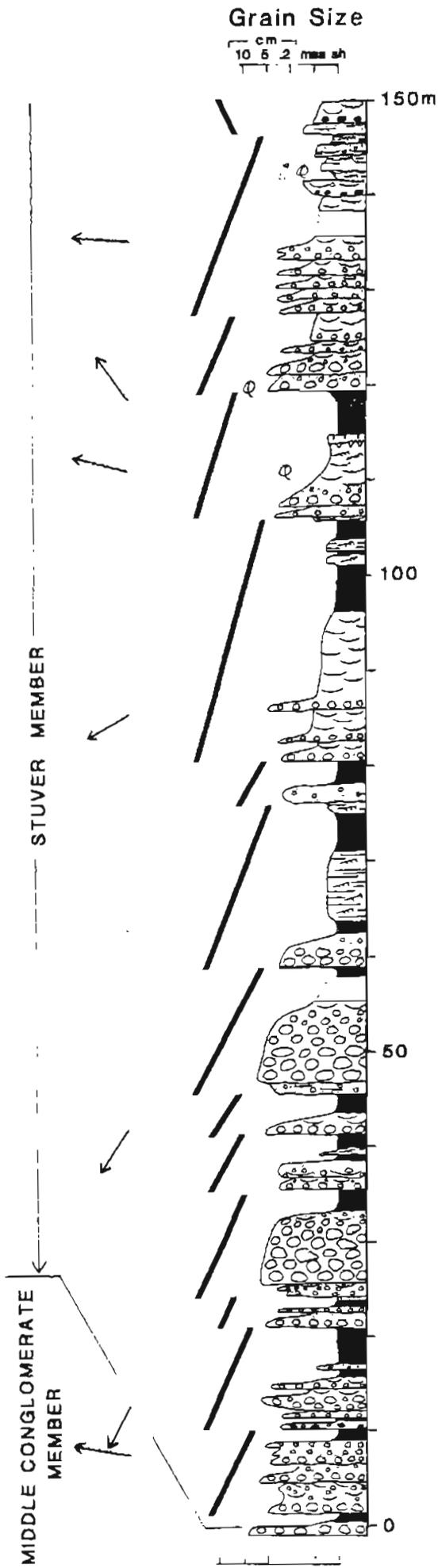
The upper cycle of section A is about 12 m thick and more complex in its internal organization. It coarsens upward over a 5 m thickness before abruptly fining upward to red shale and green siltstone. The maximum clast size is 3.5 cm in the lower thinner beds of conglomerate and 4 cm in the upper thicker bed. The floodplain sequence consists of only 1.5 m of shale.

Westerly transport of sediments is suggested on the basis of two paleocurrent measurements from the lower cycle, none from the middle cycle three from the upper cycle, and one from the overlying cycle (fig. 7). The measurements include three trough-cross strata, two clast imbrication and long-axis orientations, and one flute cast.

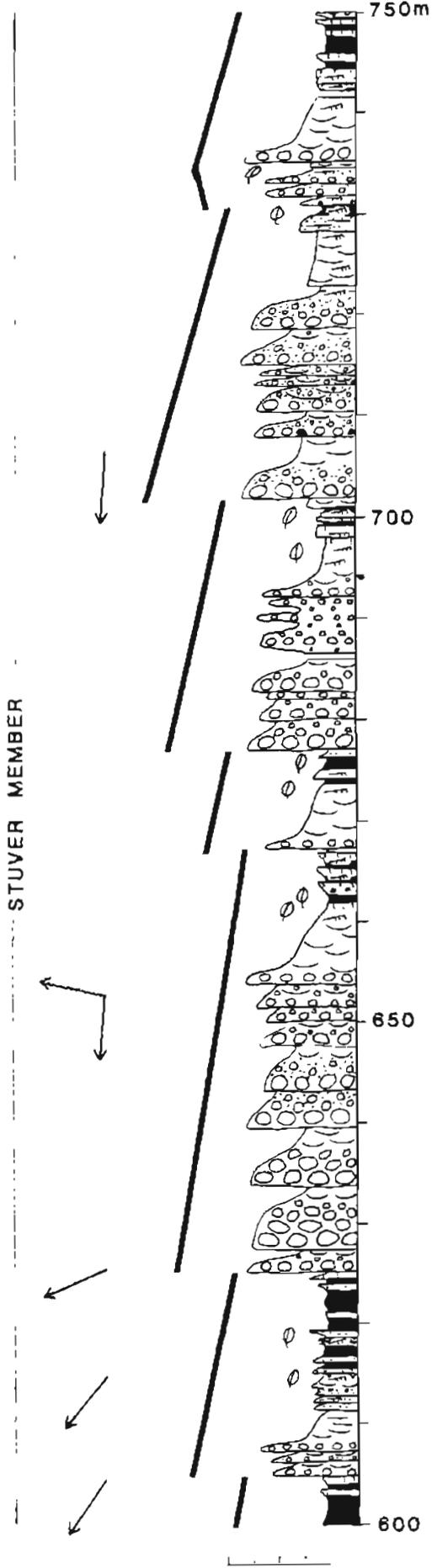
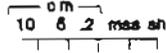
Section B is about 52 m thick and contains what we believe are the same three fining- and thinning-upward cycles. The cycles are relatively similar in character, with the lower cycle having the most clearly developed fining-upward character, the middle cycle being the thinnest and consisting mostly of a single coarse-grained bed, and the uppermost cycle being most complex, characterized in part by a coarsening- and thickening-upward character of the conglomeratic part of the cycle. No paleocurrents were obtained from the cycles in section B.

The lower cycle of section B is about 26 m thick and has a maximum clast size of 9 cm. It consists of a lower 10 m of conglomerate and sandstone and an overlying 16 m of maroon shale that contains three prominent interbeds of sandstone and soil horizons. The lower interbedded unit, about 12 m above the base of the section, consists of trough-cross-stratified, fine- to medium-grained sandstone and may be a crevasse-splay deposit. The two upper interbeds consist of fine-grained sandstone and an orange paleosol.

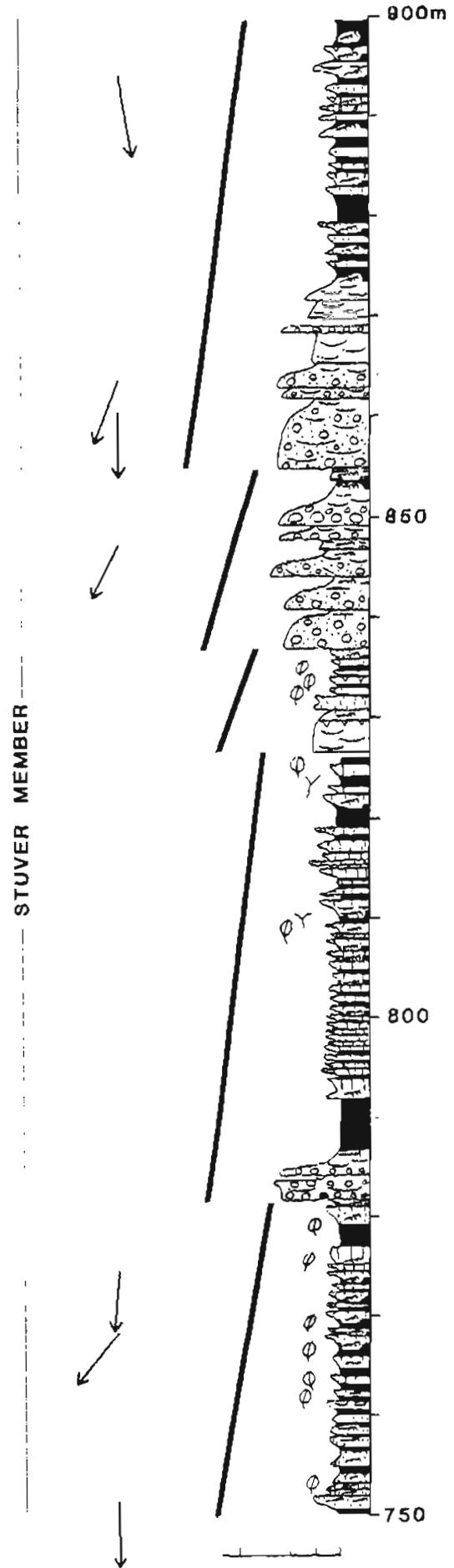
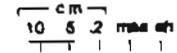
The middle cycle of section B is about 12 m thick. It consists of a basal medium-grained sandstone that coarsens upward to conglomeratic sandstone with a maximum clast size of 0.5 cm. The coarse-grained basal part of the cycle is about 2 m thick. The overlying maroon shale floodplain sequence is about 10 m thick, and contains several soil horizons and thin interbedded units of fine-grained, ripple-marked sandstone.

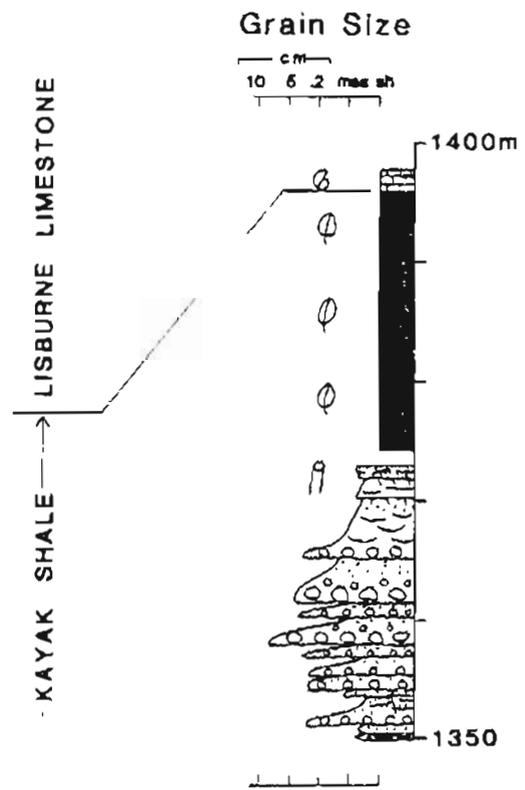
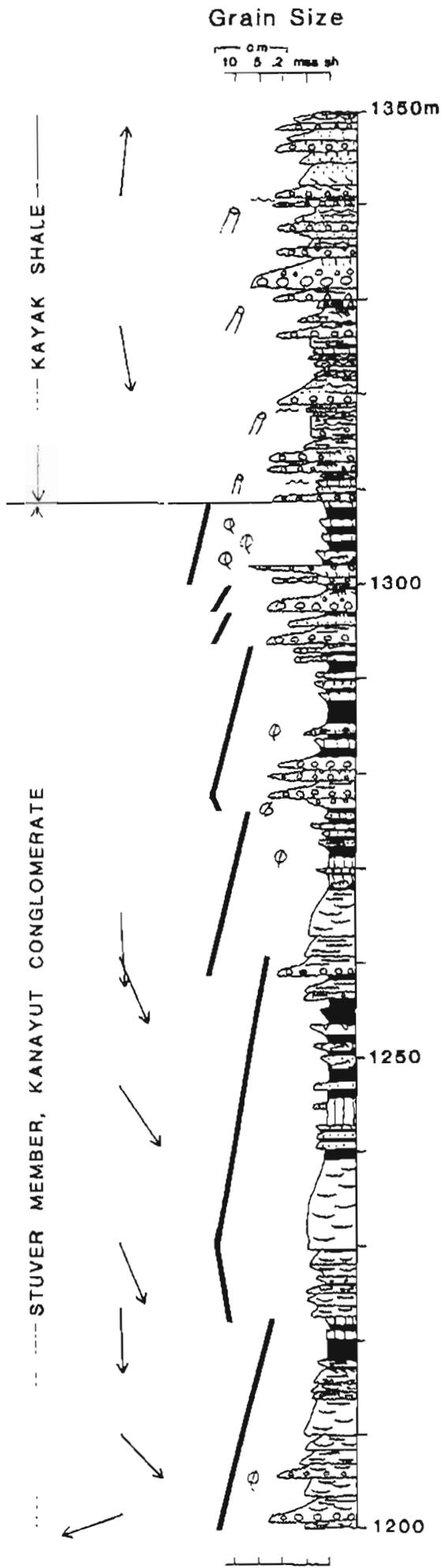


Grain Size



Grain Size





The upper cycle of section B is about 14 m thick, but contains only 0.5 m of shale at its top. The lower 6 m consists of thinly interbedded conglomerate, conglomeratic sandstone, and sandstone having a maximum clast size of 3 cm. The average bed of conglomerate in this part of the cycle is about 40 cm thick and the sandstone interbeds are characterized by trough cross-strata. The upper 9 m of the cycle contains a maximum clast size of 7 cm and an average conglomerate bed thickness of about 1.1 m, indicative of the coarsening- and thickening-upward nature of the main part of this cycle. The top of the cycle grades abruptly from conglomerate with clasts as large as 4.5 cm to cross-stratified medium-grained sandstone that is 0.25 m thick to maroon shale.

The east-to-west changes within the three cycles are complex. For the lower cycle, the thickness changes from 13 to 26 m, maximum clast size from 3 to 9 cm, and conglomerate-plus-sandstone to shale ratio from 1:1 to 1:2.5. We thus have a westward increase in thickness, maximum clast size and amount of shale, which is not wholly consistent with expected changes in the direction of sediment transport. For the middle cycle, the thickness stays the same but the maximum clast size decreases westward and the conglomerate-plus-sandstone to shale ratio diminishes, which is what one would expect from westward transport. For the upper cycle, the thickness increases from 12 to 14 m, the maximum clast size increases from 4 to 7 cm, and the conglomerate-plus-sandstone to shale ratio increases slightly from west to east.

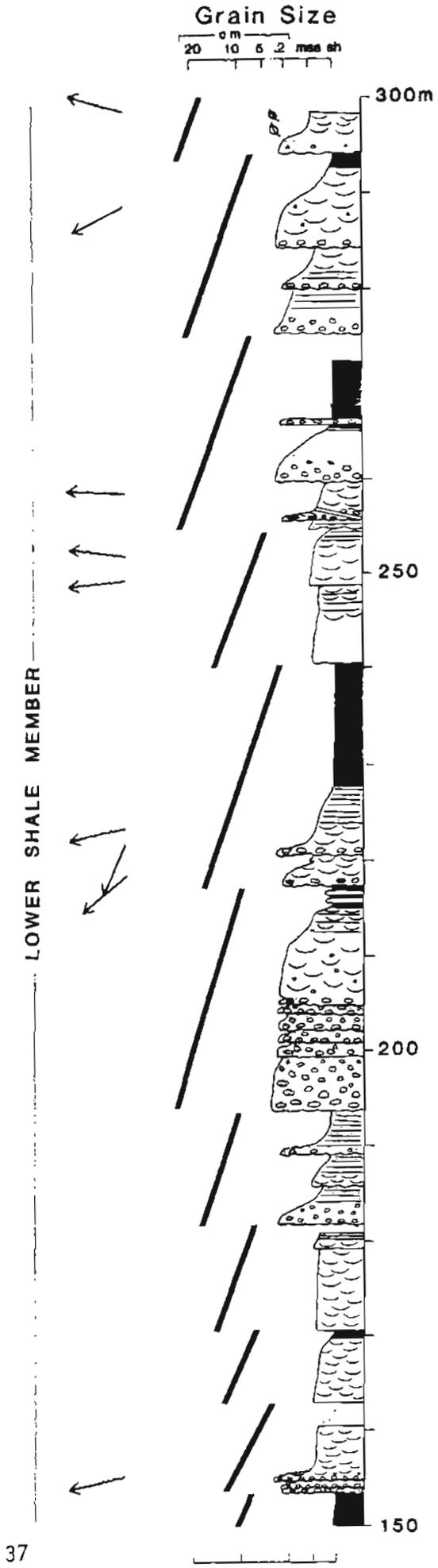
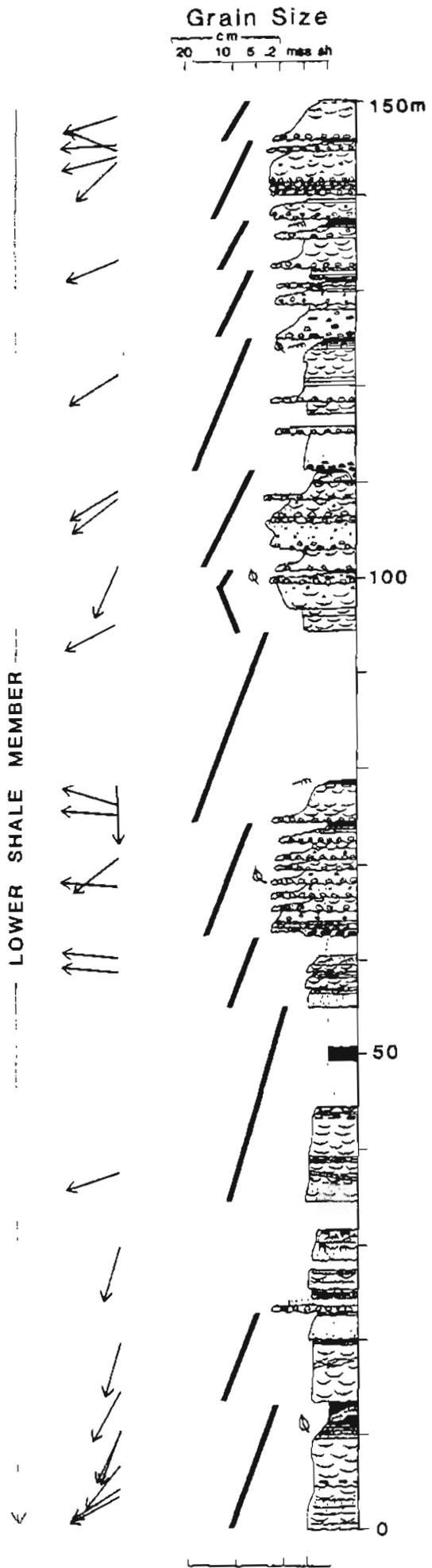
The two sections thus show variable downcurrent changes in cycle thickness and character, clast size, and conglomerate-plus-sandstone to shale ratios. No simple explanations for these changes, in terms of westerly transport of sediment by streams, is apparent. The possible overall transport of sediment toward the southeast in this area, as suggested by maximum clast size and regional paleocurrent data, may explain some of the variability. Additional studies of this type, especially involving three-dimensional analyses of cycle variability, are needed to understand major aspects of sediment distribution of the Kanayut depositional system.

Shainin Lake area

Lower shale member

A complete section of the lower shale member of the Kanayut Conglomerate was measured along the southernmost spur of an east-facing ridge near Ear Peak, east of Shainin Lake (sec. 13, T.13S., R.5E., the Chandler Lake Quadrangle). The section overlies the Hunt Fork Shale and underlies the middle conglomerate member of the Kanayut Conglomerate (fig. 8); the basal sandstone member was not recognized at this location by either Brosge and others (1979b) or us.

The section is 510 m thick and consists of thirty-five major fining-upward cycles that average 15 m in thickness. About 20 percent of the section is covered. The entire sequence generally coarsens upward as it approaches the base of the middle conglomerate member. Cycles typically begin with 1 to 5-m-thick beds of conglomerate or conglomeratic sandstone overlain by trough-cross-stratified or flat-stratified sandstone capped in some cycles by ripple-marked siltstone and red shale. The base of individual cycles may be marked by erosional scour into the underlying unit. The amount of shale, siltstone,



and fine-grained sandstone decreases upsection as the amount of coarse-grained sandstone and conglomerate increases. Rare burrows and plant fossils are present within beds of shale and fine siltstone in the lower half of the section.

The boundary of the lower shale member with the underlying Hunt Fork Shale is transitional and marked by the disappearance of marine megafossil debris and the initiation of distinct fining-upward cycles. The first distinct fining-upward cycle, at an elevation of 4820 feet (1610 m) forms the boundary between the two units. We interpret the boundary to mark the change upward from marine-dominated deposition to fluvial-dominated deposition. The distinct fining-upward cycles result from lateral and vertical accretion of meandering-stream channel and point-bar deposits. Epsilon cross-strata form inclined surfaces transecting the vertical sequence and may represent the actual inclined surfaces of point-bars within the system. Upper point-bar and floodplain deposits consist of shale and ripple-marked siltstone, with thin interbeds of fine-grained sandstone possibly representing crevasse-splay deposition.

The clast composition approximately 320 m above the base of the section is 87 percent chert, 11 percent quartz, and 2 percent quartzite. The clast composition 50 m below the base of the middle member consists of 82 percent chert, 13 percent quartz, 4 percent quartzite, and 1 percent argillite. The maximum clast size ranges from less than 1 cm near the base of the lower shale member to a maximum of 13 cm just below the base of the middle member.

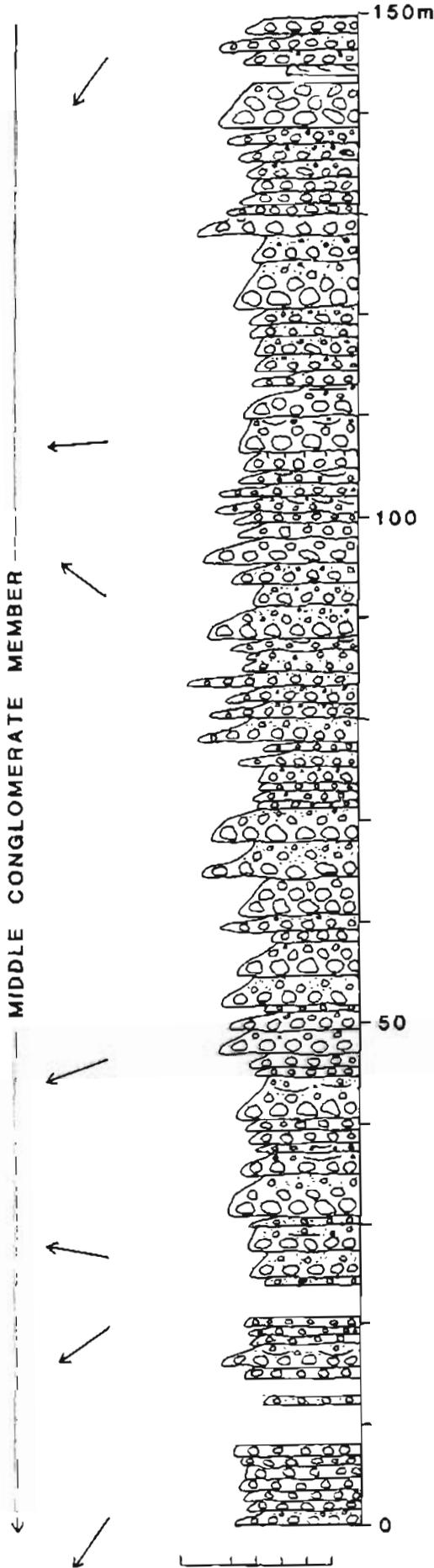
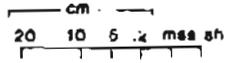
Forty-four paleocurrent measurements from the lower shale member have a vector mean and standard deviation of $243^{\circ} \pm 29^{\circ}$. The measurements include 19 trough cross-strata, 17 primary current lineations, 3 flute casts, 2 tabular cross-strata, and 2 sites of clast long-axis orientation, and 1 oscillation ripple marking. Because the unidirectional indicators have a direction of sediment transport toward the southwest, bidirectional indicators have been preferentially assigned a similar transport orientation.

Middle conglomerate member

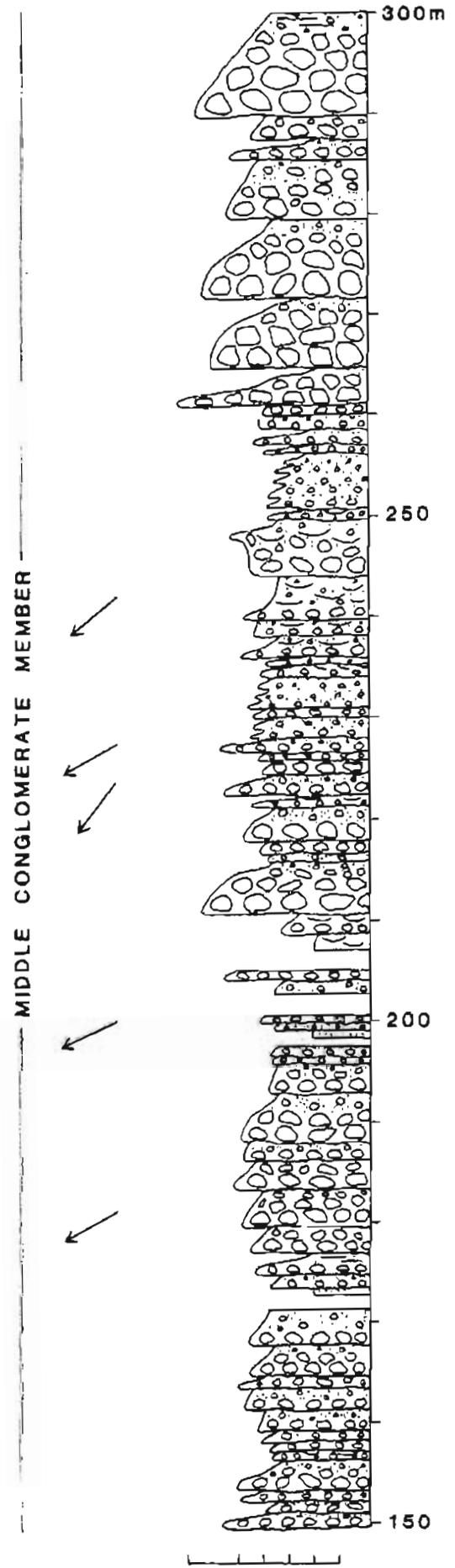
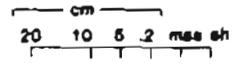
A complete section of the middle conglomerate member of the Kanayut Conglomerate was measured on the prominent cliff directly south of the intersection of Kayak Creek and Alapah Creek, about 8 km south-southeast of Shainin Lake (sec. 33, T.13S., R.5E., and sec. 5, T.14S., R.5E., Chandler Lake Quadrangle). The section is 526 m thick and is almost completely exposed, with relatively little cover (fig. 9). Almost all covered intervals appear to be sandstone. The section is near the one measured by Bowsher and Dutro (1957).

The section is relatively uniform, consisting of thick beds of conglomerate that fine upward to conglomeratic sandstone and medium- to very coarse-grained sandstone. Shale is generally absent except for some thin intervals from 450 m to 490 m above the base of the section. The coarsest and most thick-bedded conglomerate is present in the middle part of the section, from 260 m to 440 m above the base. Thinner bedded and finer grained conglomerate with greater amounts of interbedded sandstone characterize the lower and upper parts of the section and mark the transition from the lower shale member and to the Stuver Member, respectively. Paleocurrents are

Grain Size



Grain Size



relatively uniform throughout the section and indicate west-southwesterly sediment transport. The mean and standard deviation of measurements of 15 localities of clast long-axis orientation and imbrication, 2 of tabular cross-strata and 1 of trough cross-strata from the section is $249^{\circ} \pm 25^{\circ}$.

The conglomerate beds typically have erosional bases characterized by coarser conglomerate resting on finer conglomerate, conglomeratic sandstone, sandstone, or shale. The conglomerate beds are most commonly massive and characterized by normal size grading of conglomerate clasts; well-developed imbrication, and clast long-axis orientation. The matrix consists of finer conglomerate and sandstone. Clast-supported conglomerate is most typical, although in finer grained conglomerate and conglomeratic sandstone, matrix-supported conglomerate is common. However, the matrix is never mud-rich and there is no indication of sediment transport and deposition by debris flows or related processes--all coarse-grained deposits appear to have resulted from streamflow processes.

Finer conglomerate typically has parallel stratification, poorly developed and marked by interlayering of coarser and finer conglomerate. Large- and medium-scale trough cross-strata and planar cross-strata are present in the upper parts of some conglomerate beds, but are most characteristic of the conglomeratic sandstone and sandstone units that rest without erosional scour on the lower beds of massive conglomerate.

The fining-upward sequences of conglomerate to sandstone or shale average about 2-3 m in the lower 260 m of the section, about 7 m in the middle 180 m, and about 4 m in the upper 90 m. The sequences record stream deposition, most likely that of a broad braidplain, in which braided streams transporting very coarse debris spread out over a broad aggrading depositional surface. There is no evidence for deposition by meandering streams and no indication for the development of fining-upward point-bar cycles. The lack of significant thicknesses of shale indicates that fine-grained floodplain deposits did not develop. The massive and parallel-stratified bars of conglomerate probably represent longitudinal gravel bars developed on the braidplain. The trough-cross-stratified and planar cross-stratified beds probably represent transverse bars and dunes as well as deposition on the downstream flanks of the longitudinal bars. The thin shale intervals in the upper part of the section probably represent drapes of mud deposited over the bars during post-flooding lowering of stream levels. In most of the sections, these mud drapes, if deposited, were probably eroded away before or during deposition of the overlying bed of conglomerate.

Anaktuvuk Pass area

Middle conglomerate member, Anaktuvuk River

A partial section of the middle conglomerate member of the Kanayut Conglomerate was measured on the east flank of the Anaktuvuk River Valley (NW1/4, T.14S., R.3E., Chandler Lake Quadrangle). The member strikes west-northwest and dips south at a low angle, forming low benches on the side of the valley floor. The section is 162 m thick and is incomplete, with neither a lower contact with the lower shale member or an upper contact with the Stuver Member exposed (fig. 10). The measured section probably forms a portion of the middle and upper parts of the middle conglomerate member.

The section consists predominantly of gray to buff-colored conglomerate and conglomeratic medium- to very coarse-grained sandstone. Fine-grained sandstone is uncommon, siltstone rare, and shale absent. The conglomerate forms lenticular beds that are chiefly clast-supported and as thick as 4.3 m. Clasts are moderately well sorted, rounded, and as large as 23 cm in maximum dimension.

Conglomerate forms the lower part of couplets that are characterized by an upward decrease in grain size and bed thickness. The maximum couplet thickness is 6.7 m. The upper portions of couplets consist of massive, low-angle, and planar or trough-cross-stratified sandstone. Noncyclic amalgamated beds of sandstone or conglomerate and rare coarsening-upward sandstone bodies are also present in the section.

The section was probably deposited by low sinuosity braided rivers. The conglomerate and sandstone couplets probably resulted from accretionary processes related to the growth and migration of longitudinal and transverse bars. Cross-stratified sandstone probably formed by migration of dunes on the tops of the larger bars. Maximum couplet size may provide a rough indicator of paleoflow depth during bankfull discharge. Paleosol development indicates periods of bar emergence.

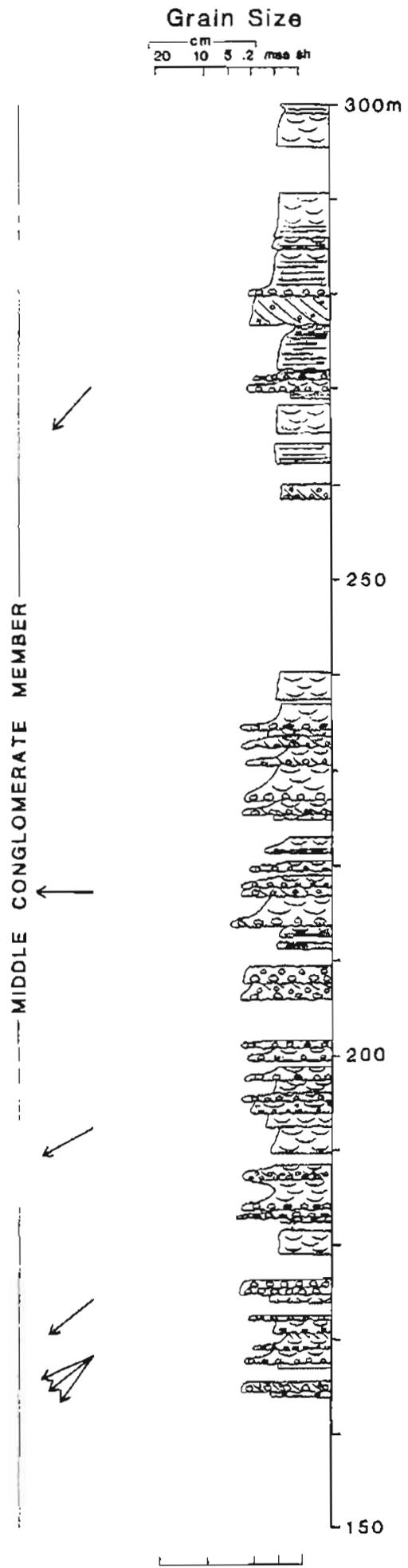
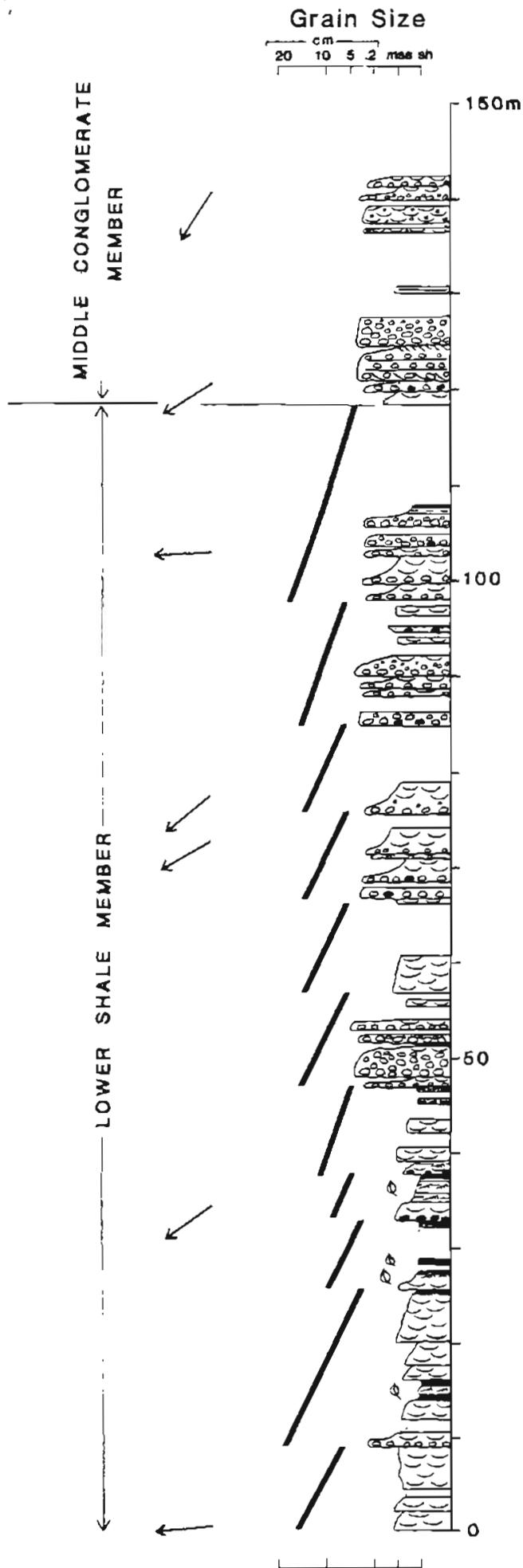
Twenty-one paleocurrent measurements were obtained from the section (eight are not shown in Fig. 10, because their exact stratigraphic position is not known) including 9 of clast imbrication and long-axis orientation, 6 of tabular cross-strata, 5 of trough cross-strata and 1 flute cast. The mean and standard deviation of all measurements are $233 \pm 50^\circ$.

Kanayut Conglomerate, John River

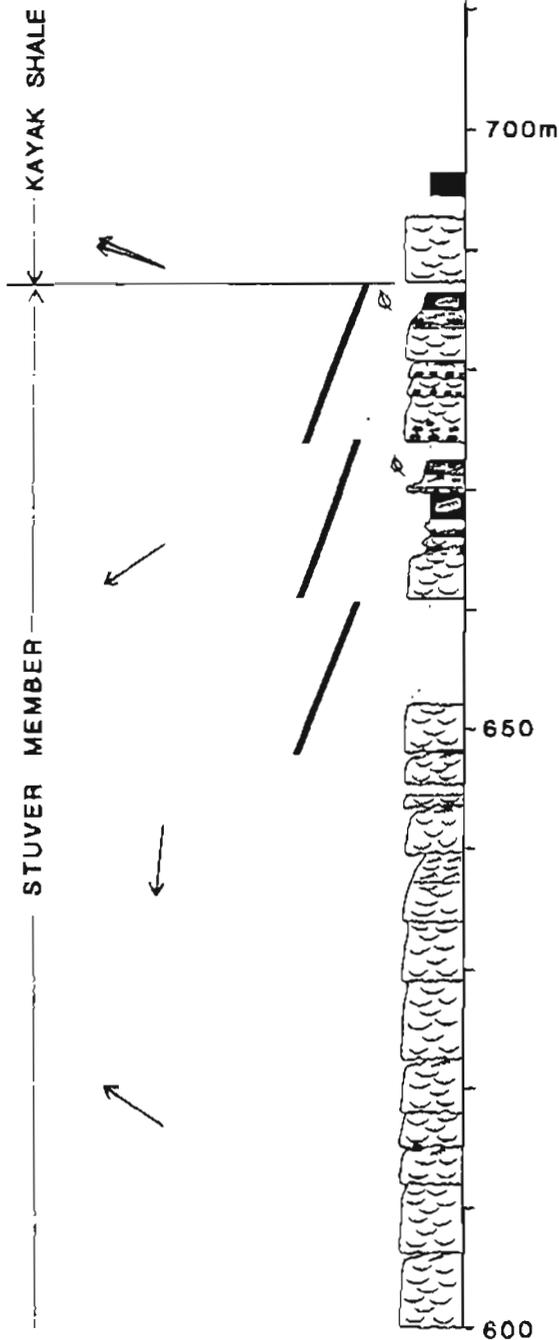
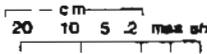
An almost complete section of gently dipping Kanayut Conglomerate was measured along a ridge extending eastward from Ekokpuk Mountain to the floor of the John River Valley (SW1/4, T.37N., R.20W., Wiseman Quadrangle). This section is the southernmost one measured and totals 697 m in thickness. Toward the south and west, the Kanayut Conglomerate in general becomes finer grained and the middle conglomerate member becomes more sand-rich and depleted in conglomerate. As a result, the members of the Kanayut become more poorly defined in this area. In this section the upper and lower parts contain abundant shaly intervals which are commonly expressed as covered intervals. We therefore tentatively assign the lower 119 m to the lower shale member, the overlying 403 m to the middle conglomerate member, the next 175 to the Stuver Member, and the upper 10 m to the basal sandstone member of the Kayak Shale (fig. 11).

The underlying Hunt Fork Shale crops out in the valley at an elevation of 2150 ft (715 m) and the first outcrops of the Kanayut are found at 2600 ft (865 m) on the ridge. The interval between the two outcrops is covered, but topographic expression suggests that the contact between the units is just below the first Kanayut exposures. There is, however, a possibility that as much as 140 m of the lower shale member of the Kanayut lie below the inferred contact and were not measured. Neither we nor Brosge and others (1979b) recognized the basal sandstone in this section.

The lower shale member consists of eleven fining- and thinning-upward cycles interpreted as the lateral and vertical accretion deposits of



Grain Size



Thirty paleocurrent measurements, consisting of 28 trough cross-strata and 2 tabular cross-strata were measured and indicate sediment transport toward the southwest. The mean and standard deviation of 15 measurements from the lower shale member are $239^{\circ} \pm 18^{\circ}$ whereas the mean and standard deviation of 8 and 7 measurements from the middle conglomerate and Stuver members are more westerly -- $271^{\circ} \pm 15^{\circ}$ and $265^{\circ} \pm 48^{\circ}$, respectively. The mean and standard deviation of all 30 measurements from the section are $253 \pm 30^{\circ}$.

Mount MacVicar area

Lower shale member

A complete section of the lower shale member of the Kanayut Conglomerate was measured along a west-trending ridge on the southeast side of Mount MacVicar (T.15S., R.3W., Chandler Lake Quadrangle). The lower shale member is 260 m thick, conformably overlies marine strata of the basal sandstone member of the Kanayut Conglomerate, and is overlain by the middle conglomerate member of the Kanayut Conglomerate (fig. 12). There is no apparent interfingering between the lower shale member and the basal sandstone member.

The section consists of 18 fining-upward cycles inferred to have been deposited by meandering rivers. The cycles range from 12 to 29 m in thickness. The coarsest cycles typically contain massive or crudely stratified conglomerate at the base, overlain by trough- cross-stratified fine- to coarse-grained sandstone, ripple-marked fine- to very fine-grained sandstone, and laminated very fine-grained sandstone, siltstone, and shale. Coarser grained lower parts of cycles contain multiple fining-upward sequences separated by erosional surfaces. The coarse-grained facies are interpreted to be river channel and lower point bar deposits and fine-grained facies to be upper point bar, levee, and floodplain deposits. Interspersed within fine-grained intervals are some massive or thickly bedded bodies of fine-grained sandstone interpreted to be orevasse-splay deposits.

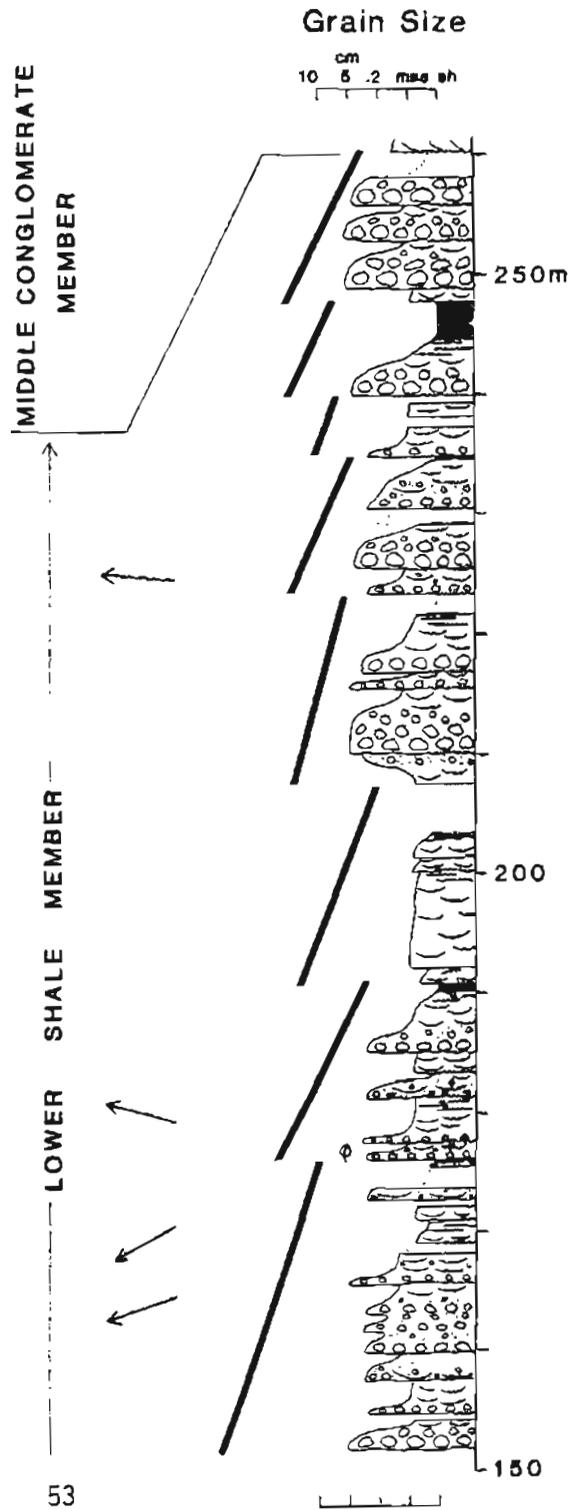
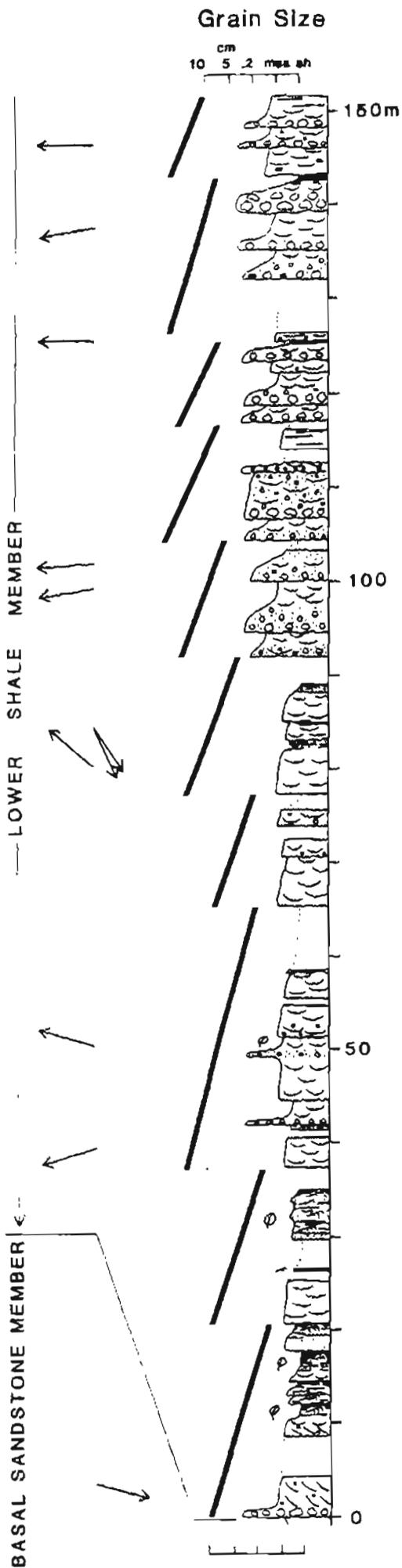
The beds of conglomerate are as thick as 3.25 m and the maximum clast size is 5 cm. The amount of conglomerate increases upsection toward the contact with the middle conglomerate member. The two lowest cycles contain the greatest thickness of floodplain deposits and contain abundant fragments of plant fossils.

Paleocurrent measurements are fairly consistent throughout the section and indicate westward sediment transport. Two measurements at about 85 m above the base of the section indicate southeastward sediment transport. The mean and standard deviation for 16 measurements of trough cross-stratification and 1 of tabular cross-stratification in the section are $261^{\circ} \pm 57^{\circ}$.

Okokmilaga Valley area

Lower shale member

A very thick and complex fining-upward cycle in the lower shale member of the Kanayut Conglomerate was measured along a ridge flank north of Kikoyat Creek on the west side of the Okokmilaga River (sec. 19, T.34N., R.5W., Killik River Quadrangle). The fining-upward cycle is 37 m thick and consists of alternating beds of conglomerate and sandstone, bounded above and below by shale (fig. 13). The thick cycle is important because (1) it has been mapped



extensively between the Killik and Okokmilaga Rivers in the central Brooks Range as a marker unit (unit Dklm of Brosge and others, 1979b), (2) it has been measured in several other sections of the lower shale member, and (3) it appears to form a correlatable unit, and (4) it forms a partly marine cycle. Scolithus-type burrows have been observed in the area where the present section was measured and in other areas where the thick cycle has been observed.

In the measured section along Kikoyat Creek, a series of fining-upward subcycles of conglomerate and cross-stratified sandstone comprise the larger cycle. The subcycles are bounded by thick sandstone beds that are present at the following intervals above the base of the cycle: 7 m, 13 m, 20 m, 27 m, and 35 m. These sandstone units are as thick as 3 m and contain trough cross-strata. The conglomeratic beds are massive, parallel-stratified, or cross-stratified, and do not generally form well-defined fining-upward components of the subcycles. The maximum clast size measured was 3 cm and clasts of this size are found throughout the section, suggesting that the section does not form an overall fining-upward cycle. Although no marine trace fossils were found in the measured section, Scolithus-type trace fossils are present in the cycle several tens of meters to the east and west of the line of section.

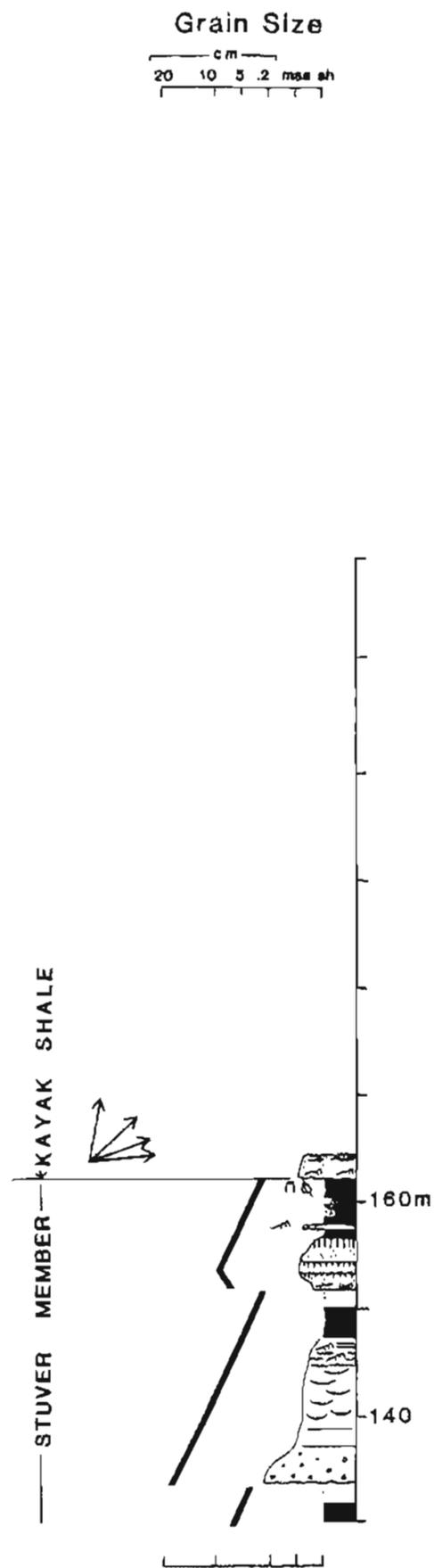
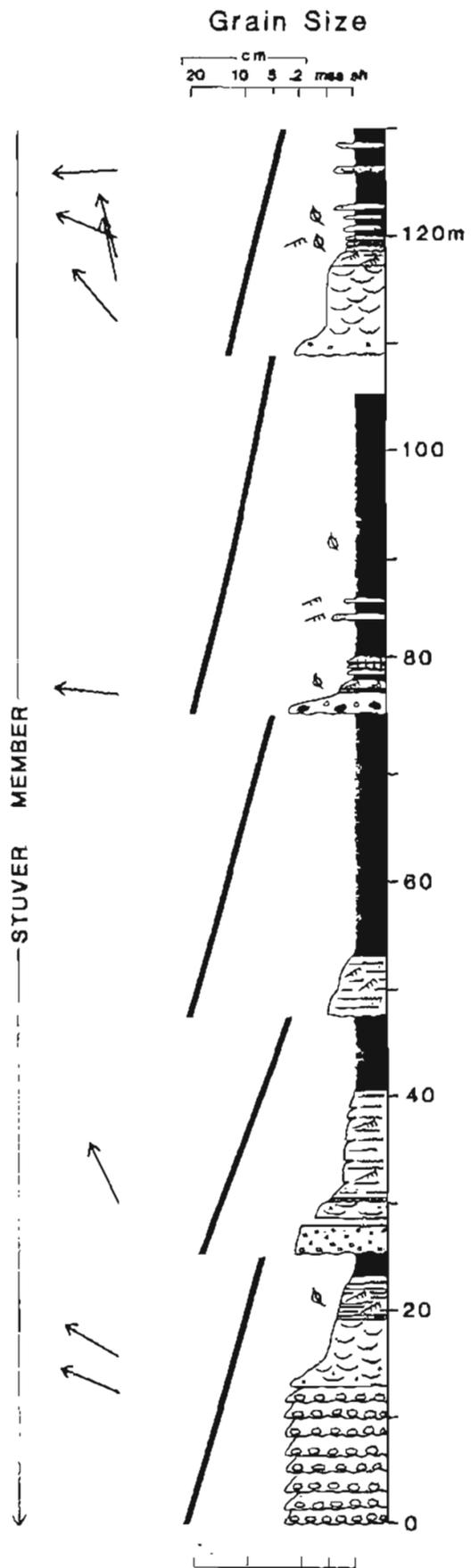
Paleocurrents from the thick cycle indicate variable transport toward the west and south. We measured 16 paleocurrent indicators (5 are not shown in fig. 13 because their exact stratigraphic position is not known), including planar and tabular cross-strata, clast imbrication and long-axis orientations, and primary current lineation. The mean and standard deviation of the measurements are $235^{\circ} \pm 65^{\circ}$. The variable nature of the paleocurrents may reflect intermixing of the fluvial transport and shallow-marine transport in the cycle.

We believe that the thick mappable cycle in the lower part of the lower shale member represents a marine incursion during deposition of the dominantly fluvial lower shale member. The unit may be recognizable as far east as the Atigun River (fig. 3). Parts of the section that contain well-sorted sandstone, laterally extensive single-pebble-thick layers of conglomerate, and variably oriented medium-scale cross-strata also argue for a marine rather than fluvial origin. In the lower part of the measured section, about 6 m above the base, is a prominent recumbently folded bed of medium- to coarse-grained sandstone with fluid-escape structures. The folds suggest slumping in a direction of 278° . Small-scale syndepositional folds of this nature have not been observed in fluvial parts of the Kanayut sequence and may also argue for a marine origin.

Stuver Member

A well-exposed and complete section of the Stuver Member of the Kanayut Conglomerate was measured along a ridge crest west of the Okokmilaga River (sec. 8, T.33N., R25E., Killik River Quadrangle). The west-dipping member crops out in a saddle west of a prominent ridge formed by the top of the middle conglomerate member at the base of the section. Individual resistant beds of the Stuver Member form topographic ridges within the section.

Approximately 162 m of section were measured, with covered intervals totalling only 5 percent (fig. 14). The well-defined boundary with the



Creeks (SE 1/4, T.33.N, R.20E., Killik River Quadrangle). The section crops out on the southern flank of an east-trending anticline and probably includes all of the lower shale member and most of the middle conglomerate member (fig. 15). The upper part of the section is truncated by the Toyuk Thrust and is poorly exposed. As a result, all of the Stuver Member and at least part of the middle conglomerate member are probably cut out. The lower 240 m of the measured section is assigned to the lower shale member and the upper 530 m to the middle conglomerate member. The middle conglomerate member in the Killik River area is finer grained and richer in sandstone compared to areas farther north and east making it less easily distinguished from the underlying and overlying units. We placed the boundary between the lower shale member and middle conglomerate member at the top of the stratigraphically highest shale unit thicker than 2 m.

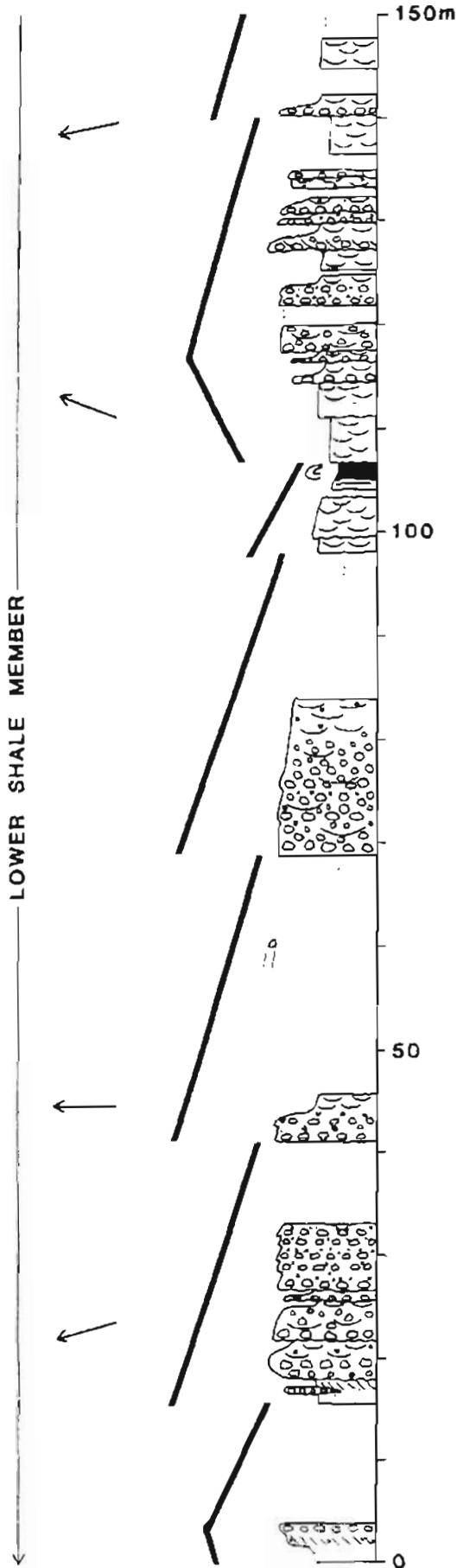
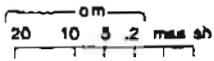
The lower 240 m of the section consists of interbedded conglomerate and sandstone with partly covered intervening intervals of fine-grained sandstone, siltstone, and shale. Strata below the base of the measured section are not exposed, but topographic expression and projection of the base of the section along strike to the northwest strongly suggests that the lowest outcrop marks the contact between the Hunt Fork Shale and the lower shale member of the Kanayut Conglomerate. The basal sandstone member of the Kanayut Conglomerate is not present. The basal 4 m of the section consists of a coarsening-upward cycle of cross-stratified medium-grained sandstone to pebble conglomerate. This cycle differs from those higher in the section and we interpret it to be the top of the marine Hunt Fork Shale.

The overlying lower shale member of the Kanayut consists of 12 fining-upward cycles which are about 25-30 m thick near the base and 15-20 m thick near the top. The lower parts of the cycles consist of massive, crudely parallel-stratified or cross-stratified pebble conglomerate and conglomeratic sandstone that is channeled into the underlying cycles. The coarse-grained units contain clasts as large as 5 cm and have numerous thinner fining-upward sequences that are separated by erosional surfaces. The coarse-grained basal parts of the cycles grade upward into trough cross-stratified, parallel-stratified, or massive medium- to coarse-grained sandstone. The upper parts of the cycles are mostly poorly exposed, but some outcrop and float show that they consist of brown to black, ripple-marked and laminated fine-grained sandstone, siltstone and shale. The fine-grained upper parts of cycles decrease in thickness upward from 10-20 m near the base to 5-10 m near the top of the lower shale member.

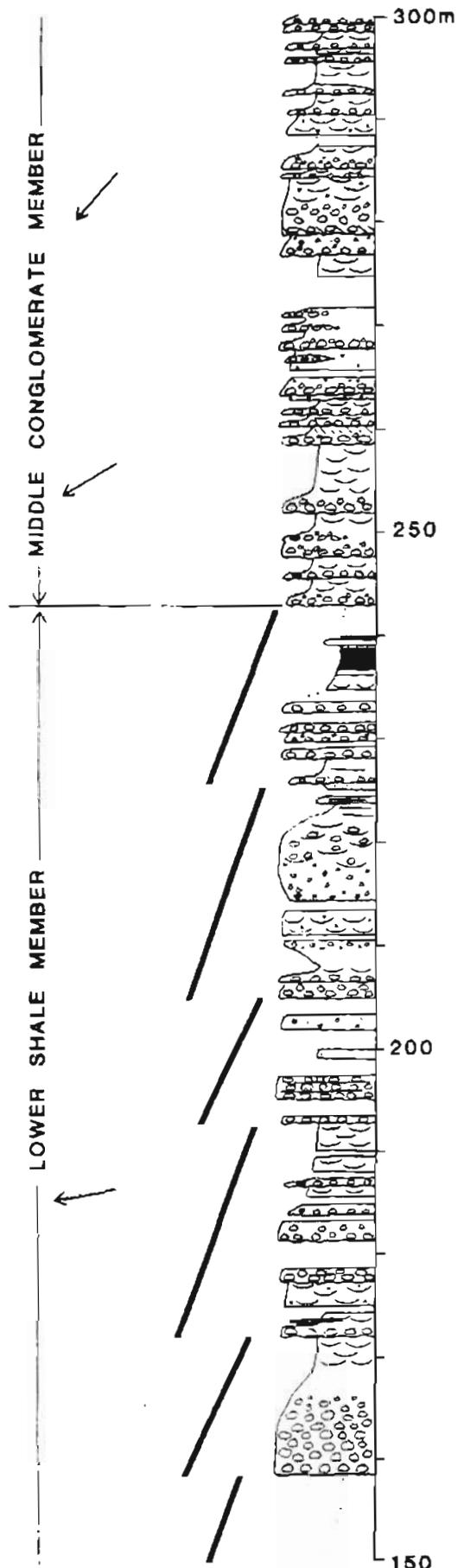
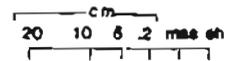
The fining-upward cycles of the lower shale member are interpreted to represent the lateral and vertical accretion deposits of meandering rivers. The coarse-grained basal parts of the cycles are inferred to be channel and lower point-bar deposits and the fine-grained upper parts to be upper point-bar, levee and floodplain deposits. *Scolithus*-type burrows in float 50 m above the base, however, may indicate that the fine-grained parts of some cycles contain marine deposits. A marine origin for some strata may also be indicated by small-scale recumbent folds which occur in very fine-grained sandstone 105 m above the base of the section.

Conglomerate-sandstone couplets predominate in the lower part of the middle conglomerate member. The couplets consist of massive to tabular cross-stratified beds of conglomerate that contain clasts as large as 4 cm and were

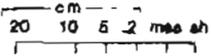
Grain Size



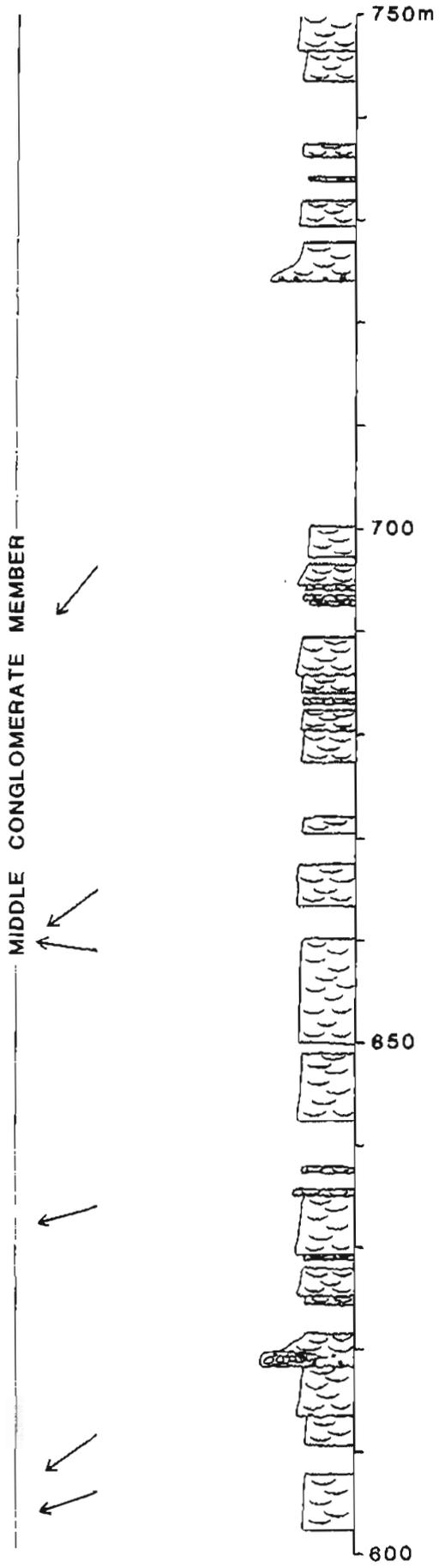
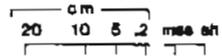
Grain Size



Grain Size



Grain Size



63

MIDDLE CONGLOMERATE MEMBER



were able to more easily sample large thickness of stratigraphic section for the largest clast contained in their conglomeratic strata. However, the map distribution of maximum clast sizes must still be interpreted cautiously with regard to paleogeography because of the presence of significant amounts of structural shortening, including at least three major thrust faults within the outcrop belt. Nevertheless, several major conclusions can be drawn from the available data.

We have compiled a map of maximum clast size data from the Kanayut Conglomerate collected during the 1978, 1979 and the 1980 field seasons (fig. 16). The 1980 data were collected primarily between the Killik River and the Sagavanirktok River. The middle conglomerate member generally contains the largest clasts and we collected most data from it.

The largest clasts we observed in the Kanayut Conglomerate are 23 cm in length and are located near Shainin Lake. To the west and south of Shainin Lake, the clast size decreases regularly and dramatically. Conglomerate is rare in the Kanayut north of the Noatak River, demonstrating a marked westward decrease in clast size away from Shainin Lake. Clast size also decreases southeast of Shainin Lake, as far east as the longitude of the Canning River, where the largest clasts are 5 cm in size. The distribution of decreasing clast sizes around the Shainin Lake region suggests that it may mark an entry site of a major trunk system into the Kanayut depositional basin.

A second major trunk system may be indicated by the clast size data in the northern and eastern outcrops of the Kanayut. Clast sizes in that region decrease from a maximum of 10 cm in the northeasternmost outcrops to less than 5 cm in the Arctic Village area. This distribution of clast sizes suggests that sediment dispersal in the eastern part of the Kanayut was from northeast to southwest.

The pattern of sediment dispersal is apparent from a contour map of clast sizes (fig. 17). This map is drawn from the information presented in figure 16 and shows the well-defined maximums of clast sizes in the Shainin Lake area and in the northeasternmost outcrops. The geometry of the contour lines around these clast size maximums indicates that sediment transport in both systems was primarily toward the southwest.

A third maximum appears to be present along the northern side of the outcrop belt near the Sagavanirktok River. It may represent a third site of entry of a major river system into the Kanayut depositional system. The minimal bowing of the contour lines south of this third clast-size maximum may indicate that it represents a smaller distributary system than the Shainin Lake and northeastern systems. In addition, the orientation of the contour lines suggests that sediment transport was mostly southerly rather than southwesterly in this possible distributary system.

CONGLOMERATE CLAST COMPOSITION

Introduction

Nilsen and others (1980a, 1980b) determined the clast composition of 79 beds of conglomerate from the Kanayut Conglomerate and associated units from

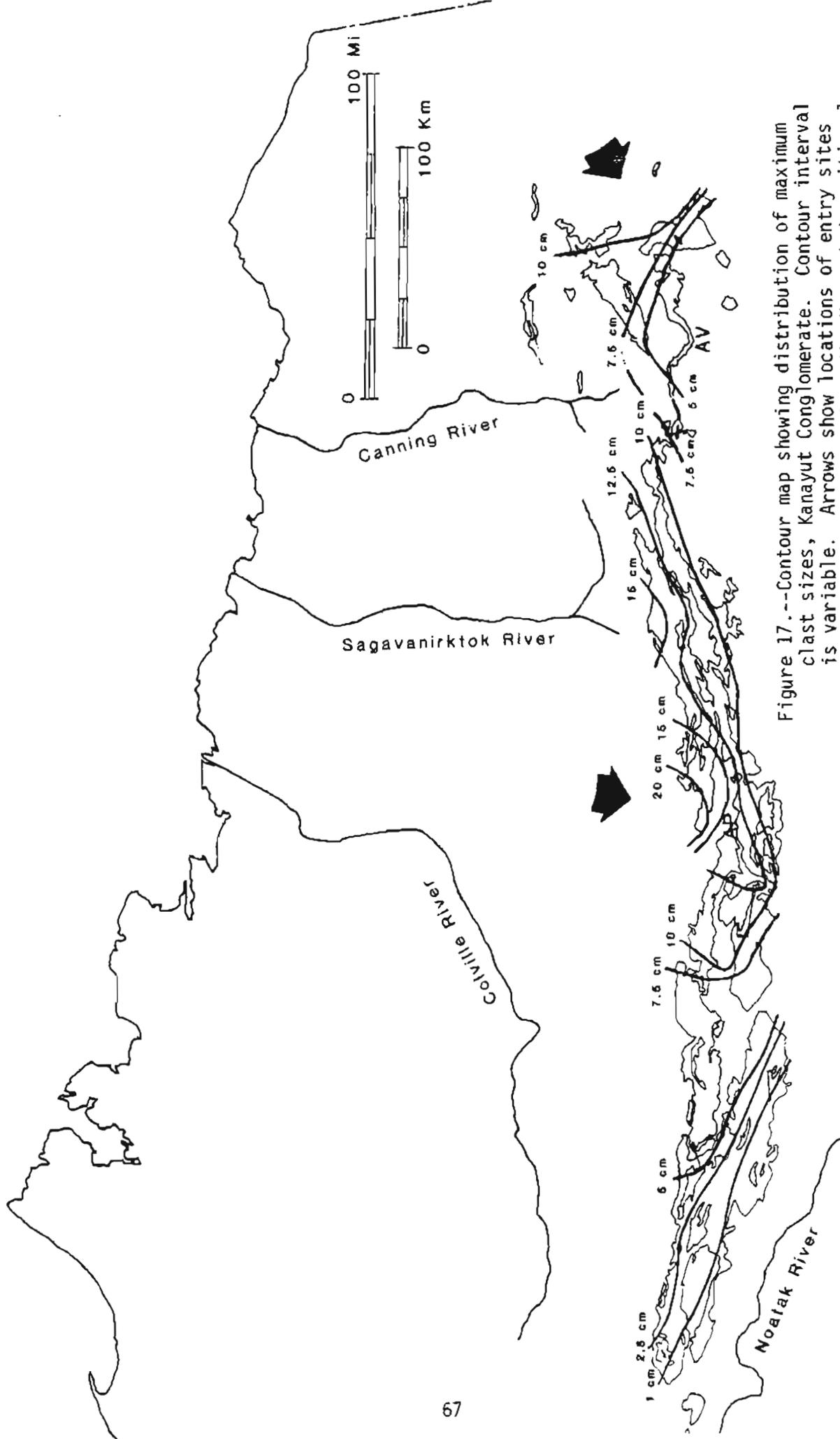


Figure 17.--Contour map showing distribution of maximum clast sizes, Kanayut Conglomerate. Contour interval is variable. Arrows show locations of entry sites of major trunk systems into the Kanayut depositional basin. Abbreviations: AV, Arctic Village; AP, Anaktuvuk Pass.

all but two of the conglomerates sampled contain more than 80 percent chert, vein quartz, quartzite, and quartzite conglomerate, and only five samples contain less than 90 percent of these constituents. The second ternary diagram has vein quartz, quartzite and chert as poles and illustrates the predominance of chert in most of the pebble counts from the Kanayut Conglomerate (fig. 19). Several pebble counts plot at the middle part of the vein quartz-chert join away from the main cluster of data. We cannot explain these variations, which appear to be random.

We have plotted clast size versus percent of each clast lithology for the combined 1978, 1979, and 1980 pebble count data from the Kanayut Conglomerate (fig. 20). This diagram again shows that the Kanayut is rich in chert and typically contains a moderate amount of vein quartz and quartzite. It also shows that there is a consistent relationship between percent of quartzite clasts and clast size (fig. 20C). Coarser parts of the Kanayut Conglomerate are enriched in quartzite. Our selective pebble count of only those clasts larger than 7 cm in longest dimension contains 67 percent quartzite plus quartzite conglomerate, strongly suggesting that quartzite and quartzite conglomerate clasts are more abundant in the coarser fractions of the conglomerate. The enrichment of quartz in this pebble count is at the expense of chert (18 percent), suggests that there is less chert in coarser conglomerate. However, this relationship is not seen in the clast size versus percent of chert diagram (fig. 20B), perhaps suggesting that provenance plays a dominant role in determining the abundance of chert in finer conglomerates.

The 1980 pebble-count data from the Kanayut Conglomerate plot within the fields for the Kanayut shown in the various diagrams previously reported by Nilsen and others (1981a) and are comparable with the earlier data in every respect. We suggest that this generally consistent conglomerate composition throughout the very extensive Kanayut depositional system indicates that the Kanayut detritus was probably derived from a single major source terrane.

Kayak Shale

Rare conglomerate in the Kayak Shale is typically associated with shallow-marine sandstone bodies in its basal sandstone. Two pebble counts from conglomerate of the basal sandstone member of the Kayak Shale were made at one location south of Galbraith Lake at the top of measured section of the Stuver Member (fig. 6) during the 1980 field season (Table 1). These counts suggest that the conglomerate of the basal sandstone member of the Kayak Shale, like the Kanayut Conglomerate, is dominated by chert clasts (83 percent), has subordinate amounts of vein quartz (average 16 percent), and locally contains minor amounts of quartzite clasts (average 1 percent). The basal sandstone member of the Kayak Shale is interpreted to reflect nearshore deposition in the regressing part of the Kanayut fluvio-deltaic complex (Nilsen and others, 1980a, 1981a). The two Kayak Shale pebble counts are consistent with this interpretation, exhibiting compositional similarity to and plotting well within the fields of the Kanayut data (figs. 18, 19, and 20).

CARBONATE CLASTS

**A. Basal Sandstone
Member, Kayak Shale**

**B. Kanayut
Conglomerate**

**IMMATURE
(OTHER LITHIC) CLASTS**

**MATURE
(QUARTZOSE) CLASTS**

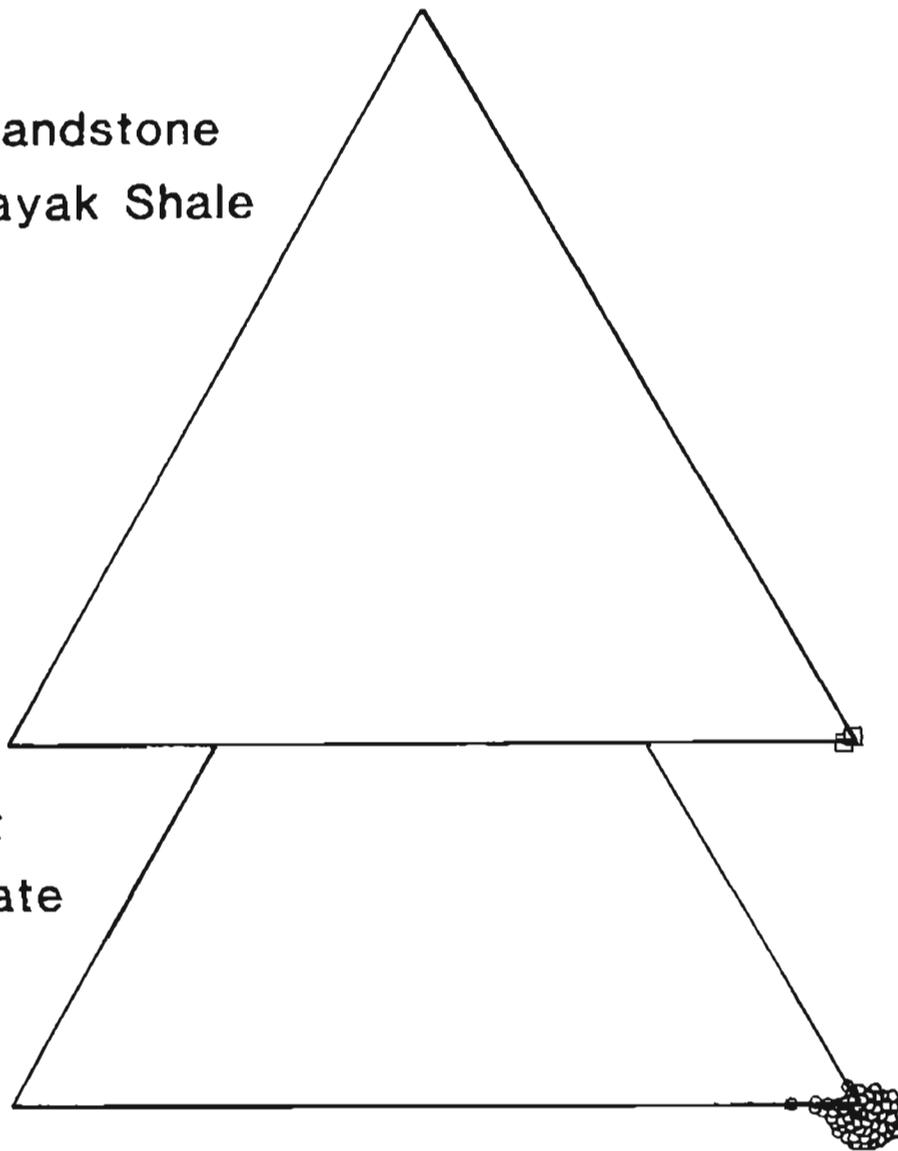


Figure 18.--Immature clast (argillite, siltstone, sandstone, and greenstone)--carbonate clast--mature clast (chert, vein quartz, and quartzite) ternary diagram from pebble counts made during the 1978, 1979, and 1980 field seasons. A, Basal sandstone member of the Kayak Shale. B, Kanayut Conglomerate.

Figure 20.--Plots of maximum clast size versus percentage of clast lithology for the Kanayut Conglomerate from pebble counts made during the 1978, 1979, and 1980 field seasons. A, Vein quartz. B, Chert. C, Quartzite. Open circles represent pebble counts with a minimum clast size larger than 1 cm; filled circles represent those with a minimum clast size larger than 7 cm.

PALEOCURRENTS

Introduction

Nilsen and others (1980) plotted 166 paleocurrent measurements made in 1978 from the Kanayut Conglomerate, Kekiktuk Conglomerate and Kayak Shale. During the 1979 field season, an additional 292 paleocurrent measurements were collected from these units and the Hunt Fork Shale and Beaucoup Formation at 63 separate locations (Nilsen and others, 1981). Data from both years were compiled and presented in map form (Nilsen and others, 1980b). An additional 316 measurements were collected during the 1980 field season from the Hunt Fork Shale, Kanayut Conglomerate, and Kayak Shale, bringing the total to 774 determinations for all three seasons (figs. 21 and 22).

Sedimentary features measured include medium-scale tabular and trough cross-strata (447 measurements), conglomerate imbrication and long-axis orientation (188 measurements), primary current lineation (83 measurements), ripple markings (33 measurements), flute marks (9 measurements), pebble trains and erosional scours (4 measurements of each), channel margins, fluid-escape folds, and ball and pillow structure (1 measurement of each). Restorations of paleocurrent directions to the horizontal were done manually on a stereonet. Computer-calculated vector means and standard deviations were determined for 37 locations at which more than four paleocurrent measurements were made and also for the total number of paleocurrent measurements from each of the various stratigraphic units. Bidirectional features such as primary current lineation were assigned a westerly or southerly sense because of the preponderance of unidirectional indicators with that orientation.

Hunt Fork Shale

Ten paleocurrent determinations were obtained from the Hunt Fork Shale at 5 separate locations during the 1979 field season (Nilsen and others, sheet 2). An additional 21 measurements collected during the 1980 field season from 7 separate locations brings the total to 31 determinations, including 17 of trough cross-strata, 8 of primary current lineations, 3 of tabular cross-strata, 2 of oscillation ripple markings, and 1 flute mark. The azimuthal vector mean and standard deviation of all measurements from the Hunt Fork Shale are $207^{\circ} \pm 54^{\circ}$, reflecting the overall southerly sediment transport direction (fig. 22). The relative consistency of these data probably indicates that sediment was chiefly transported offshore. A few measurements which show significant deviation from the mean may result from storm-generated currents affecting sediments deposited in shallower water.

Fluvial members of Kanayut Conglomerate

In our report of the 1978 field season, we plotted 158 paleocurrent measurements from the fluvial lower shale, middle conglomerate and Stuver members of the Kanayut Conglomerate (Nilsen and others, 1980a). During the 1979 field season, we measured an additional 156 paleocurrent directions at 35 separate locations. The orientation of the measurements was compiled and plotted in map form and by summary rose diagram (Nilsen and others, 1980b), sheet 1). An additional 291 determinations were collected during the 1980 field season, bringing the total number of measurements to 605 which include 249 trough cross-strata, 170 clast imbrication and long-axis orientations, 88 tabular cross-strata, 60 primary current lineations, 16 current ripple

markings, 8 flute marks, 4 pebble trains, 4 erosional scours, 2 aligned plant fragments, 2 oscillation ripple markings, 1 channel-margin orientation, and 1 fluid escape fold (fig. 21).

Unidirectional indicators show a southwest transport of sediment, and bidirectional indicators (those giving sense of transport only) show west-southwest or east-northeast transport. The azimuthal vector mean and standard deviation of all measurements from the Kanayut is $243^{\circ} \pm 45^{\circ}$. Although the local orientation of paleocurrents is as southerly as $172^{\circ} \pm 28^{\circ}$ for locations having more than four measurements, the overall southwesterly transport direction is consistent between members as shown by a calculation of $243^{\circ} \pm 47^{\circ}$ for the lower shale member (209 measurements), $245^{\circ} \pm 40^{\circ}$ for the middle conglomerate member (247 measurements), and $234^{\circ} \pm 61^{\circ}$ for the Stuver Member (128 measurements), with 31 paleocurrent determinations unassigned to stratigraphic units. These data indicate that the fluvial sediment transport direction was predominantly toward the southwest throughout the entire time of deposition of the Kanayut Conglomerate. Donovan and Tailleux (1975) previously determined southerly directions of transport for the Kanayut.

Kayak Shale

Three paleocurrent measurements from current ripple markings in the shallow-marine or intertidal basal sandstone member of the Kayak Shale were previously reported (Nilsen and others, 1980a). An additional 10 measurements were collected during the 1979 field season (Nilsen and others, 1980b, sheet 2) and 6 more during the 1980 field season. The total number of measurements include 8 current ripple markings, 5 oscillation ripple markings, 4 tabular cross-strata, 1 long-axis orientation, and 1 ball and pillow structure.

Although the sediment transport direction given by the measurements is relatively consistent at each location, there is considerable variability in data between locations (fig. 21). Because of this geographic variation, the significance of the vector mean and standard deviation ($120^{\circ} \pm 79^{\circ}$) is relatively minor. We attribute the variability of paleocurrent measurements in the Kayak Shale to alternating offshore, onshore, and longshore sediment transport related to wave-, wind-, storm-, and tide-generated currents.

SUMMARY

This report summarizes stratigraphic and sedimentologic data collected during the 1980 field season in the central Brooks Range from the Upper Devonian and Lower Mississippian Endicott Group. The Upper Devonian Kanayut Conglomerate consists of a basal marine sandstone member and three successive fluvial members. It crops out for about 1000 km along strike and about 50 km across strike in at least three major thrust plates in the central part of the Brooks Range. The basal marine sandstone member was probably deposited as channel-mouth bars, offshore bars, and spits. The overlying lower shale member is inferred to have been deposited by meandering streams over a large floodplain area. The middle conglomerate member, which contains the coarsest conglomerate and little or no shale, was probably deposited by braided streams. The Stuver Member is inferred to have been deposited by meandering streams and is overlain by intertidal and shallow-marine sandstone at the base of the Kayak Shale.

The maximum clast size of conglomerate decreases westward, southward, and eastward from the Shainin Lake area, suggesting that a major trunk stream originally entered the depositional basin in this area. A second major trunk stream probably entered the depositional basin at its northeastern end.

The composition of the Kanayut Conglomerate clasts varies little from place to place or member to member. In most of the conglomerates examined, about 80-95 percent of the pebbles are chert, about 5-15 percent quartz, and 1-5 percent quartzite. Red chert clasts are locally abundant in the Shainin Lake and Galbraith Lake area. Pebbles of argillite and other rock fragments are rare.

The orientations of cross-strata, primary current lineations, current ripple marks, and imbrication and long axes of pebbles in the three fluvial members of the Kanayut consistently show sediment transport toward the southwest across most of the central and eastern Brooks Range. Paleocurrents reported from the 1980 field season indicate more southerly transport in the Atigun River area.

The facies sequence in the Hunt Fork Shale and Kanayut Conglomerate suggests that the Kanayut comprises the marginal-marine and fluvial parts of a prograding delta system. The consistent southwestward direction of paleocurrents in the fluvial deposits, together with the southwestward decrease of grain size, suggests an eastern, northern, or northeastern source, although the allochthonous nature of the outcrop belt precludes identification of the source at present. Judging from the abundance of chert, quartz, and quartzite clasts in the Kanayut, the source terrane was probably composed mostly of slightly metamorphosed sedimentary rocks. The high roundness and sphericity of the clasts and their compositional maturity indicate that the source terrane may also have included older conglomerates.

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